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Assessment of wave runup and wave rundown based on observed long-term wave conditions

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Abstract. The article presents an analytical method that can be applied to provide first estimates of wave runup (RU) and wave rundown (RD) on shorelines and coastal structures based on observed long-term wave statistics and is supplementary to [12] who based their results on long-term wind statistics. Some recently published wave RU and wave RD formulae are used, together with joint statistics of significant wave height and spectral peak period from the Northern North Sea. Results are exemplified for the average statistical properties of wave RU and wave RD in terms of the expected values and the standard deviations, as well as values estimated from 1-, 10- and 100-years return period contour lines.

1. Introduction
When waves approach a coastline, they usually break and run up on the coast, be it a structure or a beach. For low laying land or areas particularly exposed and vulnerable to waves this can represent critical situations leading to e.g. flooding, coastal erosion, and damage to coastal infrastructure such as breakwaters, seawalls, artificial reefs and sand barriers. Some recent works addressing these issues are those of [1], [2], [3] and [4]. The research attention has increased in the recent years driven by the focus on climate change and consequently possible sea level rise and more extreme weather. For shorelines and coastal structures, it is essential to reliably assess the maximum wave runup (RU) and the maximum wave rundown (RD) for safe design and cost-efficient coastal protections.

The wave RU height and the wave RD height are defined as the vertical difference between the highest and the lowest point, respectively, and the still water level. Two components contribute to the wave RU; the wave set-up and the swash. Here the wave set-up is the mean surface elevation level with reference to the mean surface elevation in deep water caused by the radiation stress (see [5]); the swash oscillates from the wave set-up corresponding to the interception between the water and the shoreline (or structure); see [1] for further details. Common design formulae for wave RU and wave RD use a 2% exceedance value of the RU maxima at the toe of the shoreline (or structure), \( R_{2%} \), as well as a 2% exceedance value of the RD maxima, \( R_{d2%} \). Most of these design formulae are given in terms of the surf (Iribarren) number, which is defined in terms of the significant wave height \( H_s \) in deep water, the spectral peak period \( T_p \) in deep water, and the slope of the shoreline (or the structure). Two other commonly used wave RU formulae are those of [6] and [7], where the first formula recently has been applied by [8]. However, these general formulae can not be treated by the present analytical approach.
Results from some recent studies will be advocated here and described briefly in Section 2; from [9], [1], [2], [3] and [4]. The latter reference gives a review and summary of wave RU formulae. Some of these recent wave RU formulae were applied by [10], [11] and [12] using long-term variation of wave and wind conditions. Here [10] and [11] based their results on long-term wave statistics, applying the wave RU formulae by [1] and the wave RU and wave RD formulae by [2], respectively. Further, [11] also demonstrated how 100-years return period values of wave RU and wave RD can be calculated based on a joint distribution of significant wave height and spectral surf parameter. The present article is supplementary to [12] who based their results on long-term wind statistics, applying eight wave RU formulae and one wave RD formula. The main purpose here is to demonstrate how similar and some additional results of wave RU and wave RD can be derived by using long-term wave statistics.

In this article it is demonstrated how long-term wave statistics in deep water can be applied to provide estimates of the wave RU and wave RD on shorelines or coastal structures. Results are obtained by adopting published formulae; eight wave RU and two RD formulae, together with a joint distribution of $H_s$ and $T_p$ from the Northern North Sea. A procedure of estimating the wave RU and wave RD from 1-, 10- and 100-years return period contour lines and the corresponding values of $H_s$, $T_p$ and the spectral surf parameter are also provided. The present analytical method should provide a useful tool that can be used for initial estimation of wave RU and wave RD in for instance early feasibility studies and risk analysis, or for estimations in the field.

2. Background

Extreme RU and RD events have been extensively investigated over the last decades in small- and large-scale laboratory studies as well as in field campaigns. The results are given in terms of empirical formulae representing estimates of the 2% exceedance values of RU and RD maxima. The formulae used in this article are adopted from some recently published works and are summarized in the following.

