

Transmissivity Averaging in Fracture Flow on Self-affine Linear Profiles: Arithmetic, Harmonic, and Beyond

Alexandre Lavrov¹

Received: 10 January 2023 / Accepted: 31 August 2023 © The Author(s) 2023

Abstract

Numerical modeling of unidirectional flow in self-affine fractures using the lubrication approximation requires averaging of the transmissivity between the nodes. Seven averaging techniques are reviewed: arithmetic averaging of transmissivity; harmonic averaging of transmissivity; two averaging techniques derived by cell-based collocation method; global reconstruction of profile by means of multiquadrics; arithmetic averaging of aperture; harmonic averaging of aperture. In order to evaluate the performance of the seven techniques in terms of pressure errors and hydraulic aperture errors, self-affine profiles of 1024 nodes with different Hurst exponents (0.4 to 0.8) are generated. Every second node is then removed, resulting in 512-node profiles. Apertures at removed nodes are used in reference flow simulations on the 512-node profiles. Then, simulations with the seven averaging techniques are performed on 512-node profiles. Errors are computed with regard to the results obtained in the reference simulations. Reconstruction with multiquadrics is found to provide superior accuracy on self-affine profiles, followed by harmonic averaging of transmissivity or harmonic averaging of the aperture (Some of the errors analyzed in this study are minimized with the two last mentioned schemes.). Multiquadrics reconstruction is found to provide the best accuracy also on a smooth periodic profile.

Keywords Flow \cdot Self-affine linear profile \cdot Finite-difference method \cdot Transmissivity \cdot Averaging

1 Introduction

Fluid flow in the subsurface is often controlled by natural fractures. In order to predict flow and contaminant transport in such systems, hydraulic properties of individual fractures must be available. Rock fractures usually have rough walls which leads to a variable local fracture aperture, *w*, defined as distance between the walls. There is strong evidence that aperture distributions of fractures in different materials, including rocks, are self-affine, with the fractal dimension in the range of 2.2–2.8 (Bouchaud et al. 1990; Måløy et al. 1992; Mandelbrot et al. 1984; Plouraboué et al. 1995; Schmittbuhl and Måløy 1997). Hydraulic

Alexandre Lavrov alexandre.lavrov@ntnu.no

¹ Norwegian University of Science and Technology, Trondheim, Norway

properties of such fractures are described by hydraulic aperture. This is the aperture of a smooth-walled conduit that yields the same flow rate as the real, rough-walled fracture does, given the same pressure difference. Substantial effort has been invested over the past 3–4 decades into the study of hydraulic aperture of rough-walled fractures. A major part of these studies is numerical whereby a constant pressure difference is applied between two edges of a rectangular fracture or between two ends of a linear, one-dimensional fracture, and the resulting flow rate is computed by finite-difference or finite-element method. In the case of linear, one-dimensional fractures and under the assumptions of the lubrication theory approximation (Zimmerman et al. 1991), the numerical problem reduces to solving

$$\frac{d}{dx}\left(c\frac{dP}{dx}\right) = 0, \ 0 < x < L$$

$$P|_{x=0} = P_0, \ P|_{x=L} = 0$$
(1)

where *L* is the length of the fracture; *P* is the fluid pressure; c = c(x) is the local fracture transmissivity that can be defined as w^3 . Second-order finite-difference discretization of Eq. (1) on a uniform grid typically used in these models yields:

$$c_{i+\frac{1}{2}}(P_{i+1} - P_i) - c_{i-\frac{1}{2}}(P_i - P_{i-1}) = 0$$
⁽²⁾

where *i* is the node number. Both pressure and transmissivity are specified at nodes, which necessitates approximation, or *averaging*, of transmissivity, *c*, between the nodes, i.e., at locations $i \pm 1/2$. Two types of transmissivity averaging are often used in numerical models of subsurface flow: arithmetic and harmonic, given, respectively, by

$$c_{i\pm\frac{1}{2}} = \frac{1}{2} \left(c_i + c_{i\pm 1} \right) \tag{3}$$

and

$$c_{i\pm\frac{1}{2}} = \frac{2c_i c_{i\pm1}}{c_i + c_{i\pm1}} \tag{4}$$

Harmonic averaging is the method of choice in some hydrogeological simulators, e.g., MODFLOW (Langevin et al. 2017). It is often used in fracture flow models (Tsang and Neretnieks 1998) and porous media flow models (Aavatsmark 2002; Bessone et al. 2022; Roth et al. 1996; Colecchio et al. 2020). Its use for diffusion equation in strongly heterogeneous media was advocated by van Es et al. (2014). Its use in nonlinear diffusion problems was disputed by Kadioglu et al. (2008) who showed that it yields large errors when the transmissivity is P^3 . In fracture flow, harmonic averaging typically results in lower values of the hydraulic aperture than the arithmetic averaging because arithmetic average of two real numbers is always greater than or equal to their harmonic average. It was obtained by Lavrov (2022) that, in two-dimensional fracture flows, numerical schemes with arithmetic averaging perform better in terms of mass balance error.

Instead of averaging the transmissivity, some schemes make use of averaging the fracture aperture. Here, several options are available, too. The two main choices are arithmetic averaging and harmonic averaging. Arithmetic averaging of the local aperture was used, e.g., by Lenci et al. (2022a). It was pointed out by Lenci et al. (2022b) that harmonic averaging of aperture preserves the energy in the scheme.

The importance of averaging techniques in flux evaluation is not limited to fluid flow computations. Accurate evaluation of fluxes is crucial also for numerical heat transfer (Klepikova et al. 2021), solute transport (Moreno et al. 1988) and particulates transport (Lin et al. 2022; Mao et al. 2023) in fractures.

The objective of this work was to assess the performance of several averaging techniques in fracture flow on self-affine linear profiles.

2 Averaging Techniques

Seven averaging techniques were used in this study and are summarized in this Section.

2.1 Arithmetic Averaging of Transmissivity

Heuristically, arithmetic averaging of transmissivity is obtained by linear interpolation of transmissivity between the nodes, which yields Eq. (3). Another heuristic approach ending up with arithmetic averaging is by assuming the transmissivity to be constant between the nodes (Kadioglu et al. 2008); this implies discontinuous transmissivity at each node.

