
PostPrint

Estimation of sea state parameters by the wave buoy analogy with comparisons to third generation spectral wave models

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Abstract

This paper presents a study focused on sea state estimation along the route of an in-service container ship. The paper is concerned with the wave buoy analogy in which wave-induced motions of the ship are processed and analysed together with corresponding motion transfer function to give the directional wave spectrum exactly at the point of operation. In this study, a simple and inexpensive instrumentation of the vessel is considered, and wave spectrum estimation is based on measurements from one motion response unit mounted close to the forward perpendicular of the ship. The estimates by the wave buoy analogy are compared with two sets of results from third generation spectral wave models, with one set provided by a commercial supplier and with another set obtained from the Copernicus Climate Change Service Information. Motion measurements from a seven-days voyage across the Pacific Ocean are studied, and it is shown that the wave buoy analogy estimates wave conditions, in terms of sea state parameters, in good agreement with the reports by the sets of ocean wave hindcasts. Along with the comparisons, the paper discusses some of the inherent drawbacks of the wave buoy analogy, notably the fact that a ship acts as a low-pass filter.

Keywords:

Wave spectrum estimation, ship motions, wave buoy analogy, in-service data, spectral wave models, Copernicus Climate Change Service Information (ERA5)

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1. Introduction

During ship operations, real-time and on-site estimation of sea state parameters can complement the crew's decisions to maintain high safety and fuel efficiency, since the estimate will be useful for early detection of critical sailing situations and be a crucial input to motion control systems. Similarly, attained knowledge of the sea state along the exact route of merchant ships can assist shore-based performance analysis teams towards optimising vessel and fleet performance focused on fuel consumption and environmental footprints. Moreover, estimates of wave conditions can be used for virtual hull monitoring.

One means for the estimation of sea state parameters - in real time and at the precise geographical position of a ship - considers the ship itself as a (sailing) wave buoy. This particular method is often referred to as the *wave buoy analogy*. The estimation principle of the wave buoy analogy relies on the combination of measurements of wave-induced motions of the vessel and a linear assumption, allowing the motion measurements to be modelled theoretically using transfer functions and a wave (energy density) spectrum.

1.1. Scope, highlights, and objective

The present paper studies the wave buoy analogy when it is applied together with in-service data from a larger container ship. Specifically, the estimation of sea state parameters has been made using seven days of consecutive data obtained while the ship made an east-bound trip across the Pacific Ocean with measurements from the Sea of Japan to off Graham Island (Canada). The data is obtained from a simple and inexpensive instrumentation on the vessel, where one single motion response unit, placed in a point off the centreline and close to the forward perpendicular, provides the horizontal and vertical accelerations together with the pitching motion. The corresponding motion transfer functions have been obtained from linear strip theory calculations. As a side note, it should be mentioned that the motion measurements from the specific ship have recently been used in a study about wave spectrum estimation (Nielsen and Dietz, 2020), where the sensitivity to the vessel's advance speed was investigated.

It is an inherent concern about the encountered sea state during in-service conditions that the ground truth is never known. In this study, additional estimates of sea state parameters have been obtained from spectral wave models where two sets of results are introduced; the one set is made by a commercial provider and the other set has been generated using the Copernicus Climate

31 Change Service Information (2020). Altogether, the highlights of the study can be referred to as
32 (a) estimation of wave spectra, i.e. sea states, using in-service data obtained from a simple sensor
33 instrumentation on a container vessel, and (b) a comprehensive comparison between results of the
34 wave buoy analogy and corresponding ones produced by spectral wave models; notably the use
35 of the freely available ERA5 data (Copernicus Climate Change Service Information, 2020) is an
36 attractive novelty for the community working with sea state estimation, either from ship motion
37 measurements or other means (e.g., buoys, remote sensing, wave radar systems), since the ERA5
38 data facilitates a comparative basis.

39 Despite the capabilities of the wave buoy analogy and its usefulness for (real-time and on-site)
40 sea state estimation, as widely reported about in the literature including this paper, the current
41 article also has as an objective to discuss some of the inherent drawbacks and problems connected
42 to the wave buoy analogy.

43 *1.2. Composition*

44 The paper is organised in the following way. In the next section, Section 2, the methodology
45 is covered and herein the fundamentals of the wave buoy analogy are outlined. The section also
46 includes a short description of the parameters forming the background of the comparison between
47 the results of the wave buoy analogy and the results from the spectral wave models. Section 3
48 presents the considered ship and its data, including the origin of the data and how it has been
49 processed. In a model-based approach, like studied in this paper, the motion transfer functions of
50 the vessel are of fundamental importance, and Section 4 includes a numerical examination. The
51 results and corresponding discussions of the study are presented in Section 5, while a summary of
52 the paper and some concluding remarks are given in Section 6.

53 **2. Methodology**

54 This section explains the basis of the wave buoy analogy and how Bayesian modelling, sometimes
55 referred to as the Bayesian technique, can be used to solve the mathematical problem connected
56 to the wave buoy analogy. The section also contains a description of the wave data made available
57 from spectral wave models. Finally, the section introduces the sea state parameters that form
58 the basis of the comparison between the wave buoy analogy and the results of the spectral wave
59 models.

60 *2.1. The wave buoy analogy - Bayesian Modelling*

61 The assumptions and the equations governing the wave buoy analogy have been widely reported
 62 in the literature, e.g., Iseki and Ohtsu (2000); Tannuri et al. (2003); Nielsen (2006); Pascoal et al.
 63 (2007); Nielsen (2008a); Nielsen and Brodtkorb (2018). This section serves to indicate the most
 64 important aspects, while the details can be found in Nielsen (2006, 2008a).

65 The central assumption of the wave buoy analogy builds on linearity between waves and the
 66 wave-induced response of a vessel, and, in a frequency domain formulation, the combination with
 67 an assumption about stationarity implies the following model in which the (unknown) directional
 68 wave spectrum is $E(\omega_e, \mu)$,

$$R_{ij}(\omega_e) = \int_{-\pi}^{\pi} H_i(\omega_e, \mu + \beta) \overline{H_j(\omega_e, \mu + \beta)} E(\omega_e, \mu) d\mu + \varepsilon_{i,j} \quad (1)$$

69 Herein, $R_{ij}(\omega_e)$ is the response spectrum for responses i, j , where i and j correspond to any set of
 70 measured responses; say, the horizontal acceleration and the vertical acceleration, respectively, in
 71 a specific point in the ship coordinate system. The corresponding theoretical response spectrum
 72 is obtained as the product between the directional wave spectrum $E(\omega_e, \mu)$ and the multiplication
 73 of the set of transfer functions $H_i(\omega_e, \beta + \mu)$ and $\overline{H_j(\omega_e, \beta + \mu)}$ for responses i and j , with the
 74 bar denoting the complex conjugate. The mean wave-encounter angle is β and the direction of
 75 waves relative to this angle is μ , while the encounter frequency is ω_e . The error between the
 76 measured spectrum and the theoretically calculated one is $\varepsilon_{i,j}$, and it should be realised that the
 77 error in principle includes errors from sensors, transfer functions, and the model itself. Errors
 78 from sensors cannot (necessarily) be excluded, which is why fault detection techniques are relevant
 79 to consider in case of (real-time) on-board systems (Nielsen et al., 2012). Errors in the transfer
 80 functions and their influence on results can be investigated through sensitivity studies. Notably,
 81 the linear assumption, imposed through the very use of transfer functions, is a crucial factor. In
 82 this connection, reference can be made to the study by Mas-Soler and Simos (2019) addressing the
 83 nonlinearity related inaccuracies in motion RAOs when the wave buoy analogy is applied for wave
 84 spectrum estimation. The two types of errors from sensors and from the transfer functions are
 85 beyond the scope of this paper, and the paper therefore implicitly focuses only on the modelling
 86 error in the later section where data and results are presented, cf. Section 5.

87 It is noteworthy that Eq. (1) is usually formulated for three responses simultaneously which
88 leads to a set of nine independent equations¹. However, the directional wave spectrum is typically
89 discretised into K directions and, if the 360-degrees interval is spaced by, say, 10 deg, this results
90 in $K = 36$ unknown spectral components for any given frequency. Consequently, Eq. (1) expresses
91 a highly underdetermined equation system that cannot be solved by minimising the error ε , as the
92 corresponding least squares problem is ill-posed. Instead, *Bayesian modelling* can be applied to
93 solve the equation system. The main points of Bayesian modelling are presented below but, before
94 this, the effect of forward speed deserves special attention.