Two RU and one RD formulae were presented by [2]; the following two formulae for maximum wave RU

$$R_{2\%} = 1.165 H_s \xi_p^{0.77}$$  \hspace{1cm} (1)

$$R_{2\%} = (0.39 + 0.795 \xi_p) H_s$$  \hspace{1cm} (2)

and for maximum wave RD

$$R_{2\%} = (0.21 - 0.44 \xi_p) H_s$$  \hspace{1cm} (3)

Here $\xi_p$ is the spectral surf parameter defined as

$$\xi_p = m\left(\frac{H_s}{\frac{g}{2\pi T_p^2}}\right)^{\frac{1}{2}}$$  \hspace{1cm} (4)

where $m = \tan \alpha$ is the slope with an angle $\alpha$ with the horizontal, and $g$ is the acceleration due to gravity. The wave RU formulae are based on the formulations by [13] and [14], with new empirical coefficients fitted to the equations; Equations (1) and (2) are modified versions of the original formulae by [13] and [14], respectively. Equations (1) to (3) were based on large-scale laboratory experiments on a prototype-scale sand barrier with a median grain size diameter $d_{so} = 0.42$ mm, and are valid for $m$ in the range $0.088 - 0.154$ and $\xi_p$ in the range $1 - 2.9$. Furthermore, [2] suggested that the linear model in $\xi_p$ (Eq. (2)) is the most easily applicable to RU data for the range of beach slopes they considered (see [2] for further details).

The following wave RD formulae based on data from a large-scale laboratory experiment was presented by [15]
The experiment was carried out using both the uniform slopes 1/6 and 1/12 and composite slopes with steepness 1/3 for the lower slope and 1/6 for the upper slope. For the composite slopes an appropriate average slope was calculated. The slope surface was covered by an asphalt concrete layer. This formula is based on data in the spectral wave steepness range $H_s / ((g / 2\pi) / T_p^2)$ of 0.001-0.031, and $0.5 < \xi_p < 2.5$.

The following wave RU model for $\xi_p < 0.6$, based on data from small-scale laboratory experiments for the three slopes 1/50, 1/30 and 1/20, and a sand seabed with $d_{50} = 0.7$ mm was proposed by [1]

$$R_{2s4} = 4 m^{0.3} H_s \xi_p$$

Assessment of wave RU predictions on beaches on the South-East Australian coast considering 11 existing empirical formulae was performed by [4]. The data used in this comparison represent sand bottom with $d_{50}$ in the range 0.25 – 0.5 mm, $m$ in the range 0.02 – 0.16, $\xi_p$ in the range 0.32 – 1.65; the wave conditions in the assessment covered a range with average values of approximately $H_s = 1.5$ m and $T_p = 8.9$ s, which are slightly below the typical mean conditions of the region of $H_s = 1.6$ m and $T_p = 9.5$ s. However, the proposed models in Equations (9) and (10) are also considered to be valid for higher wave conditions because the formulae they are based on, cover a range up to $H_s = 5$ m and $T_p = 15$ s. Among the 11 formulae [4] considered (in addition to Equations (9) and (10)) were those by [16] and [9]. The wave RU formula by [16] is given as

$$R_{2s4} = (0.2 + 0.83\xi_p) H_s$$

and based on data from measurements at Duck, North Carolina, USA.

The wave RU formula by [9] is given as

$$R_{2s4} = (0.58 m + 0.53\xi_p) H_s + 0.45$$

and based on field measurements on the south coast of Algarve, Portugal representing the European Atlantic coast. The data they used were based on manually selected wave RU maxima representing $m$ in the range 0.04 - 0.15, $\xi_p$ in the range 0.3 – 2.9, sand seabed with $d_{50} = 0.50$ mm.

The formulae by [4] are

$$M1 R_{2s4} = 0.99 H_s \xi_p$$

$$M2 R_{2s4} = (0.16 + 0.92\xi_p) H_s$$

and are based on the best fit to the 11 formulae they considered, where the one in Equation (9) is forced through the origin (i.e. with the origin in $\xi_p = 0$ for $R_{2s4} / H_s$ versus $\xi_p$).

Moreover, [3] proposed a parameterization of wave RU from field measurements on gravel beaches and numerical calculation given as

$$R_{2s4} = C m^{0.5} T_p H_s ; \quad C = 0.33$$

The field data they used were obtained during a 2-year period including storm conditions with $H_s$ in the range 1 – 8 m, $m$ in the range 0.05 – 0.20, $\xi_p$ in the range 0.2 -1.9, $d_{50}$ in the range 2 -50 mm (see [3] for more details).

As a summary, Equations (1) to (3) and (5) to (10) can be represented as

$$R_s = a H_s + b H_s \xi_p + d$$

$$R_d = -0.1 \xi_p^{2.21} H_s$$

(5)
Table 1. Wave RU and RD formulae according to Equation (12) for models 1 to 9 and Equation (11) for model 10.

