Adaptive Sampling of Surface Fronts in the Arctic using an Autonomous Underwater Vehicle

4	Trygve Olav Fossum ^{<i>a,b</i>} , Petter Norgren ^{<i>a,b</i>} , Ilker Fer ^{<i>c,d</i>} , Frank Nilsen ^{<i>d,c</i>} , Zoe		
5	Koenig ^{c,e} and Martin Ludvigsen ^{a,b,d} ^a Department of Marine Technology,		
6	Norwegian University of Science and Technology (NTNU).		
7	^b Centre for Autonomous Marine Operations and Systems (AMOS), Trondheim,		
8	Norway.		
9	^c Geophysical Institute, University of Bergen and Bjerknes Centre for Climate		
10	Research, Bergen, Norway.		
11	^d The University Centre in Svalbard (UNIS), Longyearbyen, Norway.		
12	^e Norwegian Polar Institute, Tromsø, Norway.		

13

Abstract

Fronts between Arctic- and Atlantic-origin waters are characterized by strong lateral gradients in 14 temperature and salinity. Ocean processes associated with fronts are complex with considerable space 15 and time variability. Resolving the processes in frontal zones by observation is therefore challenging 16 but important for understanding the associated physical-biological interactions and their impact on the 17 marine ecosystem. The use of autonomous robotic vehicles and in situ data-driven sampling can help 18 improve and augment the traditional sampling practices such as ships and profiling instruments. Here 19 we present the development and results of using an autonomous agent for detection and sampling 20 of an Arctic front, integrated on board an autonomous underwater vehicle. The agent is based on a 21 subsumption architecture implemented as behaviors in a finite state machine. Once a front is detected, 22 the front tracking behavior uses observations to continuously adapt the path of the vehicle to perform 23 transects across the front interface. Following successful sea trials in the Trondheimsfjord, the front 24 tracking agent was deployed to perform a full-scale mission near 82°N north of Svalbard, close to the 25

sea ice edge. The agent was able to detect and track an Arctic frontal feature, performing a total of
 six crossings while collecting vertical profiles in the upper 90 m of the water column. Measurements
 yield a detailed volumetric description of the frontal feature with high resolution along the frontal zone,
 augmenting ship-based sampling that was run in parallel.

Index terms—Autonomous underwater vehicle (AUV), fronts, Arctic, data-driven sampling, adaptive sampling.

33

30

I. INTRODUCTION

Oceanic fronts are dynamic boundaries where different water masses meet and interact. Large changes in water properties (such as temperature, salinity, or dissolved oxygen concentration) result in elevated horizontal gradients that can be used to detect their presence. Ocean ecosystems thrive in fronts [1], which are typically associated with enhanced primary production [2], [3], high concentrations of zooplankton and fish larvae [4], and biogeochemical cycling [3]. Flow convergence along fronts also induces patchiness, vertical mixing, and increased nutrient supply [5].

The observational practices of frontal processes are usually scale dependent, as different frontal 41 processes occur across a variety of spatial scales (from 1-10 km [sub-mesoscale] up to 50 km 42 [mesoscale]) and directions (along-front and cross-frontal process scales) [6]; propagating vor-43 tices, frontal meandering, and ocean turbulent patchiness further introduce temporal variability. 44 A lack of resolution in time or space can therefore fail to capture important dynamics and 45 their variability. Traditional ocean sensing practices, such as profiling at stations occupied by 46 a ship, or transects using sensors towed behind a vessel, impose both economic and logistical 47 limitations that compromise the spatial and temporal coverage of the observations. Satellite 48 imagery from polar orbiting satellites have been extensively used to observe ocean surface fronts, 49 see e.g. [7]. As there is an inherit latency between capture and availability, surface coverage 50 only, cloud dependence, combined with the tidal movements and ocean dynamics, the use of 51 satellite-based remote sensing disallow direct use in resolving subsurface details and tracking 52 the front. Thus, a sustained focus on effective use and consolidation of sampling resources and 53 acquisition strategies has consequently emerged [8]–[11]. As a result of these developments, the 54

⁵⁵ use of robotic assets has increased, providing mobility and adaptive sampling capabilities that
 ⁵⁶ can substantially augment current ocean observation practices.

Autonomous underwater vehicles (AUVs) are effective for observing sub-mesoscale features 57 such as fronts, and have been used in numerous water-column sampling applications [8], [11]– 58 [13] and field programs. For instance, the Autonomous Ocean Sampling Network (AOSN-I/-II) 59 [14], [15] conducted a study to understand how ocean variability could be observed using a 60 coordinated network of assets. Central to the success of these efforts were decentralized adaptive 61 sampling strategies running locally on board the sensing platforms. Such sampling is capable 62 of adjusting the mission execution according to the evolving dynamics of a feature-such as a 63 front or a coherent vortex-without human intervention. 64

Polar regions, particularly in the marginal ice zone, are characterized by strong lateral gradients 65 in water mass properties in the upper water column. In the region north of Svalbard, the warm 66 and saline Atlantic Water (AW) flows into the Arctic Ocean as a boundary current leading 67 to open (sometimes fragmented), ice-free waters in the region. This has consequences for the 68 marine ecosystem, regional air-ice-sea interaction processes, and thus, vertical mixing and ocean 69 circulation [16]. The high-latitude regions are also challenging environments for AUV operations 70 due the presence of drifting sea ice and a harsh climate (low temperatures and strong winds) 71 impacting deployment and recovery operations. The lack of over-the-horizon communication 72 infrastructure (often limited to satellite only) also increases complexity and requirements for these 73 systems to be self reliant. The remote and isolated locations imply that the time available for 74 sampling is highly valuable and must be used efficiently. Autonomous assets have the potential to 75 contribute to data collection and knowledge generation by delivering target-specific and intelligent 76 sampling schemes which aim to resolve processes and scales of interest. 77

⁷⁸ Using an AUV equipped with a CTD (conductivity, temperature, depth) sensor and water qual-⁷⁹ ity sensors (more details in Section IV-A), we have developed and tested a sampling algorithm ⁸⁰ for detection, tracking, and mapping of a polar front. The aim of the algorithm is two-fold: i) ⁸¹ automatically detect and track the front based on the horizontal gradient of temperature, and ⁸² ii) refine and increase the sampling resolution both along the front and sufficiently into the ⁸³ water masses on either side of the front. The AUV was deployed and recovered during the ⁸⁴ Nansen Legacy cruise on R/V Kronprins Haakon in September 2018, from which CTD profiles were collected in parallel during the AUV operation. An overview of the experiment setup and involved systems are shown in Fig. 1(a), together with the deployment location relative to Svalbard (Fig. 1(b)).