The equation system in (1) is formulated in the 'encounter domain' as the spectral densities of the wave spectrum depend on the encounter frequency ω_e , which itself is dependent on the vessel's forward speed and the wave encounter angle. As a consequence, the absolute frequency must be used instead, and it is therefore necessary to introduce the Doppler Shift. Thus, the mapping of the absolute frequency ω (of a progressive wave) to the encountered frequency ω_e is given by,

$$\omega_e = \omega - \omega^2 \frac{U}{g} \cos \mu \quad (2)$$

95 when the ship moves with speed U and at an angle μ relative to the progressive wave; g is the ac-
96 celeration of gravity. It is noteworthy that deep-water conditions have been assumed in the present
97 formulation. In practice, the inclusion of the Doppler Shift for problems related to general ship
98 motion dynamics is not without complications (Bhattacharyya, 1978; Beck et al., 1989; Lindgren
99 et al., 1999; Nielsen, 2017, 2018), but this is beyond the scope of the present paper. For wave
100 spectrum estimation, the problem has been solved, and this is indicated in the next paragraph.

101 In matrix notation, Eq. (1) can be written

$$\mathbf{b} = \mathbf{A}\mathbf{f}(\mathbf{x}) + \mathbf{w} \quad (3)$$

102 The vector function $\mathbf{f}(\mathbf{x})$ expresses the unknown values of the wave spectrum $E(\omega, \mu)$ through
103 a non-negativity constraint $\mathbf{f}(\mathbf{x}) = \exp(\mathbf{x})$, so that $\mathbf{x} = \ln E(\omega, \mu)$. It is noted that \mathbf{x} contains
104 $M \times K$ entries, where M is the number of discrete - absolute - wave frequencies, while K was
105 defined previously as the number of discrete wave heading angles. \mathbf{w} is a Gaussian white noise

¹Cross spectral analysis on three discrete-time motion signals leads to 3 real-valued spectra and 6 (= 3×2) complex-valued spectra with both real and imaginary parts.

106 sequence vector with elements $\varepsilon_{i,j}$ which are assumed to have zero mean and variance σ^2 . The
 107 vector \mathbf{b} contains the elements of $R_{ij}(\omega_e)$, and the coefficient matrix \mathbf{A} has elements according to
 108 the multiplication between products of the complex-valued transfer functions and the frequency
 109 derivatives $\frac{d\omega}{d\omega_e}$ obtained from the Doppler shift in Eq. (2). It should be realised that the total
 110 number of elements in \mathbf{b} will be $N \times P$, where N is the number of discrete encounter frequencies
 111 and P is the number of ("fundamental") equations derived from Eq. (1); in this case $P = 3 + 6 = 9$
 112 as reported previously.

113 In principle, the wave spectrum can be estimated from the minimisation of $g^2(\mathbf{x})$

$$g^2(\mathbf{x}) \equiv \|\mathbf{A}\mathbf{f}(\mathbf{x}) - \mathbf{b}\|^2 \quad (4)$$

114 where $\|\cdot\|$ represents the L_2 norm. As mentioned above, Eq. (4) represents an ill-posed problem.
 115 However, by introducing Bayesian modelling (Akaike, 1980) and thereby imposing prior constraints,
 116 the wave spectrum - in terms of \mathbf{x} - is basically estimated by minimising (Nielsen, 2008a)

$$h(\mathbf{x}) = \|\mathbf{A}\mathbf{f}(\mathbf{x}) - \mathbf{b}\|^2 + \mathbf{x}^T(u^2\mathbf{H}_1 + v^2\mathbf{H}_2)\mathbf{x} \quad (5)$$

117 where the hyperparameters u and v control the trade-off between the good fit to the data and the
 118 prior distributions set by the matrices \mathbf{H}_1 and \mathbf{H}_2 . In qualitative terms, the additional equations
 119 imposed through Eq. (5) are established by assuming the directional wave spectrum to be a smooth
 120 (piecewise continuous) function for variations with frequency and direction. Thus, the matrices,
 121 H_1 and H_2 , are organised so that they ensure that the curvature of the wave spectrum is minimised
 122 (Nielsen, 2006).

123 In the strict application of Bayesian modelling (Akaike, 1980) it is *not* Eq. (5) which is min-
 124 imised but a certain criterion - a Bayesian Information Criterion - known as ABIC. In the specific
 125 situation related to wave spectrum estimation, the criterion can be formulated (Nielsen, 2008a)

$$\begin{aligned} \text{ABIC} = P \ln h_{\min}(\mathbf{x}) - \ln |\det(u^2\mathbf{H}_1 + v^2\mathbf{H}_2)| + \\ \ln |\det(\mathbf{A}^T\mathbf{A} + u^2\mathbf{H}_1 + v^2\mathbf{H}_2)| + C \end{aligned} \quad (6)$$

126 The independent variables in ABIC are the hyperparameters and the minimisation problem is thus
 127 highly nonlinear, not to mention that ABIC depends on the solution for which $h(\mathbf{x})$ is minimum.
 128 The customary practice is to solve the convolved problem brute-force; that is, for each (manually)
 129 selected combination of the hyperparameters, Eq. (5) is minimised. Obviously, this leads to

130 a relatively high computational burden, since a range of hyperparameters must be covered for
131 both u and v , and for each combination of the two, Eq. (5) represents an equation system with
132 $N \times P$ equations from which $K \times M$ unknowns are solved. In the past, Sparano et al. (2008) and
133 Nielsen and Iseki (2010) came beyond the computational burden by suggesting to use a fixed set of
134 hyperparameters with no account for changing operational and/or environmental conditions. The
135 selection of the fixed set of hyperparameters must be made by trial and error; and should be made
136 in a situation when the *true* sea state is available, for instance using numerical simulations based
137 on a specified sea state. It is noteworthy that the resulting increase in computational efficiency
138 comes at the price of decreased accuracy from time to time. In a study, where large amount of
139 data is analysed retrospectively for the sole purpose of comparison with other means this cost is
140 considered acceptable.

141 2.2. Results from spectral wave models

142 Two sets of additional wave estimates, produced using third generation spectral wave models,
143 have been collected. The one set is from a commercial supplier mainly offering their service in
144 connection with tasks related to vessel and fleet performance analysis. The other set of result
145 has been generated using Copernicus Climate Change Service Information (2020), noting that the
146 dataset is a climate reanalysis, named ERA5 and based on ECMWF's Earth System model IFS.
147 The name ERA refers to 'ECMWF ReAnalysis', with ERA5 being the fifth major global reanalysis
148 produced by European Centre for Medium-Range Weather Forecasts (ECMWF). An overview is
149 given by ECMWF (2020).

150 As a practical remark, in this paper, the term *hindcast* is often used as a reference to the
151 estimate obtained from one of the spectral wave models.

152 Both sets of hindcasts comprise a number of integral wave parameters, cf. subsection 2.3,
153 which are available every 60 minutes on a discrete spatial grid spaced 0.5 degrees in the Earth
154 coordinates (latitude and longitude). Thus, the sets of hindcast results are (bi)linearly interpolated
155 to the exact geographic vessel positions, cf. Section 3, for the exact time stamps in Coordinated
156 Universal Time (UTC). The frequency and directional resolutions of the computations used by
157 the commercial supplier are not known to the authors², and some additional concerns about the

²The ship data originates from April 2016, which was also the time when Maersk Line collected the wave data from the commercial supplier. Maersk Line never received the raw data (i.e., the wave spectra).

158 integral wave parameters are present, as explained in subsection 2.4.2. On the other hand, a
 159 condensed introduction to the ERA5 data is given by Hersbach et al. (2020), including ECMWF
 160 (2017), while the interested reader should consult Komen et al. (1994) for a thorough description
 161 of the equations and associated mathematical modelling related to spatio-temporal development
 162 of ocean wave spectra; as used in connection with hindcasted (and forecasted) wave spectrum
 163 estimation.

164 2.3. Sea state parameters

165 The statistics of ocean wave systems can be derived from the (directional) wave spectra char-
 166 acterising the particular wave systems. However, for a large data set with many samples of wave
 167 spectra, it is not practical to compare the spectra, one by one, and, besides, the actual wave
 168 spectra are available only for the wave buoy analogy and the ERA5 data but not for the commer-
 169 cial hindcast data, as the data supplied to Maersk Line contained integral wave parameters only.
 170 Consequently, it is decided to focus the comparative study of the different estimation methods on
 171 the basis of a set of integral wave parameters, also referred to by *sea state parameters*. On the
 172 other hand, selected samples of (directional) wave spectra by the wave buoy analogy and ERA5 are
 173 studied in the discussion of results, cf. section 5, to point out the consequence(s) of the low-pass
 174 filtering characteristics of a large ship, but a detailed comparison of the actual spectra remains as
 175 a future task.