<table>
<thead>
<tr>
<th>Model nr</th>
<th>Abbreviation</th>
<th>Ref.</th>
<th>a</th>
<th>b</th>
<th>c</th>
<th>d</th>
<th>C</th>
</tr>
</thead>
<tbody>
<tr>
<td>1, RU</td>
<td>Bl 1</td>
<td>[2]</td>
<td>0</td>
<td>1.165</td>
<td>0.77</td>
<td>0</td>
<td>-</td>
</tr>
<tr>
<td>2, RU</td>
<td>Bl 2</td>
<td></td>
<td>0.39</td>
<td>0.795</td>
<td>1</td>
<td>0</td>
<td>-</td>
</tr>
<tr>
<td>3, RD</td>
<td>Bl d</td>
<td></td>
<td>0.21</td>
<td>-0.44</td>
<td>1</td>
<td>0</td>
<td>-</td>
</tr>
<tr>
<td>4, RD</td>
<td>Sc</td>
<td>[15]</td>
<td>0</td>
<td>-0.1</td>
<td>2.21</td>
<td>0</td>
<td>-</td>
</tr>
<tr>
<td>5, RU</td>
<td>Pe</td>
<td>[1]</td>
<td>0</td>
<td>4m0.3</td>
<td>1</td>
<td>0</td>
<td>-</td>
</tr>
<tr>
<td>6, RU</td>
<td>Ho</td>
<td>[16]</td>
<td>0.2</td>
<td>0.83</td>
<td>1</td>
<td>0</td>
<td>-</td>
</tr>
<tr>
<td>7, RU</td>
<td>Vo</td>
<td>[9]</td>
<td>0.58m</td>
<td>0.53</td>
<td>1</td>
<td>0.45</td>
<td>-</td>
</tr>
<tr>
<td>8, RU</td>
<td>At1</td>
<td></td>
<td>0</td>
<td>0.99</td>
<td>1</td>
<td>0</td>
<td>-</td>
</tr>
<tr>
<td>9, RU</td>
<td>At2</td>
<td></td>
<td>0.16</td>
<td>0.92</td>
<td>1</td>
<td>0</td>
<td>-</td>
</tr>
<tr>
<td>10, RU</td>
<td>Po</td>
<td>[3]</td>
<td>-</td>
<td>-</td>
<td>-</td>
<td>-</td>
<td>0.3</td>
</tr>
</tbody>
</table>

where $R_2$ represents $R_{2\%}$ for wave RU and $R_{d2\%}$ for wave RD, and the coefficients $a$, $b$, $c$, $d$ are given in Table 1.

As referred to in Section 1, the wave RU formulae by [6] and [7] are frequently used for assessing wave RU, but can not be treated analytically. In particular, the formulations by [7] cover a wide range of conditions on different slopes on shallow and very shallow foreshores with different surface roughnesses, although some of the formulae, e.g. their Equations (5.1), (5.2), (5.5) and (5.6) have the forms of Equation (12). Alternative methods are, for instance; to apply Monte Carlo simulations, or simple methods as those of [17] or [18].

Thus, according to Equations (11) and (12), $R_2$ is given in terms of the deep water sea state parameters $H_s$ and $T_p$. Different parametric models for the joint probability density function (pdf) of $H_s$ and $T_p$ are available in the literature; see [19] for a review. The joint pdf used to exemplify the results in this article is taken from [20], given as

$$p(H_s, T_p) = p(T_p | H_s) p(H_s)$$

(13)

Here $p(H_s)$ is the marginal pdf of $H_s$, given by the following combined lognormal and Weibull distributions

$$p(H_s) = \begin{cases} 
\frac{1}{\sqrt{2\pi \kappa} H_s} \exp \left[ -\frac{(\ln H_s - \theta)^2}{2\kappa^2} \right], & H_s \leq 3.25 \text{ m} \\
\beta \frac{H_s^{\beta-1}}{\zeta^\beta} \exp \left[ -\frac{H_s}{\zeta} \right], & H_s > 3.25 \text{ m} 
\end{cases}$$

(14)

where $\theta = 0.801$, $\kappa^2 = 0.371$ are the mean value and the variance, respectively, of $\ln H_s$ and $\zeta = 2.713$, $\beta = 1.531$ are the Weibull parameters.