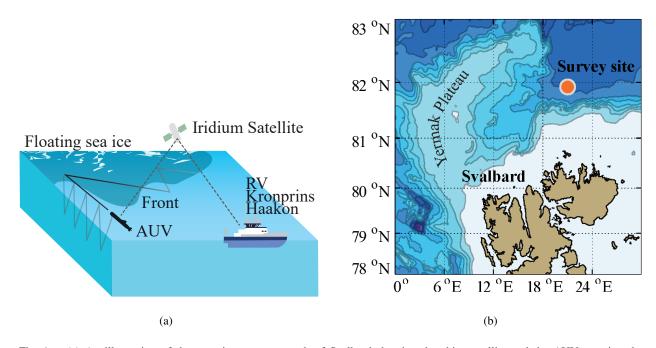


Fig. 1. (a) An illustration of the experiment setup north of Svalbard showing the ship, satellite and the AUV crossing the front between warm (light blue) and cold (dark blue) waters close to the ice edge (white). (b) Map showing the location of the experiment relative to Svalbard. Isobaths are at 500 m intervals.

II. RELATED WORK

88

Autonomous and adaptive feature tracking of frontal features was explored for a horizontal (i.e. 89 two-dimensional [2D]) single-location (across-front sampling) upwelling system in [12] using 90 the horizontal temperature gradient in a stratified water column, where the vertical temperature 91 difference is large between stratified layers compared to upwelling (mixed) water. The method 92 was successfully used to map an upwelling frontal system in Monterey Bay, California, April 93 2011, completing 14 transects across the front. The method was later extended to accommodate 94 along-front tracking in [17] for deployment on board an unmanned surface vehicle. Building 95 on this, as well as related experiments by [18], [19] developed a method for both across- and 96 along-front tracking capable of both 2D and three-dimensional (3D) tracking. The method was 97 tested in a virtual environment using ocean model data. An adaptive path planning algorithm 98

for 2D tracking of ocean fronts was also presented in [13], based on adjusting a cubic spline 99 from a set of pre-defined assumptions about the properties of a front. A zig-zag pattern was 100 then derived based on the adjusted curve to cover the across-front variation. The method was 101 restricted to tracking only, but could not find or re-locate the front. In [20] an adaptive sampling 102 and tracking algorithm of a near-shore frontal feature (a river plume) was explored outside Porto, 103 Portugal using a state-based autonomous agent on board an AUV that performs tracking on a 104 threshold detection of salinity (isopycnal) and state switching. As the method was developed 105 for a disc-shaped river plume, a constant angle increment was used to traverse the disk feature. 106 Other related work on reactive tracking of thermocline features using AUVs includes [21]–[23]; 107 each describe the procedure and results of tracking the vertical temperature gradient in the 108 water column using a state machine with different behaviors that issues depth specific set-points 109 (in each case, adaptation only occurs in the z-direction, one-dimensional [1D] tracking). An 110 important aspect for all front tracking to keep contact with the boundary feature. In [24] this 111 is addressed by adapting the orientation of the crossings to the local curvature of the boundary. 112 When available, multi vehicle approaches are able to resolve spatial features more effectively 113 and in turn provide an enhanced synopticity of the frontal boundary, as demonstrated in e.g. 114 [25], where several platforms were coordinated to sample an upwelling front in Monterey Bay. 115

The method of front tracking presented in this paper is intended for a single vehicle operating 116 in 2D using an isotherm as a frontal indicator. However, to increase robustness and adapt these 117 methods to Arctic conditions, the method introduces two distinctive elements: i) a "Regain 118 Maneuver" captures the front if it is lost by the tracking algorithm; ii) Hysteresis and double 119 detection verification elements preventing false or spurious detection of the front, which may 120 arise when passing through filaments generated from advection and turbulence. The method also 121 features a rigorous satellite-based reporting scheme informing the operator on the progress and 122 status of the mission. The latter reduces risk when operating in harsh and rapidly changing sea 123 conditions close to the ice edge and far from supporting infrastructure. 124

126 A. Front Detection and Tracking Algorithm

The method is based on a finite state machine (FSM) following classical reactive subsumptionbased approaches [26]. The main idea is to generate a zigzag maneuver triggered by a switch between two main descriptive states dependent on the AUV location relative to the front (inside/outside). As the front is partially due to melt water (with low density), the strongest frontal signature is near the surface, hence only measurements from a pre-defined depth interval (0.5-8 m) are used for detection. For the surface fronts discussed here, this depth interval can be a fixed constant or be adapted based on experience. We define the average temperature from the depth interval as

$$\mu_d = \frac{1}{n_d} \sum_{d=1}^{n_d} t_d,$$
(1)

where n_d is the number of measurements inside the depth interval, and t_d the associated temperatures. To determine the position to the AUV relative to the front we define a variable s_{front} that can take on the values $s_{front} = inside$ or $s_{front} = outside$. We also define a variable \tilde{s}_{front} as follows

$$\tilde{s}_{front} = \begin{cases} inside & \text{if } T_{isotherm} - T_{hysteresis} > \mu_d, \\ outside & \text{if } T_{isotherm} + T_{hysteresis} < \mu_d, \end{cases}$$
(2)

where $T_{isotherm}$ is the constant defining the frontal isotherm and $T_{hysteresis}$ is the constant defining the hysteresis sensitivity. A frontal crossing is confirmed if $s_{front} \neq \tilde{s}_{front}$ for two consecutive yoyo envelopes, in which case s_{front} changes from *inside* \rightarrow *outside* or from *outside* \rightarrow *inside*.