One way to properly *derive* Eq. (3) is by expanding the left-hand side of Eq. (1) as the derivative of a product:

$$c\frac{d^2P}{dx^2} + \frac{dc}{dx}\frac{dP}{dx} = 0$$
(5)

The derivation is provided in the Appendix of (Lavrov 2022) and is repeated here for reader's convenience. The aim is to derive a second-order accurate approximation of Eq. (5). A second-order approximation of the first term on the left-hand side of Eq. (5) by central difference is given by

$$c\frac{d^{2}P}{dx^{2}} = \Delta x^{-2} \left(c_{i}P_{i+1} - 2c_{i}P_{i} + c_{i}P_{i-1} \right) + O\left(\Delta x^{2}\right)$$
(6)

A second-order accurate approximation of the second term on the left-hand side of Eq. (5) can be obtained by adding Taylor expansions of $c(x + \Delta x)P(x + \Delta x)$ and $c(x - \Delta x)P(x - \Delta x)$, resulting in

$$c(x + \Delta x)P(x + \Delta x) + c(x - \Delta x)P(x - \Delta x)$$

= $2c(x)P(x) + \Delta x^2 \left[c(x)\frac{d^2P(x)}{dx^2} + 2\frac{dc(x)}{dx}\frac{dP(x)}{dx} + P(x)\frac{d^2c(x)}{dx^2} \right] + O\left(\Delta x^4\right)$ (7)

Using central difference approximations for d^2P/dx^2 and d^2c/dx^2 in Eq. (7) yields the following second-order approximation for the second term on the left-hand side of Eq. (5):

$$\frac{dc}{dx}\frac{dP}{dx} = 0.5\Delta x^{-2} \left(c_{i+1}P_{i+1} + 2c_iP_i + c_{i-1}P_{i-1} - c_iP_{i+1} - c_iP_{i-1} - c_{i+1}P_i - c_{i-1}P_i \right) + O\left(\Delta x^2\right)$$
(8)

Substitution of Eqs. (6) and (8) into Eq. (5) recovers Eq. (2) with the internodal transmissivities given by Eq. (3):

$$P_{i-1}(c_{i-1}+c_i) - P_i(c_{i-1}+2c_i+c_{i+1}) + P_{i+1}(c_i+c_{i+1}) = 0$$
(9)

Deringer



Thus, arithmetic averaging of transmissivity yields second-order accuracy as previously shown by Kadioglu et al. (2008) using a different technique.

2.2 Harmonic Averaging of Transmissivity

This type of averaging is hard to justify except heuristically. The latter can be done by imagining that the real transmissivity profile is replaced with a collection of conduits. Each conduit is centered around a grid node and has a constant aperture along its length (Fig. 1).

The fracture in Fig. 1 is assumed to be composed of three segments of constant aperture. The transmissivity of each segment is equal to the transmissivity specified in the node located in the middle of the segment, i.e., c_{i-1} , c_i , c_{i+1} . Pressures are referred to the nodes as well, i.e., P_{i-1} , P_i , P_{i+1} . Consider part of the fracture located within the shadow box. It consists of two half-segments connected in series. Pressure difference across these two half-segments is given by P_{i+1} . Due to mass conservation, the flow rate through this part of the fracture is given by $\frac{2i_ic_{i+1}}{c_i+c_{i+1}} \frac{P_i-P_{i+1}}{\Delta x}$. Thus, $c_{i\pm 1/2}$ is given by Eq. (4), with the following scheme as a result:

$$P_{i-1}c_{i-1}(c_i + c_{i+1}) - P_i(c_ic_{i-1} + 2c_{i-1}c_{i+1} + c_ic_{i+1}) + P_{i+1}c_{i+1}(c_i + c_{i-1}) = 0 \quad (10)$$

By construction, this is a *heuristic* averaging method. The fact that it is based on the assumption of constant transmissivity within the cell has been recognized for at least two decades (Hyman et al. 2002). It can be shown that harmonic averaging yields second-order accuracy (Kadioglu et al. 2008), just as arithmetic averaging does. However, the magnitudes of the errors can be considerably different on some types of problems unless very fine grids are used (Kadioglu et al. 2008).

It should be noted that the assumption about piecewise-constant aperture profile implied in Eq. (4) is not realistic for fractures. Fractures have self-affine aperture maps. Thus, aperture varies at all scales, and it is only by chance that there could be a segment of constant aperture in a real fracture.

Both arithmetic and harmonic averaging procedures result in symmetric coefficient matrices. A symmetric matrix is necessary in order to ensure theoretical mass balance with the scheme (Romeu and Noetinger 1995).

2.3 Cell-Collocation-Based Averaging Techniques

When deriving the arithmetic averaging in Sect. 2.1, another, alternative path could be taken when approximating the second term on the left-hand side of Eq. (5). Namely, we could have used central finite difference for each of the derivatives, dc/dx and dP/dx, resulting in the following scheme:

$$P_{i-1}(c_{i-1} + 4c_i - c_{i+1}) - 8c_i P_i + P_{i+1}(c_{i+1} + 4c_i - c_{i-1}) = 0$$
(11)

By construction, this is a second-order scheme.

Another way to arrive at Eq. (11) is by using the cell-collocation method introduced in (Lau and Brebbia 1978). The use of cell-collocation method for deriving finite-difference approximations is detailed by Brebbia et al. (1984). Following the procedure outlined by Brebbia et al. (1984), we approximate *P* around node *i* by

$$P = P_{i-1}\varphi_{-1} + P_i\varphi_0 + P_{i+1}\varphi_1$$
(12)

where the three basis functions are given by

$$\varphi_{-1} = \frac{\hat{x}}{2\Delta x} \left(\frac{\hat{x}}{\Delta x} - 1 \right)$$

$$\varphi_0 = \left(1 - \frac{\hat{x}}{\Delta x} \right) \left(1 + \frac{\hat{x}}{\Delta x} \right)$$

$$\varphi_1 = \frac{\hat{x}}{2\Delta x} \left(\frac{\hat{x}}{\Delta x} + 1 \right)$$
(13)

with the origin of the local coordinate system, \hat{x} , being at node *i*. Since we are working with a three-node collocation cell, let us use the following interpolant for transmissivity between nodes *i*-1 and *i*+1:

$$c = c_i + \frac{c_{i+1} - c_{i-1}}{2\Delta x}\hat{x} + \frac{c_{i+1} - 2c_i + c_{i-1}}{2(\Delta x)^2}\hat{x}^2$$
(14)

The interpolant takes on the nodal values of *c* at the nodes. Substituting Eqs. (12) and (14) into Eq. (1), and using $\hat{x} = 0$ as the collocation point (i.e., the point where the residual is equal to zero) yields:

$$c_{i}\frac{P_{i+1} - 2P_{i} + P_{i-1}}{(\Delta x)^{2}} + \frac{c_{i+1} - c_{i-1}}{2\Delta x}\frac{P_{i+1} - P_{i-1}}{2\Delta x} = 0$$
(15)

i.e.,

$$P_{i-1}(c_{i-1} + 4c_i - c_{i+1}) - 8c_i P_i + P_{i+1}(c_{i+1} + 4c_i - c_{i-1}) = 0$$
(16)

Inspection of Eq. (16) reveals that it implies the following averaging for the internodal transmissivities:

$$c_{i\pm\frac{1}{2}} = \frac{1}{4} \left(\pm c_{i+1} + 4c_i \mp c_{i-1} \right)$$
(17)

We will call this *collocation-based averaging*. It is evident from Eq. (11) that the resulting coefficient matrix is not symmetric. As a consequence, this averaging technique yielded flux errors on the order of 1% in our simulations. Therefore, this averaging technique was not further considered in this study.