$p(T_p | H_s)$ is the conditional pdf of $T_p$ given $H_s$, given by the lognormal pdf

$$p(T_p | H_s) = \frac{1}{\sqrt{2\pi \sigma T_p}} \exp \left[ -\frac{(\ln T_p - \mu)^2}{2\sigma^2} \right]$$

(15)

where $\mu$ and $\sigma^2$ are the mean value and the variance, respectively, of $\ln T_p$ given as
\[ \mu = a_i + a_2 H_s^{a_i}, \quad (a_i, a_2, a_3) = (1.780, 0.288, 0.474) \] (16)

\[ \sigma^2 = b_i + b_2 e^{b_i H_s}, \quad (b_i, b_2, b_3) = (0.001, 0.097, -0.255) \] (17)

This joint pdf of \( H_s \) and \( T_p \) was obtained as a best fit to 29 years of wave data from the Northern North Sea representing wind waves, swell waves and combinations; see [20] for more details. The unit of \( H_s \) in Equations (16) and (17) is in meter.

It should be noted that the wave \( RU \) and wave \( RD \) formulae used here are based on data from other locations than the wave data. Therefore, in a proper analysis the wave data should be from the same location as the wave \( RU \) and wave \( RD \) models are based on. However, here the present wave data are used since the main purpose is to demonstrate the application of the method.

3. Statistical properties of \( RU \) and \( RD \)

The statistical properties of \( R_2 \), from which the statistical properties of \( R_{2\%} \) and \( R_{42\%} \) follow, are derived from the joint pdf of \( H_s \) and \( T_p \), giving the joint pdf of \( R_2 \) and \( H_s \). This is achieved from Equations (11) and (12) by a change of variables from \((H_s, T_p)\) to \((R_2, H_s)\), yielding

\[ p(H_s, R_2) = p(R_2 \mid H_s) p(H_s) \] (18)

From Equations (11), (4), (12) and (13) it is noticed that this change of variables only affects \( p(T_p \mid H_s) \).

First, consider Equations (4) and (12) from which

\[ T_p = \left[ \frac{R_2 - aH_s - d}{bm^c \left( \frac{g}{2\pi} \right)^{\frac{c}{2}} H_s^{\frac{1-c}{2}}} \right]^\frac{1}{c} \] (19)

This gives the Jacobian

\[ \frac{\partial T_p}{\partial R_2} = \left[ \frac{1}{c} (R_2 - aH_s - d)^{\frac{1-c}{c}} \right] \left[ bm^c \left( \frac{g}{2\pi} \right)^{\frac{c}{2}} H_s^{\frac{1-c}{2}} \right]^{-\frac{1}{c}} \] (20)

Thus, by using this Jacobian and that \( p(R_2 \mid H_s) = p(T_p \mid H_s) \mid \frac{\partial T_p}{\partial R_2} \mid \), the following lognormal pdf of \( R_2 \) given \( H_s \) is obtained

\[ p(R \mid H_s) = \frac{1}{\sqrt{2\pi} \sigma_R R} \exp \left[ -\frac{(lnR - \mu_R)^2}{2\sigma_R^2} \right] \] (21)

where

\[ R = R_2 - aH_s - d \] (22)

and \( \mu_R \) and \( \sigma_R^2 \) are the conditional mean value and conditional variance, respectively, of \( ln R \), given as

\[ \mu_R = c\mu + ln[bm^c \left( \frac{g}{2\pi} \right)^{\frac{c}{2}} H_s^{\frac{1-c}{2}}] \] (23)
\[ \sigma_R^2 = (c \sigma)^2 \]  

(24)

where \( \mu \) and \( \sigma^2 \) are given in Equations (16) and (17), respectively.

The conditional cumulative distribution function \((cdf)\) of \( R \) given \( H_s \) is given by the standard Gaussian \( cdf \ \Phi \) as (\([21]\)\)

\[ P(R \mid H_s) = \Phi \left[ \frac{\ln R - \mu_R}{\sigma_R} \right] \]

(25)

\[ \Phi(\nu) = \frac{1}{\sqrt{2\pi}} \int_{-\infty}^{\nu} e^{-\nu^2/2} \, d\nu \]  

(26)

Then, the expected value and the standard deviation of \( R \) given \( H_s \) are given in Equations (27) and (28), respectively, as (\([21]\)\)

\[ E[R \mid H_s] = \exp(\mu_R + \frac{1}{2} \sigma_R^2) \]