176 The two sets of hindcast data contain the following sea state parameters: the significant wave
 177 height H_s , the mean energy period T_E , and the mean wave direction D_s . In case of the wave buoy
 178 analogy, which has a directional wave spectrum $E(\omega, \mu)$ as the main output, the parameters must
 179 be calculated according to their mathematical definitions,

$$H_s = 4\sqrt{m_0} \quad (7)$$

$$T_E = 2\pi \frac{m_{-1}}{m_0} \quad (8)$$

$$\widehat{D}_s = \arctan(d/c) \quad (9)$$

180 where

$$m_n = \int_0^\infty \omega^n F(\omega) d\omega \quad n = \{-1, 0\} \quad (10)$$

$$F(\omega) = \int_{-\pi}^\pi E(\omega, \mu) d\mu \quad (11)$$

$$d = \int_{-\pi}^\pi \int_0^\infty E(\omega, \mu) \sin(\mu) d\omega d\mu \quad (12)$$

$$c = \int_{-\pi}^\pi \int_0^\infty E(\omega, \mu) \cos(\mu) d\omega d\mu \quad (13)$$

It is noteworthy that Eq. (9) yields the mean wave direction \widehat{D}_s relative to the centreline of the ship, in accordance with the definition of μ , cf. Eq. (1) in Subsection 2.1. Thus, for the wave buoy analogy, the estimate of the (absolute) mean wave direction D_s is given by

$$D_{s,WBA} = \widehat{D}_s + \Xi \quad (14)$$

181 where Ξ is the heading of the ship; 0 deg is North, 90 deg is East, etc.

182 In order to distinguish the results of the estimation methods from each other, the following
 183 notations will be used in connection with the comparisons of the sea state parameters, cf. Section
 184 5: Results of the wave buoy analogy are referred to by 'WBA', the commercial hindcast data is
 185 denoted by 'HC', and the Copernicus data is referred to by 'ERA5'.

186 2.4. Notes of concern

187 2.4.1. The wave buoy analogy

188 It is important to realise that, while the integration formulas above, i.e. Eqs. (10)-(13),
 189 are given in accordance with their exact mathematical definitions, the practical computations
 190 associated with the wave buoy analogy "suffer" from the fact that the lower and upper integration
 191 limits only reflect the used cut-off frequencies as applied in the spectral calculations. This is
 192 discussed further in later sections, but the central point to note is that results of the wave buoy
 193 analogy, per se, are compromised because of two related reasons: (1) the (necessary) use of cut-off
 194 frequencies in the spectral calculations, and (2) the fact that a ship acts as a low-pass filter. In
 195 addition to these drawbacks, other issues can negatively affect results of the wave buoy analogy, as
 196 already indicated in subsection 2.1, and the later sections of the paper elaborate on this together
 197 with the comparisons of the hindcast studies (ERA5 and HC).

As indicated in subsection 2.2, directional wave spectra from the commercial supplier are not available. Unfortunately, it is another concern that the total wave system is decomposed into partitions of integral parameters for swells and wind sea, respectively. For comparative reasons, it is therefore necessary to calculate equivalent wave parameters of the total wave system. In this case, the (total) significant wave height is obtained by

$$H_{s,HC} = \sqrt{H_{s,wind}^2 + H_{s,swell}^2} \quad (15)$$

The "total mean" relative direction $D_{s,HC}$ is approximated by introducing a weighted average considering the relative direction of the individual components (swell and wind sea) together with their energy content represented by the significant wave height. In this calculation, special care must be shown because directionality is circular - that is, defined on the interval $[0,360[$ deg, where 0 deg and 360 deg correspond to the same point - and this must be accounted for in the calculation. The weighting is according to ratios of the squared values of significant wave height and, schematically, the definition of $D_{s,HC}$ is,

$$D_{s,HC} = \frac{H_{s,wind}^2}{H_{s,tot}^2} \cdot D_{s,wind} + \frac{H_{s,swell}^2}{H_{s,tot}^2} \cdot D_{s,swell} \quad (16)$$

It is possible to approximate the "total mean" energy period $T_{E,HC}$ in a similar way, and the calculation follows from

$$T_{E,HC} = \frac{H_{s,wind}^2}{H_{s,tot}^2} \cdot T_{E,wind} + \frac{H_{s,swell}^2}{H_{s,tot}^2} \cdot T_{E,swell} \quad (17)$$

199 Later, in the comparisons of the three sets of results (WBA, ERA5, HC) it must thus be kept in
200 mind that the HC estimates of D_s and T_E , in the strict sense, are not (fully) consistent with the
201 estimates by WBA and ERA5.

202 3. Case ship and in-service data

203 The case ship is a 7,200 TEU container vessel. The vessel's main particulars are listed in Table
204 1, and plan views of the vessel are shown in Figure 1.

205 Wave-induced motions of the ship have been measured with a motion sensor (XSSENS, MTi-
206 30-6A5G4), and the recordings for the study were made on an east-bound route across the Pacific

Table 1: Main particulars of the example ship.

Length between perpendiculars, L_{pp}	332 m
Breadth moulded, B_m	42.8 m
Design draught, T_d	12.2 m
Deadweight (at T_d),	76,660 tonnes
Block coefficient, C_B	0.65

207 Ocean, see Figure 2. The motion sensor was mounted close to the bow, off the centreline, with the
 208 exact position known by the authors. The particular sensor provides drift-free 3D orientation as
 209 well as calibrated 3D acceleration, 3D rate of turn (rate gyro) and 3D earth-magnetic field data.
 210 For the purpose of sea state estimation, the vessel’s pitching motion and the horizontal and vertical
 211 accelerations have been used. The corresponding transfer functions have been calculated with an
 212 in-house linear strip theory code based on Salvesen et al. (1970), see also Section 4. In the study,
 213 the advance speed of the vessel is, as a reasonable approximation, assumed to be constant with a
 214 value of $U = 21.0$ knots at all times in the seven-days sailing period. Figure 3 shows the logged
 215 speed and, although smaller variations occur, it can be seen that it is indeed a fair assumption
 216 to use exclusively a speed of 21 knots for all 30-minutes samples forming the data stream. It is
 217 beyond the scope of the present paper, but Nielsen and Dietz (2020) discuss in detail the influence
 218 of forward speed when the wave buoy analogy is applied for wave spectrum estimation. One
 219 important finding from Nielsen and Dietz (2020) is noteworthy though; it is important to realise

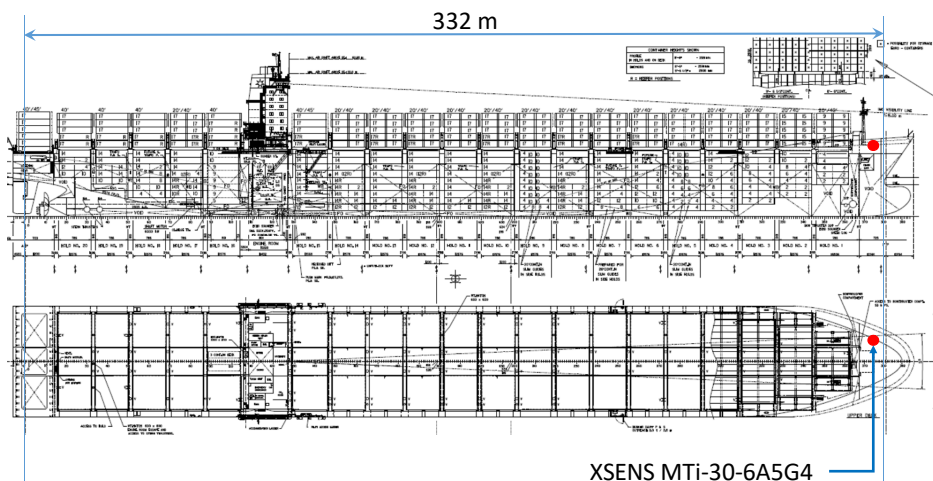


Figure 1: Plan views of the vessel with the location of the motion sensor (XSENS) indicated.

220 that the speed to use together with the wave buoy analogy must be the logged *speed-through-water*
 221 (STW). This point induces itself some further discussions (Antola et al., 2017; Hasselaar and den
 222 Hollander, 2017; Taudien and Bilén, 2018; Oikonomakis et al., 2019); how reliable is the logged
 223 STW from in-service vessels(?) As already indicated, the detailed discussions about forward-speed
 224 is out of the scope of this paper. Herein, reference is instead given to Nielsen and Dietz (2020) and,
 225 at the same time, noting that as part of another study (Nielsen et al., 2019b) it has been validated
 226 that the logged STW, cf. Figure 3, from the measurement period is reliable.