The FSM, its switching logic, and flow diagram are shown in Fig. 2. To accompany the diagram, a step-by-step description is given to further explain the adaptive behavior.

- ¹³² Step-by-step description of the FSM:
- 133 1) *Search mode* Run along a pre-defined transect (chosen by the operator) towards the front.
- a) Each minute, check if there are measurements available from the predefined depth interval. If observations are available, calculate the mean temperature μ_d .
- b) Check if the observed mean temperature μ_d is outside the hysteresis threshold according to Eq. (2), and determine the current \tilde{s}_{front} state.

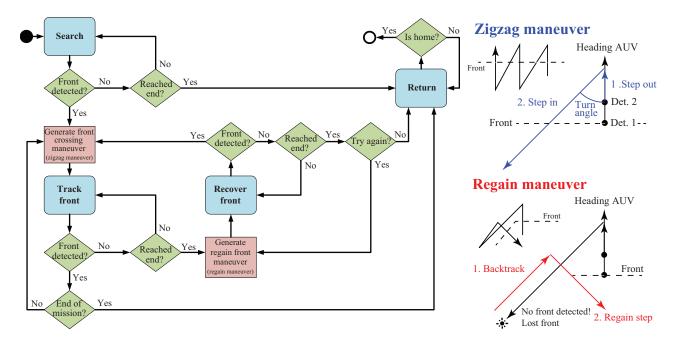


Fig. 2. The finite state machine with the associated states, actions, and decisions, as well as an illustration of the zigzag and regain maneuvers.

138	c) If $s_{front} \neq \tilde{s}_{front}$ is True: Increase the detection counter.
139	Else: Reset the detection counter.
140	d) If the detection counter is greater than 2:
141	i) Change state to <i>Track front</i> .
142	ii) Switch the current state s_{front} : inside \rightarrow outside or outside \rightarrow inside.
143	iii) Generate new zigzag maneuver.
144	2) Track front mode - Cross-front using zigzag maneuver.
145	a) Perform step a)-c) of Search mode.
146	b) If detection counter is greater than 2:
147	i) Increase the front crossing counter.
148	ii) Switch the current state s_{front} : inside \rightarrow outside or outside \rightarrow inside.
149	iii) Generate new zigzag maneuver.
150	c) If maximum number of front crossings is reached: Switch state to Return using
151	predefined home waypoint.
152	d) If end is reached without a front is detected: Switch state to Recover front.

153	i) Change state to <i>Recover front</i> .
154	ii) Generate new regain maneuver.
155	3) Recover front mode - Attempt to recover front using regain maneuver.
156	a) Perform step a)-c) of Search mode.
157	b) If the detection counter is greater than 2:
158	i) Change state to <i>Track front</i> .
159	ii) Generate new zigzag maneuver.
160	c) If end is reached without front detection:
161	i) If regain attempts is less than maximum: Generate new regain maneuver.
162	ii) Else: Switch state to Return using predefined home waypoint.

163 Description of Zigzag and Regain Maneuvers:

Fig. 2 also presents insight into the generation of the zigzag maneuver, as well as the maneuver used to regain the front if it is lost. The predetermined zigzag patterns assume that the front is approximately orthogonal to the heading when detection (which has to be verified twice [Det. 1 and Det. 2 in the figure]) was made. This requires using a *turn angle* (set to 45°) and a predetermined *step in* which allows the pattern to steer towards another crossing of the front. A *step out* is also taken initially after detection in order to sample well within the different sides of the front.

Finally, the *Regain* maneuver is an important addition to the tracking behavior, allowing the 171 AUV to regain the front; this is important if the curvature of the front is high, as a fixed angle 172 zigzag maneuver will have difficulties following the front boundary. The maneuver is initiated 173 after a completed zigzag maneuver without any front detection. The essence is to back-track to 174 the place where the font was last detected and approach this point with a different heading. The 175 maneuver uses a pre-defined *backtrack*-step and *regain*-step (shown in red text in Fig. 2) set to 176 be half the length of the step in distance. Execution of the backtrack-step is followed by a 90° 177 turn angle, before the regain-step, resulting in a path that ends where the last front crossing was 178 detected. This maneuver can also be run multiple times if necessary (see e.g. Fig 3). Continuing 179 to run the maneuver will cause the AUV to loiter around in a square pattern (due to the regain 180 back-tracking step) until a certain number of repetitions have been made (five in our case), at 181

¹⁸² which point the AUV returns to a pre-determined home location.

183 B. Sea trial in the Trondheimsfjord

Before the deployment in the Arctic, a field trial was made in the Trondheimsfjord using 184 a simulated front. Measurements were fed to the sensor from a simulation engine that relied 185 on calculating the radial distance from the AUV to a pre-determined coordinate. If the AUV 186 was within a given radius, shown in grey in Fig. 3, measurements would change to reflect that 187 the AUV had crossed the front. The Zigzag maneuver would then be triggered and the AUV 188 would turn to cross the front once again. The simulated front was programmed to be circular 189 with a radius of about 250 m and hence had a large curvature. The assumption that the heading 190 would be orthogonal does not hold for a fixed turn angle and thus the AUV would eventually 191 lose the front and would have to execute the Regain maneuver. The resulting AUV track from 192 the trial is shown as a black line in Fig. 3. As indicated, the AUV successfully navigated the 193 simulated front, executing both the Zigzag and Regain maneuvers. The trial also shows that the 194 AUV successfully regained track of the front after a failed detection attempt. Due to the small 195 scale of the front, the 45° turns of the AUV can be seen clearly. 196

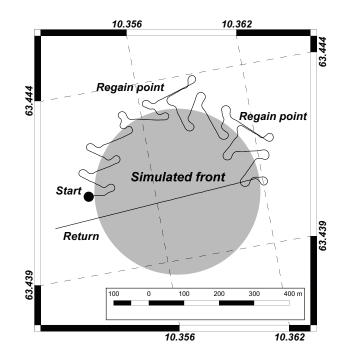


Fig. 3. The AUV path in the Trondheimsfjord. The AUV tracks a simulated front using sensor spoofing, crossing the front, losing contact, and regaining contact using the *Regain* maneuver.