(27)

\[ \sigma[R \mid H_s] = \left( (e^{\sigma_R^2} - 1) \exp(2\mu_R + \sigma_R^2) \right)^{1/2} \]

(28)

Thus, from Equation (22) it follows that

\[ E[R_2 \mid H_s] = E[R \mid H_s] + aH_s + d \]

(29)

\[ \sigma[R_2 \mid H_s] = \sigma[R \mid H_s] \]

(30)

Similarly, by taking Equation (11) as \( R_2 = C m^{0.5} H_s T_p \), a change of variables from \((H_s, T_p)\) to \((H_s, R_2)\) only affects \( p(T_p \mid H_s) \). Then, by using the Jacobian \( |\partial T_p / \partial R_2| = 1/(C m^{0.5} H_s) \), this yields a lognormal pdf of \( R_2 \) given \( H_s \) in the form

\[ p(R_2 \mid H_s) = \frac{1}{\sqrt{2\pi} \sigma_{R_2}} \exp \left[ -\frac{(\ln R_2 - \mu_{R_2})^2}{2\sigma_{R_2}^2} \right] \]  

(31)

where the conditional mean value and the conditional variance, respectively, of \( \ln R_2 \), are given as

\[ \mu_{R_2} = \mu + \ln(C m^{0.5} H_s) \]

(32)

\[ \sigma_{R_2}^2 = \sigma^2 \]  

(33)

where \( \mu \) and \( \sigma^2 \) are given in Equations (16) and (17), respectively.

4. Example of results

4.1. Conditional statistical values of \( RU \) and \( RD \) within a sea state

Here the average statistical properties of \( RU \) and \( RD \) expressed in terms of the expected values and the standard deviations are provided. Examples of results are given for:

- Significant wave height in deep water, \( H_s = 3 \text{ m} \)
- Slope, \( m = 1/10 \)

These parameters are within the validity range of the models (except for model 5 which is valid for \( m = 1/50, 1/30, 1/20 \)), and is used here to serve the purpose of demonstrating the application of the method.
Table 2. Conditional expected values of \( RU \) for models 1, 2, 5 to 10 and \( RD \) for models 3, 4, for \( H_s = 3 \text{ m} \) and slope \( m = 1/10 \) (column 3); the corresponding expected value (EV) ±1 standard deviation (SD) (column 4); see also Table 1.

<table>
<thead>
<tr>
<th>Model nr</th>
<th>Abbreviation</th>
<th>EV (m)</th>
<th>EV ±1 SD (m)</th>
</tr>
</thead>
<tbody>
<tr>
<td>1, RU</td>
<td>Bl 1, Eq. (1)</td>
<td>2.7</td>
<td>2.2, 3.1</td>
</tr>
<tr>
<td>2, RU</td>
<td>Bl 2, Eq. (2)</td>
<td>2.9</td>
<td>2.5, 3.2</td>
</tr>
<tr>
<td>3, RD</td>
<td>Bl d, Eq. (3)</td>
<td>-0.31</td>
<td>-0.51, -0.10</td>
</tr>
<tr>
<td>4, RD</td>
<td>Sc, Eq. (5)</td>
<td>-0.15</td>
<td>-0.23, -0.07</td>
</tr>
<tr>
<td>5, RU</td>
<td>Pe, Eq. (6)</td>
<td>4.3</td>
<td>3.3, 5.2</td>
</tr>
<tr>
<td>6, RU</td>
<td>Ho, Eq. (7)</td>
<td>2.4</td>
<td>2.0, 2.8</td>
</tr>
<tr>
<td>7, RU</td>
<td>Vo, Eq. (8)</td>
<td>1.8</td>
<td>1.5, 2.0</td>
</tr>
<tr>
<td>8, RU</td>
<td>At 1, Eq. (9)</td>
<td>2.1</td>
<td>1.7, 2.6</td>
</tr>
<tr>
<td>9, RU</td>
<td>At 2, Eq. (10)</td>
<td>2.4</td>
<td>2.0, 2.9</td>
</tr>
<tr>
<td>10, RU</td>
<td>Po, Eq. (11)</td>
<td>3.1</td>
<td>2.4, 3.8</td>
</tr>
</tbody>
</table>