There are, however, at least two ways how to symmetrize Eq. (16) and thereby improve the flux error considerably. One way to symmetrize Eq. (16) is to use Eq. (17) with the lower sign for all internodal transmissivities except the one near the right boundary (see below for the boundary treatment). Essentially, this means averaging c by means of a three-node stencil where one node is located to the left of the internodal point of interest, and two nodes are located to the right of the internodal point of interest. The resulting scheme follows immediately from Eq. (16) and is given by

$$P_{i-1}(c_{i-1} + 4c_i - c_{i+1}) - (c_{i-1} + 5c_i + 3c_{i+1} - c_{i+2})P_i + P_{i+1}(c_i + 4c_{i+1} - c_{i+2}) = 0$$
(18)

(The coefficient in front of P_i ensures conservation.) We will call Eq. (18) *right-collocation-based averaging* in the remainder of this paper. The resulting coefficient matrix is symmetric.

Equation (18) has a four-node stencil (-1, 0, 1, 2) and thus, cannot be used in the rightmost node. There, Eq. (16) should be used instead. This adjustment in one node retains the symmetry of the coefficient matrix.

Another way to symmetrize Eq. (16) is by using Eq. (17) with the upper sign for all internodal transmissivities except that near the left boundary. This means averaging *c* by means of a three-node stencil where two nodes are located to the left of the internodal point of interest, and one node is located to the right of the internodal point of interest. The resulting scheme is given by

$$P_{i-1}(c_i + 4c_{i-1} - c_{i-2}) - (-c_{i-2} + 3c_{i-1} + 5c_i + c_{i+1})P_i + P_{i+1}(c_{i+1} + 4c_i - c_{i-1}) = 0$$
(19)

We will call Eq. (19) *left-collocation-based averaging* in the remainder of this paper. The resulting coefficient matrix is symmetric.

Equation (19) has a four-node stencil (-2, -1, 0, 1) and thus, cannot be used in the leftmost node. There, Eq. (16) should be used instead. This adjustment in one node retains the symmetry of the coefficient matrix.

2.4 Arithmetic Averaging of Aperture

Averaging methods considered so far in this Section were either heuristic (harmonic), or based on some formal mathematical derivation (arithmetic, collocation-based), or a combination of the two (right- and left-collocation-based). It should be remembered, however, that in the case of fracture flow the transmissivity is given by $c=w^3$ where w is the local aperture. Thus, another way to obtain the internodal transmissivities, $c_{i\pm 1/2}$, is by *reconstructing* the fracture *aperture profile* between the nodes, which would yield the estimates of the local aperture, $w_{i\pm 1/2}$.

Such a reconstruction may be performed by either local or global approximation of the aperture profile. There is no unique way to construct such an approximation. In this study, three approximations are considered: two local (arithmetic and harmonic averaging of aperture) and a global approximation by multiquadrics.

Arithmetic averaging of aperture on a uniform grid is given by

$$w_{i\pm\frac{1}{2}} = \frac{1}{2} \left(w_i + w_{i\pm 1} \right) \tag{20}$$

2.5 Harmonic Averaging of Aperture

Harmonic averaging of aperture is another way to perform local reconstruction. It is given by:

$$w_{i\pm\frac{1}{2}} = \frac{2w_i w_{i\pm1}}{w_i + w_{i\pm1}} \tag{21}$$

2.6 Global Reconstruction of Aperture Using Multiquadrics

Multiquadrics are a subclass of radial basis functions given, in general, by (Kansa 1990)

$$f(x) = \left[1 + (\hat{\varepsilon}x)^2\right]^{\beta} \tag{22}$$

where $\beta \ge -\frac{1}{2}$; $\hat{\epsilon}$ is the shape factor. In this study, multiquadrics with $\beta = \frac{1}{2}$, originally introduced by Hardy (1971), were used. By using multiquadrics as interpolants, with the grid points serving as collocation points, the approximated aperture at $i + \frac{1}{2}$ is given by (uniform grid with grid step Δx is assumed throughout this paper):

$$w_{i+\frac{1}{2}} = \sum_{k=0}^{N-1} C_k \sqrt{1 + \epsilon^2 \left(i + \frac{1}{2} - k\right)^2}$$
(23)

where N is the number of grid points (N=512 in this study); $\varepsilon = \hat{\varepsilon}\Delta x$ is the dimensionless shape factor. The condition number of the linear system used to obtain the interpolation coefficients, C_k , increases rapidly as the shape factor, ε , decreases. As ε increases, multiquadrics become more 'pointy', and the interpolation Eq. (23) approaches the arithmetic averaging of aperture, Eq. (20). Various strategies are available for choosing ε . For instance, different (e.g., random) individual shape factors can be assigned to different k (Sarra,Sturgill 2009). In this study, the simplest possible choice was made by using the same shape factor, $\varepsilon = 1.0$, for all quadrics. With this choice of ε , and N=512, the condition number is sufficiently low (2.1·10⁶).

3 Methodology

None of the averaging techniques presented in Sect. 2 appear to be preferable based on the derivation only. It should be remembered though that Eq. (1) is to be solved on self-affine linear profiles. In order to evaluate which averaging procedure suits best for solving Eq. (1) on such profiles, the following approach was employed in this study: First, a profile of 1024 nodes was constructed. Then, every second node was removed from the 1024-node profile. The hydraulic aperture of the resulting 512-node profile was computed, using the seven types of averaging introduced in Sect. 2 (arithmetic transmissivity, harmonic transmissivity, left- and right-collocation-based, reconstruction by global multiquadric approximation, arithmetic aperture, and harmonic aperture). The results of the seven computations were compared to the 'true' (reference) value which

was obtained with the internodal transmissivities corresponding to the 'true' internodal apertures, i.e., the apertures at the removed nodes of the 1024-node profile.