227 During the voyage, a total of 336 ($= 7 \times 48$) 30-minutes motion samples were collected. After
 228 the initial ("raw") sampling at 100 Hz, the motion recordings were resampled to 5 Hz as the
 229 vast majority of ocean waves are observed on the interval $[0-0.5]$ Hz; at least the waves being of
 230 importance to a +300 m container ship. Next, for each set of the 30-minutes motion samples,
 231 the cross power spectral density spectrum of the pairs of motion components was calculated using
 232 Welch's averaged, modified periodogram method. The resulting set of nine (cross) spectra, as used
 233 for wave spectrum estimation for a single motion sample, has been limited, i.e. low-pass filtered, to
 234 the encounter-frequency interval $[0.01-0.30]$ Hz, spaced 0.005 Hz, emphasising that no significant
 235 (wave-induced) motion occurs outside this interval. Finally, for each 30-minutes motion sample,
 236 the directional wave spectrum has been estimated, cf. Section 2.1, using a discretisation with
 237 $M = 30$ absolute wave frequencies and $K = 36$ (relative) wave directions on the intervals $\omega =$
 238 $[0.01;0.30]$ Hz and $\mu =]-180;180]$ deg, respectively; noting that, for the relative wave direction,

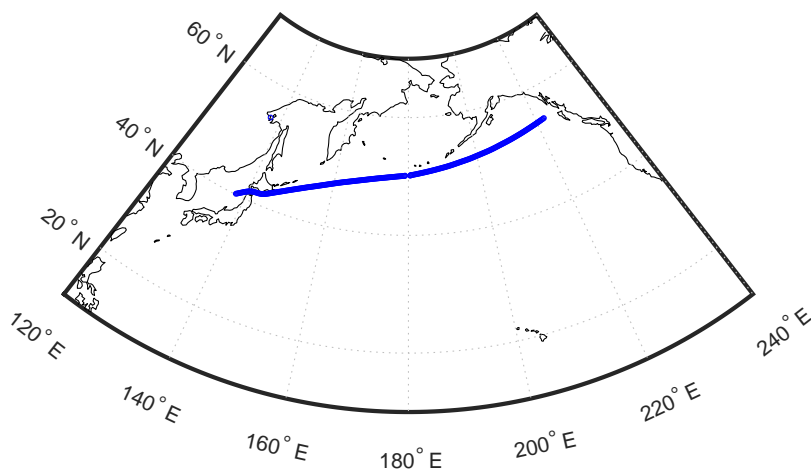


Figure 2: The analysed measurements have been recorded during an east-bound voyage across the Northern Pacific.

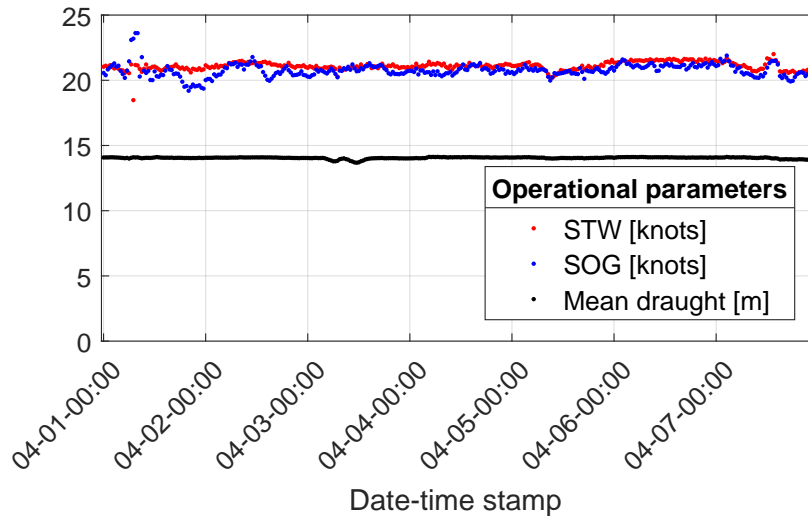


Figure 3: Measurements of advance speed during the voyage, using an acoustic Doppler current profiler for speed-through-water (STW), while GPS provides speed-over-ground (SOG). The plot includes also the logged (mean) draught amidships. The time stamps are in format 'mm-dd-hh:mm' (UTC).

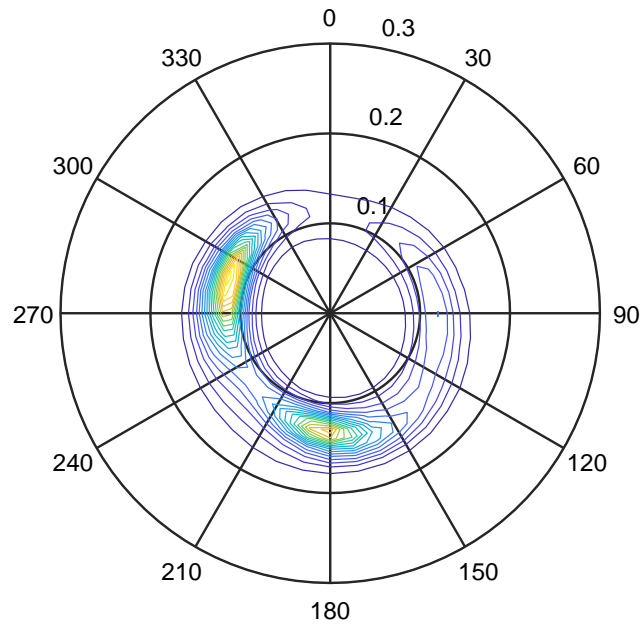


Figure 4: Directional wave spectrum estimated by the wave buoy analogy using a 30-minutes motion sample. The directional and frequency-wise energy density is indicated by colours and the directions show the directions where the energy goes to, noting that 0 deg is North, 90 deg is East, etc. In the given case, the equivalent mean wave direction is 65 deg (= where the waves come from), consistent with the sign convention of ERA5 (ECMWF, 2020).

239 equivalently wave encounter angle, $\mu = 180$ deg is head sea, $\mu = 0$ deg is following sea, while
 240 '+' and '-' are used to indicate if waves approach on the starboard or port side, respectively. An
 241 example of an estimated directional wave spectrum is shown in Figure 4. While detailed discussions
 242 are given later, in the given situation (sample 161) from 2016, April 4, UTC 08:00, the ship is on
 243 an 80 degrees course, which means that the wave system is coming in on the bow at the port
 244 side, i.e. bow-quartering sea. As a practical note, the solution is sensitive to the discretisation of
 245 the 'spectral domain' consisting of frequencies in the one dimension and wave heading angles in
 246 the other dimension; that is, the solution depends *conditionally* on the values of K and M and
 247 associated cut-off frequencies. Hereby is understood that if the discretisation is fine enough, the
 248 solution is stable and does not change (significantly) for a finer discretisation. For the specific ship
 249 and data, tests were made with $K = 18$ and $K = 72$ for selected cases leading to $K = 36$ as a good
 250 compromise (CPU time vs. accuracy). On the other hand, no sensitivity study has been made
 251 for M , and the cut-off frequencies, but based on the results in the next section, dealing with the
 252 motion transfer functions, the selected discretisation is considered appropriate.

253 4. Motion transfer functions

254 The motion transfer functions are of fundamental importance to the results of the wave buoy
 255 analogy. It is therefore useful to examine the behaviour of the transfer functions used for the
 256 ship in study. Figure 5 shows the modula of the three specific motions considered in the present
 257 study that uses pitch, vertical acceleration, and horizontal acceleration; repeating that the motion
 258 response unit is placed in a point close to the forward perpendicular, slightly off the centreline. As
 259 mentioned previously, the transfer functions have been computed with an in-house code, I-ship,
 260 based on the linear strip theory formulation by Salvesen et al. (1970).

261 In Figure 5, the modula of the transfer functions are displayed for heading angles 0-330 deg,
 262 spaced with 30 deg. In a linear theory, the assumption about rigid body motions³ means that the
 263 local lateral motion in an arbitrary point will be a (linear) coupling of sway, roll, and yaw, and
 264 thus the lateral motion is a combination of asymmetric motion components exclusively which, in
 265 turn, implies that the *modulus* of the local lateral motion is symmetric with respect to incoming

³Herein, the coordinate system is a standard right-handed with surge in the forward direction of the vessel, sway to port side, and heave upwards.

266 waves (port side vs. starboard side). On the other hand, the local vertical motion in any point will
 267 be a coupling of heave (symmetric), roll (asymmetric), and pitch (symmetric) which means that