For these conditions the expected value (EV) and the expected value plus and minus one standard deviation (SD), i.e. EV±1 SD, of \( RU \) and \( RD \) can be calculated. The results are given in Table 2 showing that EV for wave \( RU \) ranges from 1.8m to 4.3m, with the lowest value based on model 7 and the highest value based on model 5; the coefficient of variation (SD/EV) varies from about 0.1 to about 0.2 with the lowest value for model 2 and the highest for model 5. Furthermore, EV for wave \( RD \) are -0.31m and -0.15m for models 3 and 4, respectively, while SD/EV is about 0.05 and 0.5, respectively. Based on the results for EV±1 SD it appears that there is overlap between the values obtained by the two wave \( RD \) models 3 and 4, suggesting that the results are consistent. For the values obtained by the eight wave \( RU \) models there is overlap between models 1, 2, 6, 8, 9, 10; between models 7 and 8; between models 5 and 10. It should also be noted that the EV ±1 SD value of model 7 coincides with the EV – 1 SD values of models 6 and 9. Overall, this suggests that the results from models 1, 2, 6, 8, 9, 10 are consistent. It should be recalled that these results are not general, but valid for \( H_s = 3 \text{ m} \) and the slope \( m = 1/10 \) (which is outside the validity range of model 5 and thus the result for this model is less reliable). Moreover, these results are associated with the present \( p(H_s, T_p) \), and thus other joint pdfs from the same locations as the wave \( RU \) and wave \( RD \) models are based might give different results.

4.2. Estimation of \( RU \) and \( RD \) from n-years return period contour lines of \( p(H_s, T_p) \)

N-years return period values of \( RU \) and \( RD \) can be estimated from the n-years return period contour lines of \( p(H_s, R_2) \) in Equation (19). However, an alternative is to estimate these values of \( RU \) and \( RD \) from the n-years return period contour lines of \( p(H_s, T_p) \) in Equations (13) to (17).
Figure 1. 1-year, 10-years and 100-years contour lines of $H_s (m)$ (vertical axes) and $T_p (s)$ (horizontal axes) from the inner to outer contours, respectively, with the tangent lines corresponding to: (a) RU model 1; (b) RU model 2; (c) RD model 3; (d) RD model 4; (e) RU model 5; (f) RU model 6; (g) RU model 7; (h) RU model 8; (i) RU model 9; (j) RU model 10. See also Tables 1 and 3.
Figure 1 depicts the 1-year, 10-years and 100-years return period contour lines by assuming a sea state duration of 3 hours, represented by the inner to the outer contours, respectively, obtained using the IFORM method described in [22]. This method is based on transformation of the joint cdf of $H_s$ and $T_p$ to the standard Gaussian cdf, from which the return period contour lines are determined as circles.

First, consider Equation (11) which by solving for $H_s$ gives

$$H_s = \frac{R_2}{C m^{0.5} T_p^{-1}}$$

(34)

For a given value of $R_2$, this is a curve in the $(H_s, T_p)$ plane. The values of $R_2$ which implies that these curves will have tangent points with the 1-year, the 10-years and the 100-years contours can be determined by iteration, and thus the corresponding tangent points. The results are shown graphically in Figure 1j by the three curves giving tangent points to the 1-, 10-, 100-years contours for $R_2 = 16.2$ m, 20.5 m and 24.7 m respectively, corresponding to the 1-, 10-, 100-years return period values of $R_2$ based on model 10 and the slope $m = 1/10$. However, how realistic these large values of $R_2$ are, is questionable; more discussion is given subsequently. It is observed that these values of $R_2$ are governed by $H_s$ in the sense that the tangent points are located close to the maximum values of $H_s$ along the respective return period contours. The corresponding coordinates of the tangent points for $H_s, T_p$ are given in Table 3, together with the corresponding values of $\xi_p$ according to Equation (4).

Second, consider Equation (12), which for a given value of $R_2$ represents a curve in the $(H_s, T_p)$ plane. Also in this case the values of $R_2$ corresponding to the tangent points to the 1-, 10-, 100-years contours are found by iteration. The results are shown graphically in Figures 1a, b, e – i for RU and in Figures 1c, d for RD by the three curves giving tangent points to the 1-, 10-, 100-years contours. The values of $R_2$ and the corresponding values of $H_s, T_p$ and $\xi_p$ for each of the RU and RD models are given in Table 3 for the slope $m = 1/10$. It is observed that these values of $R_2$ corresponding to RU are governed by $H_s$. However, the values of $R_2$ corresponding RD (Figures 1c, d) are related to tangent points given by combined lower $H_s$ and higher $T_p$ values.