Profiles of 1024 nodes each were constructed as follows: Five square self-affine aperture maps of 1024×1024 nodes were generated using the spectral synthesis method (Huang et al. 1992). The Hurst exponent of these surfaces was H=0.4, 0.5, 0.6, 0.7 and 0.8. Lower Hurst exponent means that the map appears rougher. It is currently believed that most fractures have H=0.8 which is sometimes considered a "universal value" (Lenci et al. 2022b). However, lower values of H, down to 0.45, have been measured as well (Lenci et al. 2022b). The value of the Hurst exponent set in the spectral generator was verified for each surface using a recently introduced version of box-counting method (Ai et al. 2014) and was double-checked with the variogram method (Ruello et al. 2006). From each of these surfaces, 10 linear profiles of 1024 nodes were extracted in the x-direction and 10 linear profiles of 1024 nodes were extracted in the y-direction. Each of these linear profiles was offset to make its average height equal to zero. The resulting profile was rescaled so that its range be equal to a given value, Δw . ('Range' here refers to the difference between the highest peak and the deepest valley.) The following values of Δw were used: 1, 2, 3, 4 and 5 mm. In the rescaled profile, the minimum height was identified. The rescaled profile was then offset so that the minimum height be equal to a given value, w_{\min} , where w_{\min} was 1, 2, 3, 4 and 5 mm. The ratio of root mean square variation (RMS) to the mean aperture varied in the range 0.0278 to 0.3091 for x-profiles and 0.0257 to 0.3008 for y-profiles. In each of these rescaled-and-offset profiles, every other node was removed thus giving rise to a 512-node profile. The pressure diffusion equation was solved on this profile with the seven averaging techniques discussed above. The pressure applied at the left-hand boundary was equal to 1, the pressure applied at the right-hand boundary was 0. In addition, on each 512-node profile, a reference simulation was performed with the 'true' internodal transmissivities based on the apertures at the removed nodes. The total number of flow computations excluding the reference simulations was thus equal to 17,500: 5 Hurst exponents $\times 20$ linear profiles (10 in x- and 10 in y-direction) $\times 5$ range values $\times 5$ offset values $\times 7$ averaging schemes. Direct solvers were used in flow computations and in solving for multiquadric coefficients.

In each simulation, the mass balance error was evaluated as follows: The flux between the two leftmost nodes, Q_l , and the two rightmost nodes, Q_r , was evaluated. The mass balance error was the magnitude of the difference between the two divided by their average, $2|Q_l - Q_r|/(Q_l + Q_r)$. It was required for the mass balance error to be below 10^{-6} for a simulation to be accepted.

In each simulation, the error in hydraulic aperture and the error's absolute magnitude were calculated. The error is given by $\left(w_h^{(a)} - w_h^{(t)}\right) / w_h^{(t)}$ where *a* stands for averaging type (arithmetic transmissivity, harmonic transmissivity, left-collocation, right-collocation, multiquadrics, arithmetic aperture, harmonic aperture); $w_h^{(t)}$ is the 'true' value of the hydraulic aperture obtained in the reference simulation.

For each value of the Hurst exponent and each type of averaging scheme, average error in hydraulic aperture and average absolute value of the error in hydraulic aperture were calculated by summing the individual values obtained in each simulation and dividing by the number of simulations (250, i.e., 10 profiles \times 5 range values \times 5 offset values). This was done separately for profiles extracted in the *x*- and *y*-directions.

In each simulation, three norms were calculated to assess the error in the pressure distribution along the profile:

$$\left\|P^{(a)} - P^{(t)}\right\|_{1} = \sum_{k} \left|P^{(a)}_{k} - P^{(t)}_{k}\right|$$
(24)

$$\left\|P^{(a)} - P^{(t)}\right\|_{2} = \left[\sum_{k} \left(P_{k}^{(a)} - P_{k}^{(t)}\right)^{2}\right]^{1/2}$$
(25)

$$\left\|P^{(a)} - P^{(t)}\right\|_{\infty} = \max_{k} \left|P_{k}^{(a)} - P_{k}^{(t)}\right|$$
(26)

where summation was performed over internal nodes of the grid (510 in total); the meaning of superscripts is explained above.¹

For each value of the Hurst exponent and each type of averaging scheme, average norms were calculated by summing the individual values obtained in each simulation and dividing by the number of simulations (250). This was done separately for profiles extracted in the x- and y-directions.

In the simulations described so far in this Section, the minimum aperture, w_{\min} , was 1, 2, 3, 4 or 5 mm and the aperture range, Δw , was 1, 2, 3, 4 or 5 mm. In order to investigate the performance of the seven schemes with smaller apertures, an additional series of simulations was conducted, with $w_{\min}=0.1$ mm and $\Delta w=1$, 2, 3, 4 or 5 mm. Only the 1024×1024 fracture with Hurst exponent equal to 0.8 was used to construct linear profiles as this *H* value is believed to be the most common for rock fractures (Lenci et al. 2022b). As before, 10 profiles were selected in *x*-direction and 10 in *y*-direction.

4 Results

4.1 Results Obtained with $w_{\min} = 1 \text{ mm to 5 mm}$

The average errors obtained with profiles extracted in the x- and y-directions are displayed in Figs. 2 and 3, respectively. It follows from these Figures that, in the case of x-profiles, the best results are obtained with multiquadric approximation of the aperture or with harmonic averaging of the aperture. The errors obtained with these two techniques are virtually indistinguishable from each other, no matter what type of error is used as a selection criterion.

Average errors obtained with left- and right-collocation-based averaging are close, as evidenced in Figs. 2 and 3. This is hardly surprising since the schemes are similar.

Examining the errors obtained in individual simulations revealed that the magnitude of error in hydraulic aperture obtained with arithmetic averaging of transmissivity is higher than that obtained with harmonic averaging of transmissivity in the majority of cases, especially for lower H (Table 1). This is also evident from Figs. 2 and 3. Thus, if the choice is to be made between arithmetic averaging of *transmissivity* and harmonic averaging of *transmissivity*, the latter appears for self-affine linear profiles to be a better choice.

¹ The norm definitions given by Eqs. (24) and (25) are common, e.g., in functional analysis (Kreyszig 1978) but are different from those usually used in numerical analysis, where it is customary to divide these by the number of nodes (LeVeque 2002). In this study, the same number of nodes was used in all simulations, and thus, definitions (24) and (25) were used.

Fig. 2 Errors averaged over 8750 simulations performed on 512-node profiles oriented in the *x*-direction: **a** basolute value of error in hydraulic aperture; **b** error in hydraulic aperture; **c** pressure solution error using L_1 ; **d** pressure solution error using L_2 ; **e** pressure solution error using L_{∞} . Results obtained on 1024-node profiles were used as reference when computing errors in each 512-node simulation

Errors obtained with arithmetic averaging of transmissivity and displayed in Figs. 2 and 3 are relatively large. In order to understand the poor performance of arithmetic averaging, it is instructive to take a closer look at individual errors in the simulations rather than at their averaged values. In Fig. 4, errors in hydraulic aperture obtained with arithmetic averaging of transmissivity and with multiquadric reconstruction (the best performing scheme) are displayed as a function of 'true' hydraulic aperture for 250 simulations obtained on H=0.4 x-profiles. Each point represents an individual simulation. As evident from Fig. 4, errors obtained with multiquadric reconstruction are distributed symmetrically around 0, while errors obtained with arithmetic averaging of transmissivity are predominantly positive. It means that arithmetic averaging of transmissivity systematically overestimates the hydraulic aperture. A graph similar to the one in Fig. 4 is obtained also if, instead of multiquadrics reconstruction, harmonic average of transmissivity is used.