IV. EXPERIMENTAL RESULTS

198 A. Experiment Setup: LAUV Harald

The AUV platform used in our experiments was a 2.4 m long, 100-m-rated OceanScan Light 199 AUV (LAUV) [27], capable of more than 24 hours of in-water operation, see Fig. 4. The payload 200 included a 16 Hz SeaBird FastCAT 49 CTD (conductivity, temperature, and depth) sensor; 201 a WetLabs Triple-Measurement Meter EcoPuck, measuring color-dissolved organic matter at 202 370/460nm, chlorophyll a fluorescence at 470/695nm, and optical backscatter; and an Aanderaa 203 4831 optode for measuring dissolved oxygen concentration. The accuracy of the CTD instrument 204 is ± 0.0003 S m⁻¹ (conductivity) and ± 0.002 °C (temperature); accuracy of the dissolved oxygen 205 sensor is less than 8μ mol L⁻¹. The chlorophyll a fluorescence sensor has a sensitivity of 0.016μ g 206 L^{-1} and the CDOM of 0.184 ppb. 207

The front tracking algorithm was hosted on a NVIDIA Jetson TX1 multicore single board computer, running through the autonomous agent architecture T-REX (Teleo-Reactive EXecutive),

197

and executing tasks continuously as sensing and control data are fed to it from the integrated control system, DUNE [28], [29]¹. Details of T-REX are beyond the scope of this work; readers are referred to [30], [31] for more information.

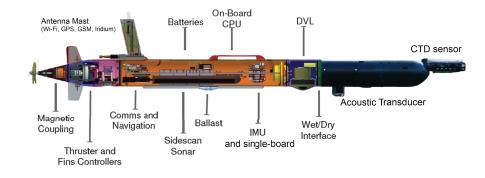


Fig. 4. The AUV platform used in the experiment. Sensors and hardware components are shown as integrated.

213 B. Experiment at Svalbard - The Nansen Legacy Campaign

The front experiment was conducted north of Svalbard at approximately 82°N (see Fig. 1(b)), 214 during cruise KH2018709 on board R/V Kronprins Haakon (14-24 September 2018). The AUV 215 deployment was augmented by hydrographic profiles collected using the ship's CTD system (Sea-216 Bird Scientific, SBE 911plus). An overview of the experiment setup and involved systems are 217 shown in Fig. 1(a). The ship maintained a safe distance from the AUV operation area during the 218 survey to avoid interference and collision (the AUV is not visible by eye from the ship bridge). 219 This distance dictated that only satellite communications could be used, making autonomy and 220 adaptation strictly necessary. 221

The AUV was deployed on the 18th of September 2018, and three missions were made between 07:00 (start first mission) and 16:34 UTC (end last mission), with duration of 238 minutes, 112 minutes, and 56 minutes, respectively. Deployment and recovery were made using a light work boat. Of the three missions, the front was detected and tracked only in the first mission (238 minutes). The latter missions failed to find the front as the front had receded toward the ice edge, moving 4 km northward; this eventually prompted the commanded abort of the AUV less

¹http://lsts.pt/toolchain

than 500 m from the ice edge in the last mission (56 minutes). Consequently, only the results from the first mission are presented. Here, the AUV successfully tracked the front along the predefined $T_{isotherm} = 1.5^{\circ}$ C contour line. A detailed description of the mission parameters used by the FSM are given in Table I.

Parameter	Value	Comment
Step out	300 m	Travel after detection
Step in	2500 m	Distance to cross front
Turn angle	45°	Turn angle
Detection temperature	$1.5 \pm 0.5^\circ C$	$T_{isotherm} \pm T_{hysteresis}$
Detection depth	0.5-8.0 m	The detection depth interval
Yoyo depths	0-90 m	Min and Max depths
Number of crossings	6	Front crossings

TABLE I		
MISSION PARAMETERS		

1) AUV Data: Fig. 5 shows the AUV path superimposed on a horizontal spatial interpolation 232 of the temperature measurements from 0.5-8 m depth. The interpolation clearly marks the front, 233 which can be seen as the dashed line. A total of six crossings made during the first mission 234 were conducted over 4 hours (23 km), with a mean speed of 1.6 m/s. As designed, the AUV 235 samples well within each water mass on both sides of the front, while tracking the orientation 236 of the front towards the northeast. The jagged AUV path is due to drift of the inertial navigation 237 system, which resets upon surfacing (hence the jump). A clear drift towards the southwest can 238 be indicative of currents influencing the AUV in this direction. The AUV covered the water 239 column from 0-90 m throughout the mission, except for the last northeast-southwest segment 240 where the AUV returned back on the surface (as a safety precaution to keep contact). 241

By combining all the measurements from the AUV profiles (yoyo from 0-100 m), a volumetric representation of the front structure can be rendered. This can be seen in Fig. 6, which shows the estimated 3D temperature distribution obtained using kriging [32] (were correlation parameters was adjusted for stratification, increasing the horizontal correlation) together with the AUV path. The temperature data were spatially averaged and discretized to a 50x50x50 cell volume grid. As the measurements were collected over a period of 4 hours, the volumetric interpolation will not

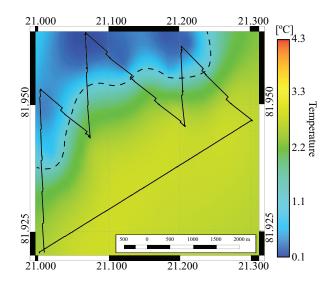


Fig. 5. Map of the AUV path overlaid on the 2 m cut plane from the interpolated (kriged) measurements, with the front indicated (dashed line).