Table 3. Extreme values of RU ($R_2$) for models 1, 2, 5 to 10 and RD ($R_2$) for models 3, 4 corresponding to the 1-10, 100-years contours of $(H_s, T_p)$ and corresponding values of $H_s, T_p, \xi_p$ (for the slope $m = 1/10$), from left to right for each variable, respectively; see also Table 1.

<table>
<thead>
<tr>
<th>Model nr</th>
<th>$R_2$ (m)</th>
<th>$H_s$ (m)</th>
<th>$T_p$ (s)</th>
<th>$\xi_p$</th>
</tr>
</thead>
<tbody>
<tr>
<td>1, RU</td>
<td>8.0, 9.3, 10.5</td>
<td>10.2, 12.0, 13.9</td>
<td>15.2, 16.3, 17.0</td>
<td>0.60, 0.59, 0.57</td>
</tr>
<tr>
<td>2, RU</td>
<td>8.8, 10.3, 11.7</td>
<td>10.3, 12.3, 14.1</td>
<td>15.0, 15.9, 16.8</td>
<td>0.58, 0.56, 0.56</td>
</tr>
<tr>
<td>3, RD</td>
<td>-1.3, -1.5, -1.8</td>
<td>3.6, 3.4, 3.6</td>
<td>19.4, 22.1, 24.4</td>
<td>1.3, 1.5, 1.6</td>
</tr>
<tr>
<td>4, RD</td>
<td>-0.72, -1.0, -1.4</td>
<td>1.7, 1.6, 1.3</td>
<td>20.1, 23.3, 26.5</td>
<td>2.0, 2.3, 2.9</td>
</tr>
<tr>
<td>5, RU</td>
<td>12.2, 14.2, 16.0</td>
<td>9.8, 11.8, 13.7</td>
<td>15.6, 16.5, 17.2</td>
<td>0.62, 0.60, 0.58</td>
</tr>
<tr>
<td>6, RU</td>
<td>7.1, 8.3, 9.4</td>
<td>10.2, 12.2, 13.9</td>
<td>15.2, 16.1, 17.0</td>
<td>0.60, 0.58, 0.57</td>
</tr>
<tr>
<td>7, RU</td>
<td>4.3, 4.9, 5.5</td>
<td>10.0, 12.0, 13.7</td>
<td>15.4, 16.3, 17.2</td>
<td>0.61, 0.59, 0.58</td>
</tr>
<tr>
<td>8, RU</td>
<td>6.0, 7.0, 7.9</td>
<td>9.8, 11.8, 13.7</td>
<td>15.6, 16.5, 17.2</td>
<td>0.62, 0.60, 0.58</td>
</tr>
<tr>
<td>9, RU</td>
<td>7.2, 8.4, 9.5</td>
<td>10.2, 12.0, 13.9</td>
<td>15.2, 16.3, 17.0</td>
<td>0.60, 0.59, 0.57</td>
</tr>
<tr>
<td>10, RU</td>
<td>16.2, 20.5, 24.7</td>
<td>10.3, 12.2, 13.9</td>
<td>15.0, 16.1, 17.0</td>
<td>0.58, 0.58, 0.57</td>
</tr>
</tbody>
</table>
It should be noted that some of the estimated RU and RD values are obtained by exceeding the surf parameter validity bounds stated in Section 2, but is used here with the purpose of demonstrating the application of the method. This is the case for models 1 and 2 for all predictions, since the models are valid for $\xi_p$ in the range 1 to 2.9; for model 4 for the 100-years return period value since $\xi_p = 2.9 > 2.5$; for model 5 for the 1-year return period value since $\xi_p = 0.62 > 0.6$. Furthermore, it is also noticed that the RU values from models 5 and 10 are large compared to the other model predictions. As referred to earlier, model 5 is valid for smaller slopes than $m = 1/10$ used here, and thus the results for this model are less reliable. If a smaller slope (i.e. in the range 1/50-1/20) is applied for the same wave conditions, it is expected to predict lower RU values (see Equation (6)). The model 10 results should be reasonable since Equation (11) is used within its range of validity. However, a firm statement on the validity of the return period results is not possible without comparison with measurements.

A similar procedure can be applied for other joint distributions of $H_s, T_p$ as well as for other $n$ – years return period contour lines to obtain the wave RU and wave RD values and the corresponding values of $H_s, T_p$ and $\xi_p$. However, as mentioned in Section 2, the wave data should be from the same location as the wave RU and wave RD models are based on.