4.2 Results Obtained with $w_{\min} = 0.1 \text{ mm}$

All average errors, except signed error in the hydraulic aperture, were at their minimum when multiquadric reconstruction was used. In particular, all pressure errors (L_1, L_2, L_{∞}) were minimized this way. Only the average signed error in hydraulic aperture was minimized with an averaging technique other than multiquadrics, namely with harmonic averaging of transmissivity. For all other types of errors, the latter was the second best averaging technique. This is in contrast to the simulations discussed in Sect. 4.1 where the second-best technique was typically harmonic averaging of the aperture.

Thus, three averaging techniques stand out for flow simulations on 1D self-affine fracture profiles: (1) multiquadrics reconstruction, (2) harmonic averaging of transmissivity and (3) harmonic averaging of aperture. Multiquadrics reconstruction of the aperture profile appears, on average, to be superior, followed by the two others. The higher accuracy of the multiquadrics reconstruction is bought by the additional computational cost this technique entails, compared to all other techniques.

5 Discussion

Multiquadrics have, in the past, demonstrated surprisingly good performance in various interpolation and reconstruction jobs even though it was pointed out that the reasons for this good performance were unknown (Kansa 1990). Our work provides yet another example of an application where reconstruction with multiquadrics can be used with advantage, compared to other, traditionally used, techniques, e.g., arithmetic or harmonic averaging of transmissivity.

The choice of the shape factor for multiquadrics in this work was based solely on the requirement of sufficiently small condition number. No attempt was made to optimize further.

Harmonic averaging of aperture or transmissivity yields virtually the same accuracy as the multiquadrics reconstruction. Computational overhead of the latter is, however, higher.







The fact that the second best option is provided by harmonic averaging schemes is consistent with the earlier results by Romeu and Noetinger (1995) who found harmonic averaging of transmissivity to be a superior scheme for linear flow in porous media. They did not consider multiquadrics reconstruction though (the multiquadrics-based method makes sense for *fracture* flow where it has a clear geometric meaning).

The focus in this paper has been on the flow along random profiles. It is instructive to compare the performance of the seven averaging techniques on a smooth profile for which an analytical solution is available. As an example, consider transmissivity given by

$$c(x) = 1 + \cos^2(2\pi x), 0 \le x \le 1$$
(27)

Equation (27) represents a smooth periodic transmissivity. The analytical solution to Eq. (1) with $P_0 = 1$, L = 1 and c(x) given by Eq. (27) is

$$P(x) = \begin{cases} 1 - \frac{1}{2\pi} \arctan \frac{\tan (2\pi x)}{\sqrt{2}}, x \in (0, 0.25] \\ 0.5 - \frac{1}{2\pi} \arctan \frac{\tan (2\pi x)}{\sqrt{2}}, x \in [0.25, 0.75) \\ -\frac{1}{2\pi} \arctan \frac{\tan (2\pi x)}{\sqrt{2}}, x \in [0.75, 1.0) \end{cases}$$
(28)

The analytical solution is displayed in Fig. 5 along with the numerical solution obtained on a 512-node grid with multiquadric reconstruction of $c^{1/3}$. Errors in hydraulic aperture obtained with the seven averaging schemes and N=512 nodes are summarized in Table 2.

From Table 2, multiquadrics approximation yields, with a good margin, the best accuracy on a smooth periodic profile. This is contrary to the results obtained with self-affine profiles where multiquadrics reconstruction and harmonic averaging of the aperture produced very similar errors. This is yet another illustration that there is apparently no universal choice of the averaging technique and the choice should be determined by the nature of the problem at hand, a conclusion made by Kadioglu et al. (2008).

Excellent performance of multiquadrics on a profile given by Eq. (27) should perhaps come as no surprise since the profile is smooth and the global reconstruction technique does a superior job here (at the expense of some computational overhead).

It should be noted that we called multiquadrics approximation, arithmetic averaging of aperture and harmonic averaging of aperture "reconstruction techniques" in this text. However, all seven averaging techniques examined in this study could be considered as reconstruction techniques since in all of them we attempt to assign transmissivity to a location between the nodes.

Uniform grids were used in this study. Uniform grids are typically used in fracture flow simulations. Averaging on nonuniform grids should be investigated in the future.

Only linear (1D) fracture profiles were considered in this study. In practice, fracture flow simulations are performed on two-dimensional aperture landscapes. Dimensionality has a crucial effect on conductivity and flow structure, as demonstrated, e.g., by Colecchio et al.(2021). The results of our study might therefore not be directly applicable to 2D grids. In particular, in 1D, even a single node with zero aperture results in a trivial solution (zero flow rate), and a simulation does not need to be performed. Treatment of zero apertures is therefore not relevant for 1D schemes. In 2D, zero apertures may still result in a conductive fracture, and a simulation *can* be performed. In this case, in order to avoid a singular coefficient matrix, zero apertures are usually replaced with very small values. Comparison of different averaging schemes on 2D grids, especially with clusters of nodes having very small apertures, should be the subject of a future work. It should be noted, however, that it might be difficult to match the number of 1D simulations used in this study (a total of 17,500) when doing 2D simulations.

The work presented in this paper focused on the effects that averaging has on the *flow*. It might be worthwhile, in the future, to examine also how these seven reconstruction techniques affect *geometric* properties of the reconstructed fractures, e.g., their correlation length, RMS, Hurst exponent, etc.

All development in this paper has been for Newtonian fluids. The situation would become more complicated with non-Newtonian fluids. This can be illustrated using the **Fig.3** Errors averaged over 8750 simulations performed on 512-node profiles oriented in the y-direction: **a** basolute value of error in hydraulic aperture; **b** error in hydraulic aperture; **c** pressure solution error using L_1 ; **d** pressure solution error using L_2 ; **e** pressure solution error using L_{∞} . Results obtained on 1024-node profiles were used as reference when computing errors in each 512-node simulation

simplest possible non-Newtonian rheology, i.e., a power-law fluid. For a linear fracture flow under the lubrication theory approximation, the fluid velocity averaged across the fracture aperture is given for such a fluid by equation (Lavrov 2023)

$$\mathbf{v} = -\frac{n}{2n+1} \frac{1}{C^{1/n}} \left(\frac{w}{2}\right)^{(n+1)/n} \left|\frac{dP}{dx}\right|^{(1-n)/n} \frac{dP}{dx}$$
(29)

where *n* is the flow index; *C* is the consistency index. Mass conservation for one-dimensional steady incompressible flow, i.e., $\nabla(w\mathbf{v}) = 0$, results in a *p*-Laplace equation:

$$\frac{d}{dx}\left[w^{(2n+1)/n}\left|\frac{dP}{dx}\right|^{(1-n)/n}\frac{dP}{dx}\right] = 0$$
(30)

Inspection of Eqs. (29-30) suggests that the very concept of local fracture conductivity as a proportionality factor between velocity and pressure gradient acquires an additional dimension here: it becomes dependent on the pressure gradient. One promising approach here could be using reconstruction of aperture instead of averaging of transmissivity. This approach was pursued, e.g., in the recent work by Lenci et al. where arithmetic averaging of *w* between nodes was used to approximate the internodal aperture (Lenci et al. 2022b). As another alternative, the method of aperture reconstruction with multiquadrics advocated in the present paper could be tried for Eq. (30). Whether this method has advantages as compared to other averaging techniques also in the case of non-Newtonian fluids should be investigated in a dedicated study.