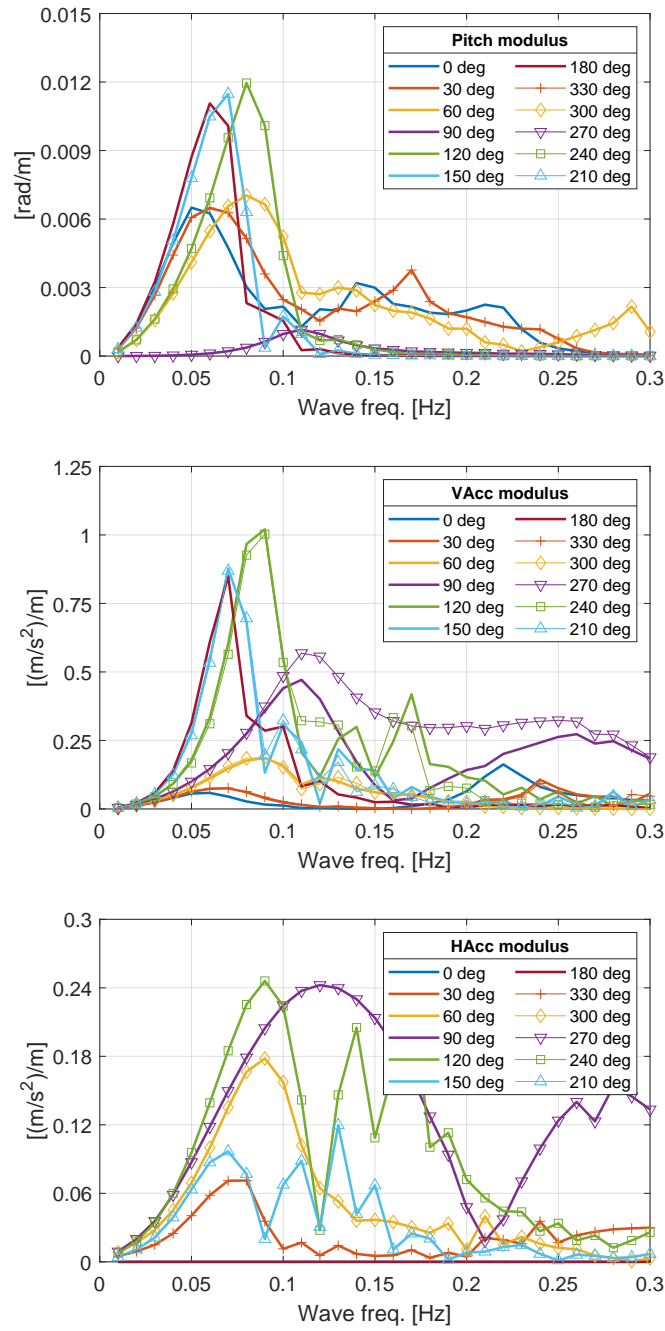


Figure 5: Modula of the three motion transfer functions (pitch, vertical acceleration, horizontal acceleration). Note, 180 deg is head sea and 0 deg is following sea.

268 the vertical motion is a combination of both symmetric motions and an asymmetric motion. In
269 turn, this implies that the modulus of the local vertical motion is itself asymmetric. Summarising,
270 the *modula* of all angular motions (roll, pitch, yaw) and all lateral local (translational) motions
271 and derivations thereof, such as the horizontal acceleration in an arbitrary but specific point, are
272 symmetric with respect to the direction of the incoming waves. However, vertical local motions
273 and derivations thereof, such as the vertical acceleration, in a point off the centreline, will be
274 asymmetric in both the argument *and* the modulus, as reflected by the middle plot in Figure 5.
275 This makes the use of the vertical acceleration (off the centreline, close to FP) advantageous, at
276 least in theory, because of the ability to distinguish between port and starboard incoming waves
277 not only by the argument but also by the modulus of the transfer function; emphasising that
278 (local) motions and other types of responses, such as wave-induced stresses (Nielsen et al., 2011;
279 Chen et al., 2019), measured in a point exactly on the centreline can be asymmetric *only* in the
280 argument.

281 Overall, it can be seen from Figure 5 that the entire set of transfer functions, considering all
282 three responses, should be useful for 'sensing' of waves on the frequency interval [0.03-0.20] Hz,
283 corresponding to waves with a period from about 30 s down to about 5 s. However, it is noteworthy
284 that cases of following sea to stern-quartering sea, i.e. $\beta \approx 0 - 45$ deg (including incoming waves
285 on either side of the vessel), do generally not impose large motions, in relative terms, which means
286 that estimation and corresponding integral wave parameters in those cases will be of a larger
287 uncertainty (Montazeri et al., 2015; de Souza, 2019) compared to estimations obtained when the
288 incoming waves approach with a mean heading $\beta \approx 45 - 180$ deg from either side of the vessel.

289 As a final remark, due to the importance of the transfer functions in connection with the wave
290 buoy analogy, it should be relevant to study the sensitivity to uncertainties in input parameters
291 such as the loading condition. However, as already discussed in Section 2, this task is left as a future
292 exercise, and the results presented in the next section are produced by taking the transfer functions
293 to be *perfect*. In any future sensitivity studies, a number of existing works are noteworthy, e.g.
294 Tannuri et al. (2003); Montazeri et al. (2016); Nielsen et al. (2018); Mas-Soler and Simos (2019);
295 Nielsen and Dietz (2020).

296 **5. Results and discussions**

297 *5.1. Comparisons of absolute values of sea state parameters*

298 As explained in Section 2, sea state parameters, equivalently integral wave parameters, have
299 been derived from the directional wave spectrum of the wave buoy analogy, and corresponding
300 estimates have also been collected from two sets of hindcast data. Comparisons of all the obtained
301 estimates are presented in Figure 6. In the plots, each point represents the result of a 30-minutes
302 period, and from the number of chronologically-ordered sample indices (x-axis) it can be seen
303 that data covers a seven-days consecutive period corresponding to the sailing time and traveled
304 distance, cf. Figures 2 and 3. Generally, reasonable agreements are found between the different
305 estimation methods (WBA vs. HC vs. ERA5), which is a finding that applies to all three sea state
306 parameters; that is, significant wave height H_s (upper plot), mean energy period T_E (middle plot),
307 and mean wave direction D_s (lower plot).

308 It is noteworthy that the commercial hindcast data (HC) has no parameters in a 5-hours period
309 around samples 182-192. While the exact reason is unknown, since the authors do not hold the
310 raw data themselves, a likely explanation could be related to the crossing of the date line. Similar
311 observations can be found in the beginning of the date stream (samples 20 to 30), where there
312 also appears to be a few values missing for the HC data provided by the commercial supplier. In
313 this case, the explanation is likely because land points are not properly treated; noting that the
314 vessel is close to land (the island of Hokkaido) during the particular time stamps.⁴ While values of
315 the wave buoy analogy are not missing at any instants, there are, however, observations of sudden
316 jumps in the data. This is primarily observed for the mean wave direction, and is likely a result of
317 modelling errors. In fact, previous reports of the wave buoy analogy have reported about problems
318 to estimate the (mean) wave direction, and it has often been found that, among the sea state
319 parameters, the largest inaccuracies are connected to the estimation of the *directional* distribution
320 of energy density (cf. Figure 4), equivalently the wave direction. The reason for this has been
321 studied and discussed by Iseki and Nielsen (2015); Hong et al. (2018, 2019), and it is considered
322 that short-term variability, due to aleatory uncertainty, in the actual wave elevation sequences is
323 responsible for the problems associated with (incorrect) estimation of the wave direction from time
324 to time. The explanation is that short-term variability severely affects the phase difference between

⁴Thanks to an anonymous reviewer to point out the problems with the date line and land points.

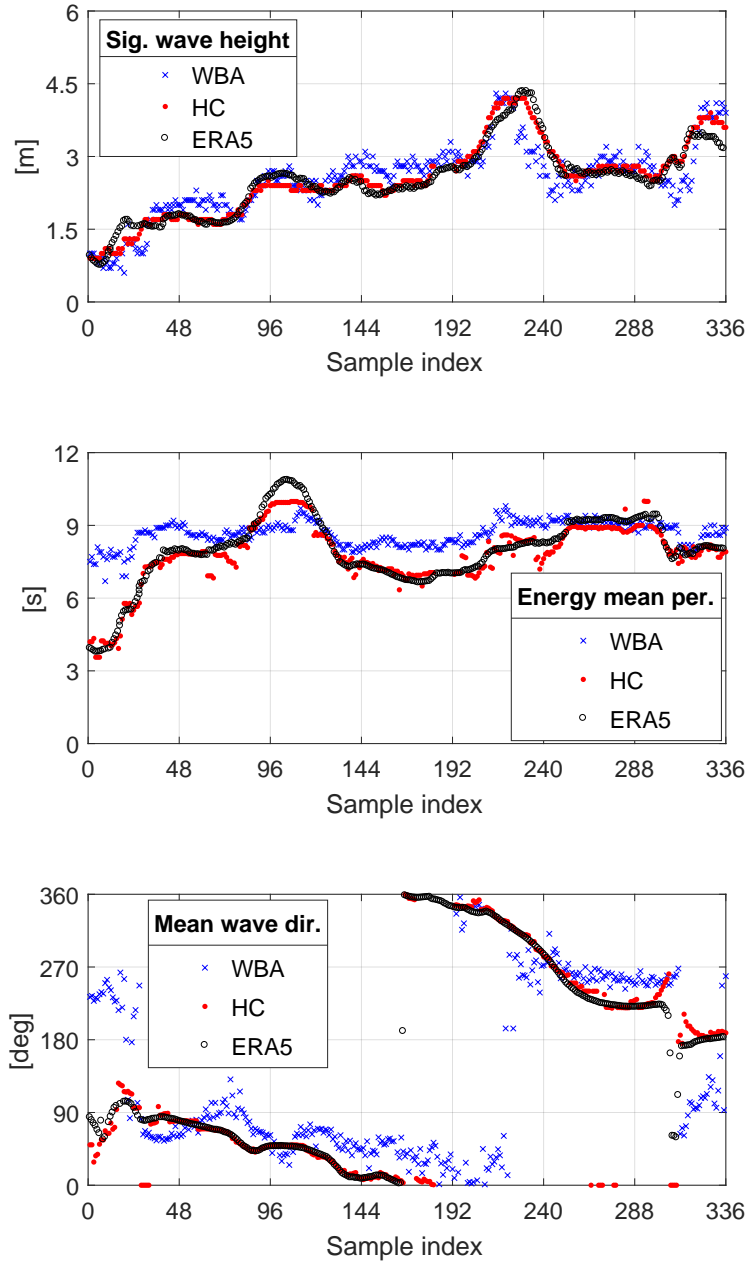


Figure 6: Estimated sea state parameters as obtained from the wave buoy analogy (WBA), the commercial hindcast (HC), and the Copernicus Climate Change Service (ERA5). The upper plot shows the result for significant wave height (H_s), the middle shows mean energy period (T_E), and the lower plot shows the mean wave direction D_s (where the waves come from).