4.3. Discussion

This section provides some comments on this approach versus a procedure which applies more complete computational demanding methods without using empirical formulae. For assessing the maximum wave RU and wave RD heights, common practice would be to start with available data on joint statistics of $H_s$ and $T_p$ (or other characteristic wave periods); preferably within directional sectors at a nearby location offshore (in deep water). The next step would be to apply a wave simulation model that includes effects of dissipation due to bottom friction and wave breaking, and thereby to obtain the joint statistics of $H_s$ and $T_p$ at the coastal (shallow water) site; then finally to use this as input for computing the maximum wave RU and wave RD heights. In general this practice would also include sites exposed to sea states with combined wind waves and swell waves from different directions. In this article an alternative is presented providing an analytical method which can be used to estimate maximum wave RU and wave RD heights from given deep water values of $H_s$ and $T_p$, exemplified by including results using a joint distribution of $H_s$ and $T_p$ representing wind waves, swell waves and combinations.

The transition from deep water to the coastal site is assumed to be smooth, i.e. by neglecting wave energy dissipation effects over changing bed conditions with varying intermediate and shallow water depths. As a result, several effects affecting the estimated wave RU and wave RD heights are neglected, e.g.: that the wave field is inhomogeneous; from where the waves are coming and the location of the assessment site; return flows from dissipation effects which in turn will affect the local wave conditions. However, the approach enables analytical estimates of the maximum wave RU and wave RD heights, which are appropriate for making quick estimates. Then, these estimates can be used to compare with more complete computationally demanding methods. Under field conditions such an easily accessible and simple tool might also be useful as there is usually limited time and access to computational resources. Although the presented results are based on specific wave RU and wave RD formulae and a joint $(H_s, T_p)$ distribution from another location than the wave RU and wave RD models are based on, the method can also be applied for other wave RU and wave RD formulae that can be treated analytically, joint distributions of sea state wave parameters, or for a given deep water wave spectrum including directional spreading effects. However, in such cases numerical calculations are most probably required.

It is important, however, to assess the accuracy of this simple approach versus common practice, which is only possible to quantify by comparing with such methods over a wide parameter range, also including
a sensitivity analysis of the results regarding the assumptions considered, but this is beyond the scope of this article.

5. Conclusions
An analytical method that can be applied to provide estimates of wave RU and wave RD on shorelines and coastal structures is presented. Results are achieved by adopting eight wave RU formulae and two wave RD formulae together with observed long-term wave statistics from deep water in the Northern North Sea. Expected values and variances of the 2% exceedance values of the wave RU maxima and the wave RD maxima are calculated. The main purpose is to demonstrate how long-term wave statistics can be used to obtain first estimates of wave RU and wave RD.

Examples of results for typical realistic conditions are also provided by presenting the conditional expected value (EV) plus and minus one standard deviation (1 SD) of the 2% exceedance values of the wave RU maxima and the wave RD maxima within a sea state. By comparing the results from the eight wave RU models and the corresponding EV ± 1 SD, it is overlap between models 1, 2, 6, 8, 9, 10; between models 7 and 8; between models 5 and 10; as well as between the two RD models 3 and 4. Overall, this suggests that the results from the RU models 1, 2, 6, 8, 9, 10 as well as from the RD models 3, 4 are consistent. However, a firm statement of this is not possible without comparison with measurements.

A procedure of estimating the 2% exceedance values of the wave RU maxima and the wave RD maxima as well as the corresponding values of significant wave height, spectral peak period and spectral surf parameter from 1-, 10-, and 100-years return period contour lines is also demonstrated. The estimated return period values seem to be reasonable, except for the estimates based on the wave RU model 5 which is valid for smaller slopes than that used here. However, a firm statement of this also requires comparison with measurements.

Overall, the estimates of wave RU and wave RD are associated with a large amount of uncertainty which is partly because the wave data are from another location than the wave RU and wave RD models are based on, and partly because the models are based on data from different locations. However, the present analytical method is convenient due to its simplicity, but it should only be used for early estimates due to the uncertainties related to the results. The method should provide an appropriate and useful tool for initial estimates of wave RU and wave RD in for instance early feasibility studies and risk analysis related to the assessment of climate change effects in coastal vulnerability studies, or for estimation in the field.

References
[22] DNV 2014 Recommended Practice DNV- C205 Environmental Conditions and Environmental Loads. DNV GL AS.