It should be stressed that conclusions drawn in our study pertain to *self-affine* linear profiles. It is straightforward to construct an example of profile where local averaging techniques, e.g., arithmetic or harmonic averaging of aperture or transmissivity, should perform *at least* as well as multiquadrics. An example I will use here is inspired by the recent work on 'binary media' (Colecchio et al. 2021) where fractures were constructed such that local aperture could take only two values,: w_l or w_u . Let us construct a special version of such 'binary media', namely a linear profile of 100 nodes where all nodes with index i < 50 have $w = w_l$, while all nodes with i > 49 have $w = w_u$. On such a profile, any local averaging method, e.g., arithmetic or harmonic averaging of aperture or transmissivity, should perform *at least* as well as any global reconstruction technique such as multiquadrics reconstruction.

Moreover, the very *methodology* employed in our study, with removing every other node and using the original profile for a reference computation, was designed for selfaffine profiles and may become faulty in other cases, e.g., when the aperture profile has a distinctive regular structure. As an illustration, let us again construct a 'binary' profile of a 100 nodes where every node with an even index has $w = w_l$, while every node with an odd index has $w = w_u$. This resembles the profile studied, e.g., by Haugerud et al. (2022). This is the profile to be used for the reference simulation. Now, if every other node is removed from this profile and a flow simulation is performed, all seven averaging methods considered in this study will perform poorly. From the averaging point of view, there is no way to distinguish between two possible original profiles once



🖄 Springer



Fig. 3 (continued)

every other node has been removed: the one where all apertures were the same and the one where even nodes had $w = w_l$ and odd nodes had $w = w_u$. Once the nodes have been removed, the information is lost (cf. Nyquist theorem). This brings us to the next issue, that of correlation length in the original profile. In practice, various degree of correlation may exist in an aperture profile, with the correlation length being determined by various factors, e.g., by the magnitude of the shear displacement (Brown et al. 1986; Wang et al. 1988). For instance, if two identical (matching) self-affine fracture surfaces are put together with zero offset, the resulting aperture profile will be w = const (i.e., infinite correlation length), and any local averaging technique in a numerical flow simulation will perform perfectly [cf. also the results of Nicholl et al. (1999) where several local averaging techniques were tested and their results converged at high mesh resolution, namely when the grid spacing became smaller than 1/5 of the correlation length]. Global reconstruction techniques, e.g., based on multiquadrics, would have no

Table 1 Fraction of simulations in which the magnitude of error in hydraulic aperture obtained with arithmetic averaging of transmissivity was greater than with harmonic averaging of transmissivity x.	Type of profile	Hurst exponent, H	Fraction of simulations	
	x-profiles	0.4	0.536	
		0.5	0.516	
		0.6	0.508	
		0.7	0.500	
		0.8	0.500	
	y-profiles	0.4	0.568	
		0.5	0.548	
		0.6	0.536	
		0.7	0.528	

0.02

Fig. 4 Signed error in hydraulic aperture obtained in individual simulations with arithmetic averaging of transmissivity and with multiquadric reconstruction vs. 'true' hydraulic aperture. Data for simulations with H = 0.4 are displayed. Each point represents an individual simulation



Error in hydraulic aperture 0.015 0.01 0.005 0 0.005 O -0.01 'True' hydraulic aperture, mm arithmetic avreaging of transmissivity reconstruction with multiquadrics 1 --- Analytical solution 0.8 Multiquadrics reconstruction ressure 0.6 0.4 0.2 0 0 0.2 0.4 0.6 0.8 1 х

0.8

0.544

advantage in this case. Thus, there might be an extra angle to the problem of choosing between different averaging techniques in fracture flow modeling. The role of correlation length and profile structure in this choice should be examined in the future.

Finally, it should be emphasized that this study makes use of the lubrication theory approximation when formulating the flow problem. Our findings suggest that multiquadrics

Averaging method	<i>c</i> ,	- 1 - 73		
	Pressure error L_1	Pressure error L_2	Pressure error L_{∞}	
Arithmetic transmissivity	4.887.10-4	2.410.10-5	1.524.10-6	
Harmonic transmissivity	$4.612 \cdot 10^{-4}$	$2.334 \cdot 10^{-5}$	$1.579 \cdot 10^{-6}$	
Right-collocation	$9.500 \cdot 10^{-4}$	$4.708 \cdot 10^{-5}$	$3.075 \cdot 10^{-6}$	
Left-collocation	9.500.10-4	$4.708 \cdot 10^{-5}$	$3.075 \cdot 10^{-6}$	
Multiquadrics reconstruction	$2.041 \cdot 10^{-4}$	$1.086 \cdot 10^{-5}$	9.186·10 ⁻⁷	
Arithmetic aperture	$4.796 \cdot 10^{-4}$	$2.368 \cdot 10^{-5}$	$1.504 \cdot 10^{-6}$	
Harmonic aperture	4.700.10-4	2.343.10-5	1.524.10-6	

Table 2 Errors in the pressure distribution obtained with six averaging techniques on the benchmark problem [Eqs. (1), (27)]. Errors are with regard to the analytical solution [Eq. (28)]

reconstruction provides superior results *when the flow is modeled using the lubrication approximation*. This does *not* suggest that multiquadrics reconstruction provides the most accurate flow solution in general. The effect of out-of-plane flow on the accuracy of 1D flow simulations has been recognized in the research community for at least 25 years by now (Oron and Berkowitz 1998; Nicholl and Detwiler 2001; Brush and Thomson 2003; Wang et al. 2015, 2018, 2020; He et al. 2021). An earlier attempt to reconcile the use of Eq. (1) with experimental results obtained using Hele-Shaw cells with artificial rough surfaces (not self-affine), while solving the flow with different averaging techniques, can be found in the work by Nicholl et al. (1999) There, corrections were introduced in order to account for out-of-plane flow dynamics, all with limited success.