²⁴⁸ be accurate as a temporal snapshot. This time distortion can be improved by using Lagrangian ²⁴⁹ measurement strategies or more complex correlation functions, see e.g. [33]. The front is shown ²⁵⁰ as an iso-surface with its outline marked for clarification. The front has an intrusion of warm ²⁵¹ water at 40 m that arises from the frontal dynamics; this is further discussed in detail in [7].

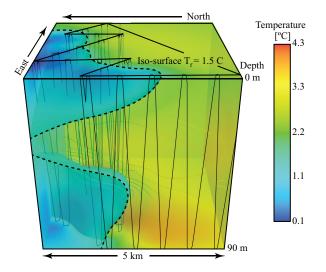


Fig. 6. The volumetric temperature distribution as estimated from kriging (see e.g. [34]). The vertical saw-tooth pattern of the AUV path can be seen in black, crossing back and forth the warm/cold temperature front. Side view.

2) Comparison with Ship Data: The AUV and shipboard measurements are not fully contem-252 poraneous or co-located, hence a detailed comparison or validation was not attempted. However, 253 the near-surface measurements of the frontal region and the Atlantic and polar origin waters from 254 the thermosalinograph of the ship are consistent with the AUV measurements (Fig. 7). Using data 255 points collected within 2 km and 2 hours by the two platforms, average temperature measured by 256 the AUV between 3 and 5 m depth, and from the thermosalinograph (water intake depth 4 m), 257 root-mean-squared difference was 0.07°C, increasing from 0.05°C on the warm side to 0.09°C 258 on the cold side. A comparison with the CTD profiles is not attempted because the distance to 259 the nearest CTD station was larger than 2 km. When combined, the measurements obtained from 260 the two platforms allow for a larger spatial coverage. The AUV therefore substantially augments 261 the sampling by performing an efficient track using the adaptive sampling method. 262

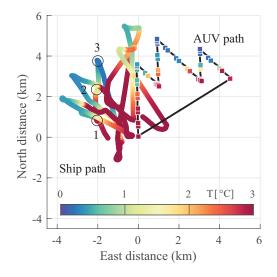


Fig. 7. Near-surface temperature measured by the ship's thermosalinograph (color coded) along the ship's track together with the AUV path (black). CTD stations 1 to 3 across the front are marked with black circles. Horizontal distance is referenced to the AUV deployment location.

Without continuous profiling by an undulating towfish, an underwater glider, or an AUV as employed here, typical shipboard sampling would be conducted using the ship's CTD. We exemplify this using three CTD profiles collected from the ship (Fig. 7, stars). They are spaced out on the warm side, at the front, and the cold side of the boundary. The profiles down to 500 m depth are shown in Fig. 8. Although a higher spatial resolution can be achieved by collecting shallower profiles, such measurements would compromise the much-needed deep hydrographic measurements during a cruise. An AUV thus provides high horizontal resolution in parallel to ship's operations, making this an attractive technology for process studies.

3) Observations: The profiles (Fig. 8) reflect the complex structure of water masses found 271 north of Svalbard that result from the different routes Atlantic Water (AW) can take to reach this 272 area. The different routes determine the water mass modification that AW experiences along its 273 path. On the warm side of the front, between 50 and 250 m depth, we see nearly pure AW with 274 a temperature above 3°C and salinity above 34.9, implying negligible dilution with surrounding 275 waters. This suggests that these waters likely followed a short route across the Yermak Plateau 276 (shown in Fig. 1(b)). Modified AW with a reduced temperature is found below this depth on 277 both sides of the front. Relatively diluted waters could represent water that has been modified 278 along a longer route, e.g., around or across the Yermak Plateau. When warm AW meets and 279 melts sea ice, a colder and relatively less saline water mass is formed in the surface layer. The 280 melt water reduces the density in the upper layer, increases the vertical stratification, and can 281 protect sea ice from further melting. The relatively fresh layer in the upper 50 m on the warm 282 side can be associated with melting. The most pronounced stratification on the cold side of the 283 front with lowest upper layer salinity, however, is the cold and fresh Polar Water that has been 284 formed in the Arctic Ocean. The gradient from warm and saline AW toward the cold and less 285 saline waters is notable, and generates the front sampled. The vertical structure of the gradients 286 of temperature, salinity and density between the warm and cold sides of the front is, however, 287 different. While the lateral density gradient is pronounced only in the upper 30 m, temperature 288 and salinity profiles show large differences reaching as deep as 200 m. 289

The AUV, with its high spatial resolution, allowed for a detailed mapping of the across-front 290 structure. Figure 9 presents cross-front sections of the different parameters, with the warm side 291 located at distance y < 0 km and the cold side at y > 0 km, where y is the cross-front distance 292 in km. Details about the construction of the sections can be found in [7]. The AUV mission 293 was designed to cross the surface temperature front, but it also criss-crossed a deeper front. The 294 front separates the warm and saline AW coming directly from Fram Strait from the colder and 295 fresher Polar Water that has been formed in the Arctic Ocean. The front has a distinct signature in 296 Chlorophyll a fluorescence, with larger concentrations on the warm side. A subsurface maximum 297 of Chlorophyll a was found at about 30 m depth on the cold side. Northeasterly winds during 298