325 the wave-induced motion components, such as heave and roll, and, thus, short-term variability is
326 harmful for the estimation of the directional distribution of energy density. This is so because
327 it is the phase difference between the motion components that gives the cross spectra used in
328 the governing equation, cf. Eq. (1), and it is in turn the cross spectra that facilitate (accurate)
329 estimation of wave direction.

330 One particular observation deserves additional comments: When the data around samples 216-
331 240 is studied, there appears to be something looking almost like a "time shift" in the estimates
332 of H_s ; comparing WBA with both sets of hindcasts. However, it is believed that the apparent
333 time shift is simply a coincidence, since it has not been possible to find any explanation related
334 to a mismatch in time/position. In addition, the "time shift" is not observable in the estimates of
335 T_E and D_s , them being neither worse nor better. On the other hand, somewhat remarkable, the
336 behaviour ("time shift") occurs as the wave system gradually changes from propagating from a
337 northerly (360 deg) to propagating from a westerly (270 deg) direction, which means that, relatively,
338 the vessel goes from being in beam sea to being in following sea, noting that the ship sails East (cf.
339 the route map in Figure 2). The change in wave direction leads to changes in the motion dynamics
340 of the vessel, and the effect(s) of 'wave filtering' by the wave buoy analogy is therefore the likely
341 cause for the particular observation resembling a "time shift" in H_s .

342 The agreement between the estimation methods is visualised in Figure 7 that shows correlation-
343 types of plots and, thus, can be used to directly evaluate the methods against each other. Not
344 surprisingly, the best agreement is observed between the two sets of hindcast data; noting that the
345 hindcast results are based on the same kind of modelling using the full *energy balance* equation
346 (Komen et al., 1994). Generally, the deviations between the two sets of hindcasts (ERA5 vs HC) are
347 small, notwithstanding it is believed that the more significant deviations are due to the calculation
348 of *equivalent* total integral wave parameters for the data by the commercial supplier, cf. subsection
349 2.4.2. The results of the wave buoy analogy agree reasonably well with the hindcast data when the
350 significant wave height is considered, and there appears to be no particular trend as the scatter is
351 random for the range of wave heights from about 0.5 m to about 4.5 m. Having a focus on the
352 wave period, i.e. T_E , it is evident that the results from the wave buoy analogy and from the sets of
353 hindcast are less consistent. Notably, it can be seen that the wave buoy analogy tends to produce
354 (too) high periods, except from a few cases around samples 85-120 (see later). This observation is

355 (again) a consequence of the fact that any ship acts as a low-pass wave filter, since the resulting
 356 wave-induced motions of a ship depend on its size relative to the wave length. In practice, this

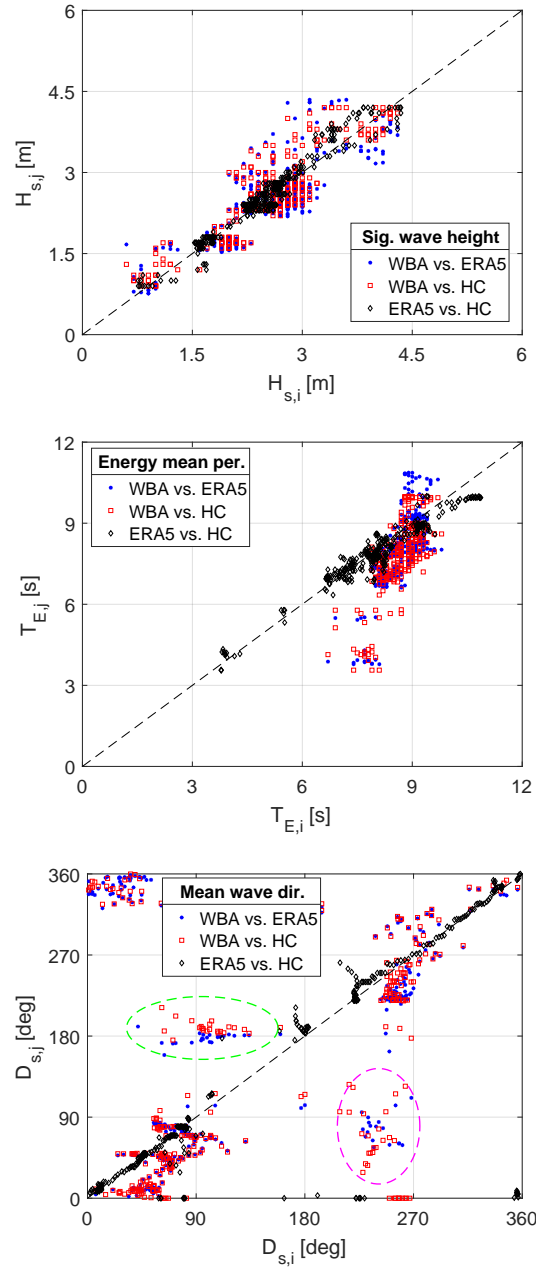


Figure 7: Direct comparisons between the estimation methods for significant wave height (top), mean energy period (middle), and mean wave direction (bottom). The pairwise comparisons are indicated by the legends, where the former estimation method, i , for a compared pair, ' i vs. j ', is given on the x-axis, while estimation method j is given on the y-axis.

357 means that, when the wave buoy analogy is applied with larger ships, there is a tendency that the
 358 higher-frequency wave components of a wave spectrum are "filtered away", and the result is that
 359 the tail of the wave spectrum is cut short.

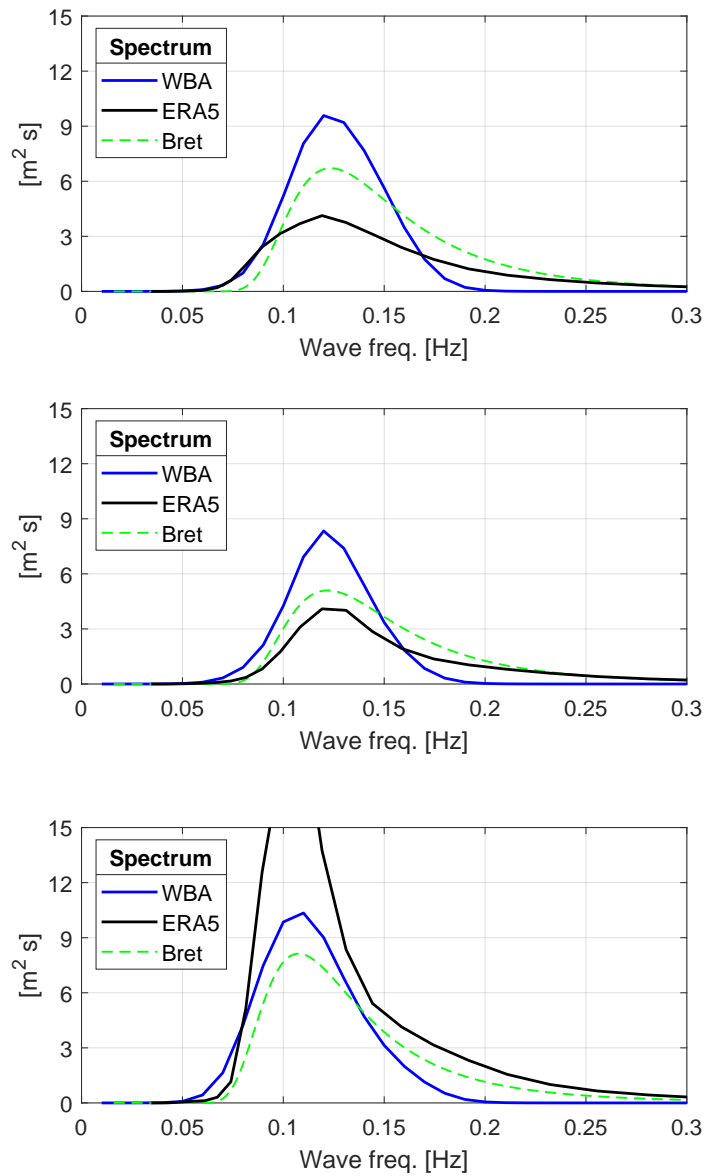


Figure 8: Examples of 1-D wave spectra to illustrate the (low-pass) filtering aspects of the wave buoy analogy compared to ERA5. The spectra correspond to samples 137, 161, and 233. The plots include comparisons to a standard spectral shape of a Bretschneider spectrum (Bret) with identical wave parameters H_s and T_p as produced by the WBA.