6 Conclusions

The following conclusions are drawn:

- In most cases studied here, global reconstruction of the aperture profile using multiquadrics provided superior accuracy in terms of both the modeled hydraulic aperture and the pressure error for fluid flow on one-dimensional self-affine aperture profiles.
- The second best averaging technique was either harmonic averaging of the aperture or harmonic averaging of transmissivity. The former can outperform multiquadrics with regard to certain types of errors when the ratio of the minimum aperture to the aperture range is sufficiently large. The latter can outperform multiquadrics reconstruction with regard to certain types of errors when the ratio of the minimum aperture to the aperture range is sufficiently soft errors when the ratio of the minimum aperture to the aperture range is sufficiently soft errors when the ratio of the minimum aperture to the aperture range is sufficiently small.
- On a smooth periodic profile, multiquadrics reconstruction outperformed all other techniques examined here, with a good margin.

Acknowledgements The author is grateful to the anonymous reviewers for useful comments and suggestions.

Authors' Contributions Not applicable.

Funding Open access funding provided by NTNU Norwegian University of Science and Technology (incl St. Olavs Hospital - Trondheim University Hospital). No funds, grants, or other support was received.

Data Availability Material All the necessary data are provided in the text of the article.

Code Availability Not applicable.

Declarations

Conflict of interest The authors have no conflicts of interest to declare that are relevant to the content of this article.

Open Access This article is licensed under a Creative Commons Attribution 4.0 International License, which permits use, sharing, adaptation, distribution and reproduction in any medium or format, as long as you give appropriate credit to the original author(s) and the source, provide a link to the Creative Commons licence, and indicate if changes were made. The images or other third party material in this article are included in the article's Creative Commons licence, unless indicated otherwise in a credit line to the material. If material is not included in the article's Creative Commons licence and your intended use is not permitted by statutory regulation or exceeds the permitted use, you will need to obtain permission directly from the copyright holder. To view a copy of this licence, visit http://creativecommons.org/licenses/by/4.0/.

References

- Aavatsmark, I.: An introduction to multipoint flux approximations for quadrilateral grids. Comput. Geosci. 6(3), 405–432 (2002). https://doi.org/10.1023/A:1021291114475
- Ai, T., Zhang, R., Zhou, H.W., Pei, J.L.: Box-counting methods to directly estimate the fractal dimension of a rock surface. Appl. Surf. Sci. 314, 610–621 (2014)
- Bessone, L., Gamazo, P., Dentz, M., Storti, M., Ramos, J.: GPU implementation of explicit and implicit Eulerian methods with TVD schemes for solving 2D solute transport in heterogeneous flows. Comput. Geosci. 26(3), 517–543 (2022). https://doi.org/10.1007/s10596-022-10136-8
- Bouchaud, E., Lapasset, G., Planès, J.: Fractal dimension of fractured surfaces: a universal value? Europhys. Lett. 13(1), 73–79 (1990). https://doi.org/10.1209/0295-5075/13/1/013
- Brebbia, C.A., Telles, J.C.F., Wrobel, L.C.: Boundary Element Techniques: Theory and Applications in Engineering. Springer, Berlin (1984)
- Brown, S.R., Kranz, R.L., Bonner, B.P.: Correlation between the surfaces of natural rock joints. Geophys. Res. Lett. 13(13), 1430–1433 (1986). https://doi.org/10.1029/GL013i013p01430
- Brush, D.J., Thomson, N.R.: Fluid flow in synthetic rough-walled fractures: Navier-Stokes, Stokes, and local cubic law simulations. Water Resour. Res. (2003). https://doi.org/10.1029/2002WR001346
- Colecchio, I., Boschan, A., Otero, A.D., Noetinger, B.: On the multiscale characterization of effective hydraulic conductivity in random heterogeneous media: a historical survey and some new perspectives. Adv. Water Resour. 140, 103594 (2020). https://doi.org/10.1016/j.advwatres.2020.103594
- Colecchio, I., Otero, A.D., Noetinger, B., Boschan, A.: Equivalent hydraulic conductivity, connectivity and percolation in 2D and 3D random binary media. Adv. Water Resour 158, 104040 (2021). https://doi. org/10.1016/j.advwatres.2021.104040
- Hardy, R.L.: Multiquadric equations of topography and other irregular surfaces. J. Geophys. Res. 76(8), 1905–1915 (1971). https://doi.org/10.1029/JB076i008p01905
- Haugerud, I., Linga, G., Flekkøy, E.G.: Solute dispersion in channels with periodic square boundary roughness. J. Fluid Mech. 944, A53 (2022). https://doi.org/10.1017/jfm.2022.522
- He, X., Sinan, M., Kwak, H., Hoteit, H.: A corrected cubic law for single-phase laminar flow through roughwalled fractures. Adv. Water Resour. 154, 103984 (2021). https://doi.org/10.1016/j.advwatres.2021. 103984
- Huang, S.L., Oelfke, S.M., Speck, R.C.: Applicability of fractal characterization and modelling to rock joint profiles. Int. J. Rock Mech. Min. Sci. Geomech. Abstr. 29(2), 89–98 (1992). https://doi.org/10.1016/ 0148-9062(92)92120-2
- Hyman, J., Morel, J., Shashkov, M., Steinberg, S.: Mimetic finite difference methods for diffusion equations. Comput. Geosci. 6(3), 333–352 (2002). https://doi.org/10.1023/A:1021282912658
- Kadioglu, S.Y., Nourgaliev, R.R., Mousseau, V.A.: A comparative study of the harmonic and arithmetic averaging of diffusion coefficients for non-linear heat conduction problems. Report INL/EXT-08– 13999. In. Idaho National Laboratory, Idaho Falls, Idaho (2008)