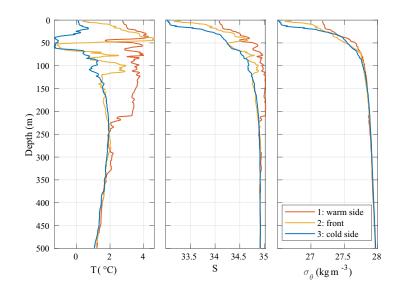


Fig. 8. Temperature, salinity, and potential density anomaly profiles collected at stations 1 to 3, in the warm side, at the front, and in the cold side of the front.

the cruise forced the warm and dense mixed layer to downwell beneath the cold, lighter side, supported by current observations from a ship-mounted acoustic Doppler current profiler. At around 40 m depth, a warm intrusion is observed in Fig. 9, in close agreement with the intrusion captured at 50 m in the shipborne CTD profile at the front.

```
303
```

V. DISCUSSION

For a subsumption architecture such as the one presented here (see Fig. 2), the sensory 304 information couples directly with action selection, with a limited state-based model of the 305 environment. There is no need for a complex internalized model as there is no planning or 306 deliberation involved, but only actions in response to sensor values. Constructing a successful 307 sampling approach rests upon the practicability of decomposing the problem into different sets of 308 behaviors that represent certain action-response pairs that are triggered by the incoming data. The 309 frontal structure studied is simple enough that this is possible. However, handling off-nominal 310 endogenous or exogenous conditions, including the presence of multiple fronts, can result in 311 complications as the potential growth of the number of states and the associated intelligent 312 switching needed to resolve conflicts or priorities. Introduction of a more elaborate environmental 313 model would alleviate the complexity of the FSM. A natural extension would then be a more 314

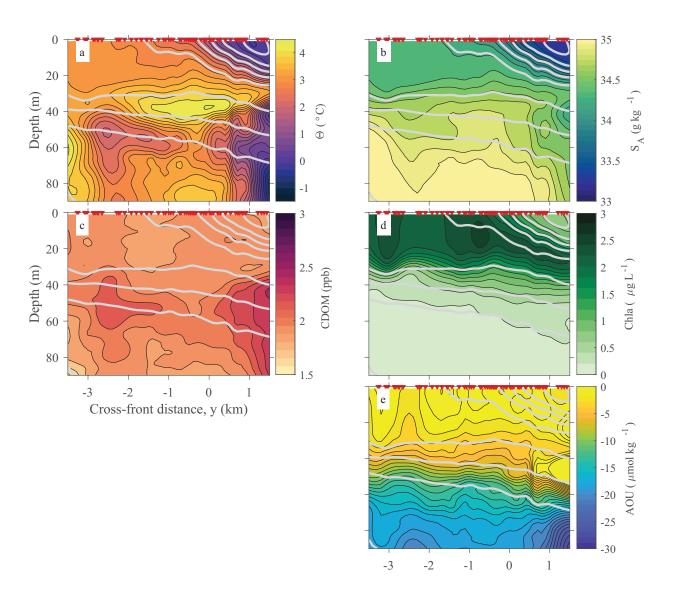


Fig. 9. Cross-front sections based on AUV data of a) Conservative Temperature (Θ), b) Absolute Salinity (S_A), c) CDOM (ppb), d) chlorophyll *a* fluorescence and e) apparent oxygen utilisation (AOU). The red triangles are the location of the AUV profiles. The grey lines are isopycnals (every 0.1 kg m⁻³)

deliberative hybrid system, where a state-based reasoning would only occur at the highest level, leaving more detailed planning to more information-theoretic approaches (see [35]).

In the FSM implementation shown here, prior information of the front dynamics, such as the estimated temperatures on each sides of the front, is necessary to define the hysteresis used for detection. This can be deducted automatically by having the AUV search for a gradient with a predefined magnitude instead, which would simplify the configuration of the FSM. One could also foresee an extension with more elaborate behaviors for keeping track of the frontal boundary (e.g. [21]). Additionally, an online estimate of the front dynamics could be deduced from either current measurements or more elaborate statistical models (see e.g. [36]).

The spatial scales associated with oceanographic processes (fronts, eddies etc.) are large (10-324 100 km) compared to the typical distance covered by short AUV missions. While a single 325 transect can cover about 30 km in a 6 hour mission at 1.5 m/s, multiple crossings of a feature 326 will be limited in scale. Targeted measurements such as those conducted here, however, return 327 high resolution observations which cannot be collected otherwise. A description of evolution of 328 processes in response to external forcing in the upper layers of the ocean, however, will require 329 coverage by a network of AUVs or other platforms, such as ships, to capture the larger scales, 330 see e.g. [25]. 331

Doing repeated autonomous missions with AUVs in the Arctic will also prompt the need for more elaborate instrumentation and acoustic navigation infrastructure, mission optimization related to energy conservation and front coverage, risk management, and fault tolerant control. One could also foresee including detection and avoidance of sea ice to ensure a more safe operation.

337

VI. CONCLUSION

Operating in harsh environments without human supervision and limited communications is 338 especially important for regions such as the Arctic. We have, in this paper, presented a method for 339 autonomous adaptive sampling of frontal features, based on a subsumption architecture. The state-340 based sampling agent uses a classical zigzag maneuver to track the front, as well as featuring 341 the capability to recover tracking if the gradient signature is lost. During field experiments 342 north of Svalbard (82°N), the agent successfully detected and tracked along an Arctic frontal 343 feature close to the sea ice edge for several kilometers, making a total of six crossings while 344 performing vertical profiles in the water column. Measurements yield a detailed volumetric 345 estimate of the frontal feature with high resolution along the frontal zone, which augments ship-346 based sampling that was run in parallel. The sampling agent revealed cross-frontal structures, 347 both horizontally and vertically, of the complex water mass compositions found north of Svalbard 348

that would not be possible through ship-based sampling only. A deeper understanding of small 349 scale processes at frontal systems will improve ecosystem models in these highly productive 350 areas. The AUV is central here as it provides substantial value through efficient and targeted 351 sampling of dynamic processes. Linking the high-resolution observations to external forcing, 352 however, requires multi-platform sampling strategies (e.g. the use of multiple AUVs, gliders 353 and a ship that operate simultaneously). The results demonstrate a framework for conducting 354 interdisciplinary oceanographic data collection in the Arctic, combining new technologies to 355 achieve a detailed picture of water-column processes. 356

357

VII. ACKNOWLEDGEMENT

The authors would like to thank the R/V Kronprins Haakon crew for the field support. This work was part of the *Nansen Legacy Program*², project number # 276730, and AMOS ³, Center of Excellence, project number # 223254.