360 This particular finding can easily be observed from plots of (1-D) wave spectra, and Figure
361 8 illustrates a few (arbitrarily) selected cases from the considered data of this study. The plots
362 contain comparisons between WBA and ERA5⁵ shown together with a Bretschneider spectrum
363 (Beck et al., 1989) produced with identical wave parameters as estimated by the WBA. From
364 the plots in Figure 8 it is noted that the specific vessel does not really respond to waves with a
365 frequency higher than about 0.20 Hz (for any wave heading) resulting in a tail on the WBA spectra
366 which is cut short at this frequency. This "cut-off" frequency is confirmed by inspection of the
367 motion transfer functions referring to Figure 5. In the shown cases in Figure 8, the ERA5 spectra
368 are seen to have tails that match well the Bretschneider shape. It is thus an inherent problem of
369 the wave buoy analogy that it produces wave spectrum estimates where the characteristic wave
370 frequency (respectively wave period), tends to be on the lower side (respectively higher side).
371 The particular disadvantage will be the most pronounced in developing wave systems where the
372 waves are relatively short compared to vessel size. In this context it is important to mention that
373 techniques against the low-pass filtering characteristics of ships in connection with the wave buoy
374 analogy have been studied (Nielsen, 2008b; de Souza et al., 2018). The idea is to use other types
375 of responses than merely global wave-induced motions but, as this requires additional sensors not
376 installed on the specific ship of this study, no further attention is given to the topic. It is noteworthy
377 that all the shown WBA 1D wave spectra in Figure 8 have a single and distinct peak, although the
378 Bayesian technique, as indicated in Section 2, allows several peaks in the solution, frequency-wise
379 as well as directional-wise, corresponding to a mixed sea consisting of both swell(s) and wind waves
380 from multiple directions. However, in the particular cases in Figure 8, the Bayesian technique -
381 apparently - 'prescribes' a (unimodal) sea consisting of wind waves exclusively. On the other hand,
382 the middle 1D spectrum (sample 161) in Figure 8 is the integrated version of the directional wave
383 spectrum shown in Figure 4, where it can be seen that two (distinct) spectral peaks exist at different
384 directions but at the same frequency, indicating that, indeed, it is a mixed sea with waves from two
385 different directions. In fact, the two other cases of 1D spectra in Figure 8 also represent sea states
386 with waves coming from more than just one direction. Based on the preceding discussion about
387 filtering (and the possibility to estimate multi-modal wave spectra with the Bayesian technique), it

⁵It should be acknowledged that the authors were kindly supplemented the (directional) ERA5 spectra by an anonymous reviewer.

388 would obviously be interesting to study the actual wave spectra obtained from the hindcast studies,
389 and subsection 5.3 contains preliminary results in this direction. To finish the discussion about
390 T_E - and the tendency to overestimate - it should be noted that there is a sequence of samples
391 ($\sim 85 - 120$) where T_E is consistently underestimated, thus contradicting the above discussion.
392 It has not been possible to properly explain this observation, especially since there appears to be
393 noting peculiar in the estimates of the two other parameters (H_s and D_s). On the other hand,
394 the "inconsistency" coincides exactly with a period, i.e. samples 85-120, where the mean wave
395 direction, as reported by ERA5 and the HC result, initially drops a little bit and then remains to
396 be fairly constant around 45 deg, corresponding to waves coming from northeast. Having the ship's
397 course in mind (sailing eastwards), the underestimation of T_E is therefore happening in cases of
398 bow-quartering waves on port side. The data does, unfortunately, not include cases corresponding
399 to bow-quartering waves on starboard side, so it is left as a future work, by analysing data from
400 other voyages, to study if there is any relation between the underestimation of T_E and the vessel
401 being in bow-quartering waves.

402 Returning to Figure 7, the plot at the bottom shows the agreement between the wave buoy
403 analogy and the hindcast data when the mean wave direction is considered. Despite the apparent
404 scatter, the agreement is fair for most of the data, as directional ambiguity implies that wave
405 directions 0 deg and 360 deg are identical; both values represent waves propagating from North.
406 On the other hand, a mismatch is observed for the cluster of points located within the green-dashed
407 ellipse. It is seen that, for this cluster, the wave buoy analogy estimates wave directions primarily
408 in the range 90-135 deg, i.e. coming from east-southeast, while the hindcast data reports wave
409 directions in a quite narrow range around 180 deg. It is noteworthy that the particular cluster of
410 points corresponds roughly to samples 300-336, and the disagreement is seen easily also in Figure 6
411 in the bottom plot. Similarly, in Figure 7, there is a cluster of points, located within the magenta-
412 dashed ellipse, where the agreement between the wave buoy analogy and the hindcast data is poor.
413 In this case, the wave buoy analogy makes estimates of (mean) wave directions mainly in the
414 range 225-270 deg, i.e. waves coming from west-southwest, while the hindcast data reports wave
415 directions around 45-150 deg; that is, a 180 deg mismatch in some cases. In fact, for the magenta-
416 dashed ellipse and with the ship's route in mind, cf. Figure 2, the estimates by the wave buoy
417 analogy represent following to stern-quartering waves, whereas the reports from the hindcast data

Table 2: Mean value and standard deviation of the absolute errors, cf. Eq. (18), between the sea state parameters estimated by the different estimation methods, when compared pairwise.

Error	ΔH_s [m]		ΔT_E [s]		ΔD_s [deg]	
	Mean	Std	Mean	Std	Mean	Std
WBA vs. HC	0.00	0.36	0.84	0.98	3.6	59
WBA vs. ERA5	0.06	0.43	0.66	1.05	15	52
HC vs. ERA5	-0.05	0.17	0.14	0.36	-3.4	27

418 correspond to head to bow-quartering waves. The disagreement of the magenta-dashed cluster of
419 points in Figure 7 is observed as well from Figure 6 where the particular cluster of points roughly
420 corresponds to samples 1-30. The discussed disagreements in the wave heading has been indicated
421 (already) in connection with the examination of the motion transfer functions, cf. Section 4, where
422 it was reported that the vessel, according to the motion transfer functions, sees relatively little
423 response around wave headings from following to stern-quartering waves.

424 5.2. Error statistics of comparisons

The comparative study of the estimation methods has been summarised in Table 2, where the statistics of the errors between corresponding sea state parameters are shown. In the table, the mean value and the standard deviation of the absolute errors between different estimation methods i and j are presented; noting that errors are calculated between the pairs of methods 'WBA vs. HC', 'WBA vs. ERA5', and 'HC vs. ERA5'. Thus, the single error ϵ between a pair of estimates, i vs. j , for a given parameter α , for time sample k is defined by

$$\epsilon_{\alpha,k}^{\{i \text{ vs. } j\}} = \alpha_k^{\{i\}} - \alpha_k^{\{j\}}, \quad \alpha \equiv \{H_s, T_E, D_s\}, \quad k = 1 : N_t \quad (18)$$

425 The mean values and standard deviations of the error are obtained by summing up over all time
426 samples; $N_t = 336$ is the total number of 30-minutes time samples in the data (seven days). It is
427 decided to focus on absolute errors rather than normalised ones. Moreover, it is noteworthy that
428 the ambiguity in wave direction is accounted for by subtracting or adding 360 degrees if the error is
429 larger or smaller, respectively, than 180 degrees. The basis of the 'summarising numbers' in Table
430 2 is presented in Figure 9 that shows the complete sets of *normalised* errors relative to the ERA5
431 data; additional comments are given later.

432 Table 2 (and Figure 9) confirm the previous findings reported together with Figures 6 and 7.
433 Specifically, it is observed that, on average, the wave buoy analogy yields estimates of significant

434 wave height in good agreement with the two sets of hindcasts (HC and ERA5) with mean values
 435 on the errors at 0.0012 m and 0.06 m. However, the two sets of hindcast data are generally more
 436 consistent in their estimates with a smaller standard deviation on the error being 0.17 m. It is also
 437 evident from the numbers in Table 2 that the two sets of hindcast results (HC vs. ERA5) are in
 438 better agreement when estimates of mean energy periods and mean wave directions are considered
 439 as the mean values of the errors are 0.14 s and -3.4 deg, respectively, on T_E and D_s . When
 440 results from the wave buoy analogy are compared to the hindcast results on these parameters,
 441 it is interesting to observe that the mean values of the errors, i.e., WBA vs. HC, and WBA vs.
 442 ERA5, are at a reasonable level with values 0.84 s and 0.66 s on T_E and 3.6 deg and 15 deg on