- Kansa, E.J.: Multiquadrics—a scattered data approximation scheme with applications to computational fluid-dynamics—I surface approximations and partial derivative estimates. Comput. Math. Appl. 19(8), 127–145 (1990). https://doi.org/10.1016/0898-1221(90)90270-T
- Klepikova, M., Méheust, Y., Roques, C., Linde, N.: Heat transport by flow through rough rock fractures: a numerical investigation. Adv. Water Resour. 156, 104042 (2021). https://doi.org/10.1016/j.advwatres. 2021.104042
- Kreyszig, E.: Introductory Functional Analysis with Applications. Wiley, Hoboken (1978)
- Langevin, C.D., Hughes, J.D., Banta, E.R., Niswonger, R.G., Panday, S., Provost, A.M.: Documentation for the MODFLOW 6 Groundwater Flow Model. U.S. Geological Survey, Reston, Virginia (2017)
- Lau, P.C.M., Brebbia, C.A.: The cell collocation method in continuum mechanics. Int. J. Mech. Sci. 20(2), 83–95 (1978). https://doi.org/10.1016/0020-7403(78)90070-X
- Lavrov, A.: Comparison of symmetric and asymmetric schemes with arithmetic and harmonic averaging for fracture flow on Cartesian grids. Transp. Porous Media 142(3), 585–597 (2022). https://doi.org/ 10.1007/s11242-022-01760-0
- Lavrov, A.: Flow of non-Newtonian fluids in single fractures and fracture networks: Current status, challenges, and knowledge gaps. Eng. Geol. 321, 107166 (2023). https://doi.org/10.1016/j.enggeo. 2023.107166
- Lenci, A., Méheust, Y., Putti, M., Di Federico, V.: Monte Carlo simulations of shear-thinning flow in geological fractures. Water Resour. Res. 58(9), e2022WR032024 (2022a). https://doi.org/10.1029/ 2022WR032024
- Lenci, A., Putti, M., Di Federico, V., Méheust, Y.: A lubrication-based solver for shear-thinning flow in rough fractures. Water Resour. Res. 58(8), e2021WR031760 (2022b). https://doi.org/10.1029/ 2021WR031760
- LeVeque, R.J.: Finite Volume Methods for Hyperbolic Problems. Cambridge University Press, Cambridge (2002)
- Lin, C., Taleghani, A.D., Kang, Y., Xu, C.: A coupled CFD-DEM simulation of fracture sealing: effect of lost circulation material, drilling fluid and fracture conditions. Fuel **322**, 124212 (2022). https:// doi.org/10.1016/j.fuel.2022.124212
- Måløy, K.J., Hansen, A., Hinrichsen, E.L., Roux, S.: Experimental measurements of the roughness of brittle cracks. Phys. Rev. Lett. 68(2), 213–215 (1992). https://doi.org/10.1103/PhysRevLett.68.213
- Mandelbrot, B.B., Passoja, D.E., Paullay, A.J.: Fractal character of fracture surfaces of metals. Nature 308(5961), 721–722 (1984). https://doi.org/10.1038/308721a0
- Mao, S., Wu, K., Moridis, G.: Integrated simulation of three-dimensional hydraulic fracture propagation and Lagrangian proppant transport in multilayered reservoirs. Comput. Methods Appl. Mech. Eng. 410, 116037 (2023). https://doi.org/10.1016/j.cma.2023.116037
- Moreno, L., Tsang, Y.W., Tsang, C.F., Hale, F.V., Neretnieks, I.: Flow and tracer transport in a single fracture: a stochastic model and its relation to some field observations. Water Resour. Res. 24(12), 2033–2048 (1988). https://doi.org/10.1029/WR024i012p02033
- Nicholl, M.J., Detwiler, R.L.: Simulation of flow and transport in a single fracture: macroscopic effects of underestimating local head loss. Geophys. Res. Lett. 28(23), 4355–4358 (2001). https://doi.org/ 10.1029/2001GL013647
- Nicholl, M.J., Rajaram, H., Glass, R.J., Detwiler, R.: Saturated flow in a single fracture: evaluation of the Reynolds equation in measured aperture fields. Water Resour. Res. 35(11), 3361–3373 (1999). https://doi.org/10.1029/1999WR900241
- Oron, A.P., Berkowitz, B.: Flow in rock fractures: The local cubic law assumption reexamined. Water Resour. Res. 34(11), 2811–2825 (1998). https://doi.org/10.1029/98WR02285
- Plouraboué, F., Kurowski, P., Hulin, J.-P., Roux, S., Schmittbuhl, J.: Aperture of rough cracks. Phys. Rev. E 51(3), 1675–1685 (1995). https://doi.org/10.1103/PhysRevE.51.1675
- Romeu, R.K., Noetinger, B.: Calculation of internodal transmissivities in finite difference models of flow in heterogeneous porous media. Water Resour. Res. 31(4), 943–959 (1995). https://doi.org/10.1029/ 94WR02422
- Roth, C., Chilès, J.-P., de Fouquet, C.: Adapting geostatistical transmissivity simulations to finite difference flow simulators. Water Resour. Res. 32(10), 3237–3242 (1996). https://doi.org/10.1029/ 96WR01828
- Ruello, G., Blanco-Sanchez, P., Iodice, A., Mallorqui, J.J., Riccio, D., Broquetas, A., Franceschetti, G.: Synthesis, construction, and validation of a fractal surface. IEEE Trans. Geosci. Remote Sens. 44(6), 1403–1412 (2006). https://doi.org/10.1109/TGRS.2006.870433
- Sarra, S.A., Sturgill, D.: A random variable shape parameter strategy for radial basis function approximation methods. Eng. Anal. Boundary Elem. 33(11), 1239–1245 (2009). https://doi.org/10.1016/j. enganabound.2009.07.003

- Schmittbuhl, J., Måløy, K.J.: Direct observation of a self-affine crack propagation. Phys. Rev. Lett. 78(20), 3888–3891 (1997). https://doi.org/10.1103/PhysRevLett.78.3888
- Tsang, C.-F., Neretnieks, I.: Flow channeling in heterogeneous fractured rocks. Rev. Geophys. 36(2), 275–298 (1998)
- van Es, B., Koren, B., de Blank, H.J.: Finite-difference schemes for anisotropic diffusion. J. Comput. Phys. 272, 526–549 (2014)
- Wang, J.S.Y., Narasimhan, T.N., Scholz, C.H.: Aperture correlation of a fractal fracture. J. Geophys. Res. Solid Earth 93(B3), 2216–2224 (1988). https://doi.org/10.1029/JB093iB03p02216
- Wang, L., Cardenas, M.B., Slottke, D.T., Ketcham, R.A., Sharp, J.M., Jr.: Modification of the local cubic law of fracture flow for weak inertia, tortuosity, and roughness. Water Resour. Res. 51(4), 2064– 2080 (2015). https://doi.org/10.1002/2014WR015815
- Wang, Z., Xu, C., Dowd, P.: A modified cubic law for single-phase saturated laminar flow in rough rock fractures. Int. J. Rock Mech. Min. Sci. 103, 107–115 (2018). https://doi.org/10.1016/j.ijrmms.2017.12. 002
- Wang, Z., Xu, C., Dowd, P., Xiong, F., Wang, H.D.: A nonlinear version of the Reynolds equation for flow in rock fractures with complex void geometries. Water Resour. Res. 56(2), 149 (2020). https://doi.org/ 10.1029/2019WR026149
- Zimmerman, R.W., Kumar, S., Bodvarsson, G.S.: Lubrication theory analysis of the permeability of roughwalled fractures. Int. J. Rock Mech. Min. Sci. Geomech. Abstr. 28(4), 325–331 (1991). https://doi.org/ 10.1016/0148-9062(91)90597-F

Publisher's Note Springer Nature remains neutral with regard to jurisdictional claims in published maps and institutional affiliations.