361

References

- [1] I. M. Belkin, P. C. Cornillon, and K. Sherman, "Fronts in large marine ecosystems," Progress in Oceanography, vol. 81,
- no. 1, pp. 223 236, 2009, comparative Marine Ecosystem Structure and Function: Descriptors and Characteristics.
 [Online]. Available: http://www.sciencedirect.com/science/article/pii/S0079661109000330
- J. Le Fèvre, "Aspects of the biology of frontal systems," ser. Advances in Marine Biology, J. Blaxter and A. Southward,
 Eds. Academic Press, 1987, vol. 23, pp. 163 299. [Online]. Available: http://www.sciencedirect.com/science/article/
 pii/S0065288108601091
- [3] C. B. Woodson and S. Y. Litvin, "Ocean fronts drive marine fishery production and biogeochemical cycling," *Proceedings* of the National Academy of Sciences, vol. 112, no. 6, pp. 1710–1715, 2015.
- [4] P. Munk, C. J. Fox, L. J. Bolle, C. J. G. Van Damme, P. Fossum, and G. Kraus, "Spawning of north sea fishes
 linked to hydrographic features," *Fisheries Oceanography*, vol. 18, no. 6, pp. 458–469, 2009. [Online]. Available:
 https://onlinelibrary.wiley.com/doi/abs/10.1111/j.1365-2419.2009.00525.x
- [5] E. D'Asaro, C. Lee, L. Rainville, R. Harcourt, and L. Thomas, "Enhanced turbulence and energy dissipation at ocean
 fronts," *science*, vol. 332, no. 6027, pp. 318–322, 2011.
- [6] J. R. Taylor and R. Ferrari, "Ocean fronts trigger high latitude phytoplankton blooms," Geophysical Research Letters,
- vol. 38, no. 23, 2011. [Online]. Available: https://agupubs.onlinelibrary.wiley.com/doi/abs/10.1029/2011GL049312
- [7] Z. Koenig, I. Fer, E. Kolås, T. O. Fossum, P. Norgren, and M. Ludvigsen, "Observations of turbulence at a near-surface temperature front in the arctic ocean," *Journal of Geophysical Research: Oceans*, vol. 125, no. 4, p. e2019JC015526,

²https://www.arvenetternansen.com ³https://www.ntnu.edu/amos

- 2020, e2019JC015526 10.1029/2019JC015526. [Online]. Available: https://agupubs.onlinelibrary.wiley.com/doi/abs/10.
 1029/2019JC015526
- [8] N. E. Leonard, D. A. Paley, F. Lekien, R. Sepulchre, D. M. Fratantoni, and R. E. Davis, "Collective Motion, Sensor Networks, and Ocean Sampling," *Proceedings of the IEEE*, vol. 95, no. 1, pp. 48–74, jan 2007.
- 383 [9] M. L. Seto, Marine robot autonomy. Springer, 2013.
- [10] S. Frolov, B. Garau, and J. Bellingham, "Can we do better than the grid survey: Optimal synoptic surveys in presence of variable uncertainty and decorrelation scales," *Journal of Geophysical Research: Oceans*, vol. 119, pp. 5071–5090, 2014.
- [11] J. Das, F. Py, J. B. J. Harvey, J. P. Ryan, A. Gellene, R. Graham, D. A. Caron, K. Rajan, and G. S. Sukhatme,
 "Data-driven robotic sampling for marine ecosystem monitoring," *The International Journal of Robotics Research*,
 vol. 34, no. 12, pp. 1435–1452, 2015. [Online]. Available: http://ijr.sagepub.com/content/34/12/1435.full
- Y. Zhang, M. A. Godin, J. G. Bellingham, and J. P. Ryan, "Using an autonomous underwater vehicle to track a coastal
 upwelling front," *IEEE Journal of Oceanic Engineering*, vol. 37, no. 3, pp. 338–347, July 2012.
- [13] R. N. Smith, F. Py, P. Cooksey, G. Sukhatme, and K. Rajan, "Adaptive Path Planning for Tracking Ocean Fronts with an
 Autonomous Underwater Vehicle," in *International Symposium on Experimental Robotics (ISER)*, Morocco, June 2016.
- [14] T. B. Curtin, J. G. Bellingham, J. Catipovic, and D. Webb, "Autonomous oceanographic sampling networks,"
 Oceanography, vol. 6, no. 3, pp. 86–94, 1993. [Online]. Available: http://www.jstor.org/stable/43924649
- [15] S. R. Ramp, R. E. Davis, N. E. Leonard, I. Shulman, Y. Chao, A. R. Robinson, J. Marsden, P. F. Lermusiaux, D. M.
 Fratantoni, J. D. Paduan, F. P. Chavez, F. L. Bahr, S. Liang, W. Leslie, and Z. Li, "Preparing to predict: The Second
 Autonomous Ocean Sampling Network (AOSN-II) experiment in the Monterey Bay," *Deep-Sea Research Part II: Topical Studies in Oceanography*, vol. 56, no. 3-5, pp. 68–86, 2009.
- E. Carmack, I. Polyakov, L. Padman, I. Fer, E. Hunke, J. Hutchings, J. Jackson, D. Kelley, R. Kwok, C. Layton *et al.*,
 "Toward quantifying the increasing role of oceanic heat in sea ice loss in the new arctic," *Bulletin of the American Meteorological Society*, vol. 96, no. 12, pp. 2079–2105, 2015.
- 402 [17] Y. Zhang, B. Kieft, C. Rueda, T. O'Reilly, J. Ryan, T. Maughan, C. Wahl, and F. Chavez, "Autonomous front tracking by
 403 a wave glider," in *OCEANS 2016 MTS/IEEE Monterey*, Sep. 2016, pp. 1–4.
- [18] C. J. Cannell and D. J. Stilwell, "A comparison of two approaches for adaptive sampling of environmental processes using
 autonomous underwater vehicles," in *Proceedings of OCEANS 2005 MTS/IEEE*, Sep. 2005, pp. 1514–1521 Vol. 2.
- [19] S. Petillo, H. Schmidt, P. Lermusiaux, D. Yoerger, and A. Balasuriya, "Autonomous amp; amp; adaptive oceanographic
 front tracking on board autonomous underwater vehicles," in *OCEANS 2015 Genova*, May 2015, pp. 1–10.
- [20] J. Pinto, R. Mendes, J. C. B. da Silva, J. M. Dias, and J. B. de Sousa, "Multiple autonomous vehicles applied to plume
 detection and tracking," in 2018 OCEANS MTS/IEEE Kobe Techno-Oceans (OTO), May 2018, pp. 1–6.
- [21] N. A. Cruz and A. C. Matos, "Reactive auv motion for thermocline tracking," in *OCEANS'10 IEEE SYDNEY*, May 2010,
 pp. 1–6.
- Y. Zhang, J. G. Bellingham, M. Godin, J. P. Ryan, R. S. McEwen, B. Kieft, B. Hobson, and T. Hoover, "Thermocline
 tracking based on peak-gradient detection by an autonomous underwater vehicle," in *OCEANS 2010 MTS/IEEE SEATTLE*,
 Sep. 2010, pp. 1–4.
- [23] S. Petillo, A. Balasuriya, and H. Schmidt, "Autonomous adaptive environmental assessment and feature tracking via
 autonomous underwater vehicles," in *OCEANS'10 IEEE SYDNEY*, May 2010, pp. 1–9.
- 417 [24] N. A. Cruz and A. C. Matos, "Autonomous tracking of a horizontal boundary," in 2014 Oceans-St. John's. IEEE, 2014,
- 418 pp. 1–6.

- [25] A. Branch, M. M. Flexas, B. Claus, A. F. Thompson, Y. Zhang, E. B. Clark, S. Chien, D. M. Fratantoni, J. C. Kinsey,
 B. Hobson, B. Kieft, and F. P. Chavez, "Front delineation and tracking with multiple underwater vehicles," *Journal of Field*
- 421 Robotics, vol. 36, no. 3, pp. 568–586, 2019. [Online]. Available: https://onlinelibrary.wiley.com/doi/abs/10.1002/rob.21853
- R. Brooks, "A robust layered control system for a mobile robot," *IEEE journal on robotics and automation*, vol. 2, no. 1,
 pp. 14–23, 1986.
- [27] A. Sousa, L. Madureira, J. Coelho, J. Pinto, J. Pereira, J. Sousa, and P. Dias, "LAUV: The man-portable autonomous
 underwater vehicle," in *Navigation, Guidance and Control of Underwater Vehicles*, 2012, vol. 3, pp. 268–274.
- [28] J. Pinto, P. Calado, J. Braga, P. Dias, R. Martins, E. Marques, and Sousa, "Implementation of a control architecture for
 networked vehicle systems," *IFAC Proceedings Volumes*, vol. 45, no. 5, pp. 100–105, 2012.
- I. Pinto, P. S. Diasand, R. Martins, J. Fortuna, E. Marques, and J. Sousa, "The LSTS toolchain for networked vehicle
 systems," in *MTS/IEEE Oceans*. IEEE, 2013, pp. 1–9.
- [30] F. Py, K. Rajan, and C. McGann, "A Systematic Agent Framework for Situated Autonomous Systems," in *9th International Conf. on Autonomous Agents and Multiagent Systems (AAMAS)*, Toronto, Canada, May 2010.
- [31] K. Rajan and F. Py, "T-REX: Partitioned Inference for AUV Mission Control," in *Further Advances in Unmanned Marine Vehicles*, G. N. Roberts and R. Sutton, Eds. The Institution of Engineering and Technology (IET), August 2012.
- [32] N. A. C. Cressie and C. K. Wikle, *Statistics for Spatio-Temporal Data.*, ser. Wiley Series in Probability and Statistics.
 Wiley, 2011.
- [33] R. Graham, F. Py, J. Das, D. Lucas, T. Maughan, and K. Rajan, "Exploring Space-Time Tradeoffs in Autonomous
 Sampling for Marine Robotics," in *Experimental Robotics*, 2013, vol. 88, pp. 819–839. [Online]. Available: http://link.springer.com/10.1007/978-3-319-00065-7
- [34] J. Eidsvik, T. Mukerji, and D. Bhattacharjya, *Value of Information in the Earth Sciences: Integrating Spatial Modeling and Decision Analysis*. Cambridge: Cambridge University Press, 008 2015. [Online]. Available: https:
 //www.cambridge.org/core/books/value-of-information-in-the-earth-sciences/61119AB2F707D557E49E00BF9FD6FE39
- 442 [35] T. O. Fossum, J. Eidsvik, I. Ellingsen, M. O. Alver, G. M. Fragoso, G. Johnsen, R. Mendes, M. Ludvigsen, and K. Rajan,
- "Information-driven robotic sampling in the coastal ocean," *Journal of Field Robotics*, vol. 35, no. 7, pp. 1101–1121,
 2018. [Online]. Available: https://onlinelibrary.wiley.com/doi/abs/10.1002/rob.21805
- 445 [36] T. O. Fossum, C. Travelletti, J. Eidsvik, D. Ginsbourger, and K. Rajan, "Learning excursion sets of vector-valued gaussian
- random fields for autonomous ocean sampling," 2020.