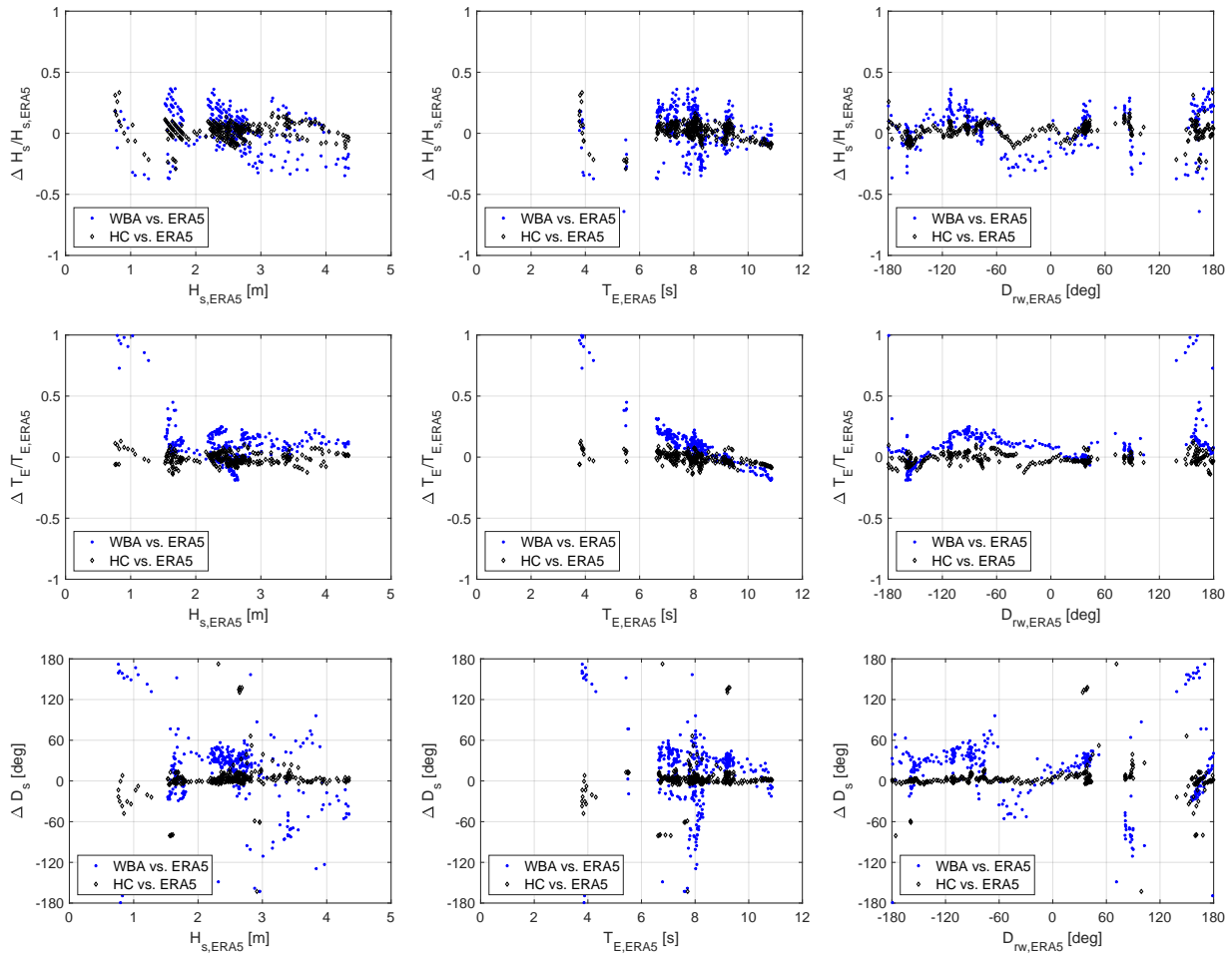


Figure 9: Illustration of the relative error between results of the wave buoy analogy and the ERA5 data (WBA vs. ERA5), and between results of the two sets of hindcasts (HC vs. ERA5). The relative wave direction $D_{rw,ERA5}$ is computed similar to Eq. (14).

443 D_s . On the other hand, the results are less consistent resulting in larger standard deviations. The
444 reason for this observation is primarily related to the two clusters of data marked by the green and
445 magenta-dashed ovals in connection with the bottom plot of Figure 7.

446 As mentioned previously, Figure 9 presents plots of the single errors used for the production
447 of Table 2. Specifically, plots are shown for the normalised errors of H_s (top row) and T_E (middle
448 row), using the ERA5 data as basis, and for the absolute errors of D_s (bottom row); in all cases as
449 function of the ERA5. The left-side column of plots shows the errors as function of the significant
450 wave height, while the middle column of plots shows the errors as function of the mean energy
451 period. The right-side column of plots shows the errors as function of the relative wave direction,
452 where 180 deg represents head waves and 0 deg is following waves, with + and - to distinguish
453 between waves approaching on the starboard side and port side, respectively. It can be seen that
454 the only notable trend, as discussed earlier, is observed for errors in the mean energy period, where
455 it is evident that the wave buoy analogy, due to filtering characteristics of the motion transfer
456 functions, cf. Section 4, produces results higher than the ERA5 data, when the wave period is low.
457 Moreover, it can be observed that cases, where WBA overestimates the most the T_E parameter,
458 correspond not only to low wave periods, but also to low significant wave heights (left middle plot).
459 Somewhat peculiar it can be seen that the wave buoy analogy and also the commercial hindcast
460 yield a mean wave energy period which is lower than the estimate from the ERA5 data for the
461 higher wave periods; or, in other words, the ERA5 mean energy period might be slightly on the
462 larger side, when the wave period is high(er).

463 5.3. Additional discussions about the effect of wave filtering

464 It is evident that wave filtering affects the estimates by the wave buoy analogy. The effect is an
465 inherent concern that results because of the motion characteristic of the given vessel, cf. section
466 4. As such, different vessels of different dimension will not imply the same "deterioration", and
467 estimates are typically better/worse depending on the relative size of the vessel compared to wave
468 length (Nielsen et al., 2019a). In general, it is difficult to evaluate exactly to which degree the
469 estimates by the wave buoy analogy are (negatively) affected by filtering. However, by use of the
470 ERA5 data it is possible to get an indication.

471 In the preceding, the integral parameters associated with the ERA5 spectra were obtained
472 *per download*, directly from the (CDS) data store, noting that the parameters are pre-computed

473 (*in* the CDS) by integrating over the full frequency range, i.e. up to 0.55 Hz, for the given the
 474 spectra (ECMWF, 2017). As an alternative, the ERA5 parameters can be obtained by limiting the
 475 integration up to 0.2 Hz; corresponding to imposing a cut-off frequency approximately equivalent
 476 to the "filtering-induced" cut-off frequency used with the WBA. The results of these computations
 477 are shown in Figure 10 with curves for the significant wave height and the mean energy period,
 478 where it is observed that the cut-off frequency, by nature, implies an increase in H_s and a decrease
 479 in T_E . It is observed that the impact seems mostly on the mean wave period, while the impact on
 480 significant wave height is less pronounced. The consequence of this finding, in relation to the wave
 481 buoy analogy, is illustrated in Figure 11. In the plots, results are included for both sets of the ERA5
 482 parameters; i.e. the normal integration over the full frequency range and the integration limited
 483 to an upper cut-off frequency of 0.2 Hz. By visual inspection, it is clear that the agreement in H_s
 484 reduces slightly ("normal" vs. "limited integration"), while the agreement in T_E improves. The
 485 agreement can be quantified by calculating normalised root mean squared (RMS) errors, making
 486 the error relative to the true ERA5 parameter obtained from integration over the full frequency
 487 range. The results are $\text{RMS}_{H_s} = 0.14$ (normal) and $\text{RMS}_{H_s}^* = 0.18$ (limited) for significant wave
 488 height, and $\text{RMS}_{T_e} = 0.15$ (normal) and $\text{RMS}_{T_e}^* = 0.09$ (limited) for mean energy period. Thus,
 489 Figure 11 supports the previous findings in the sense that the main concern with the wave buoy
 490 analogy is the, at times, inconsistent *distribution* of energy densities.

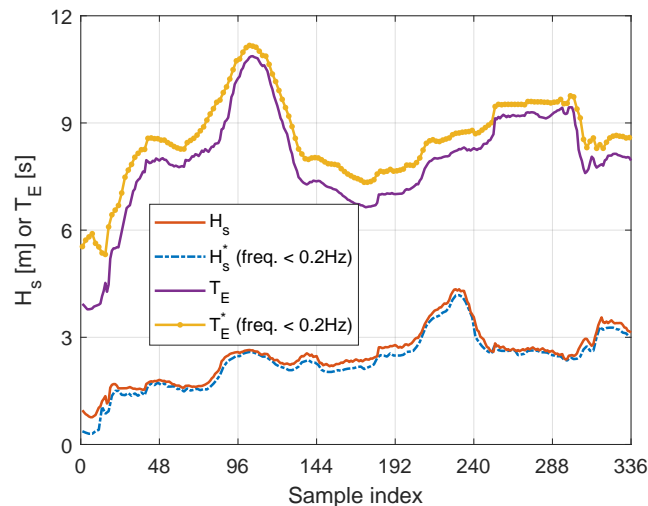


Figure 10: The effect on ERA5 H_s and T_E by imposing a cut-off frequency at 0.2 Hz.

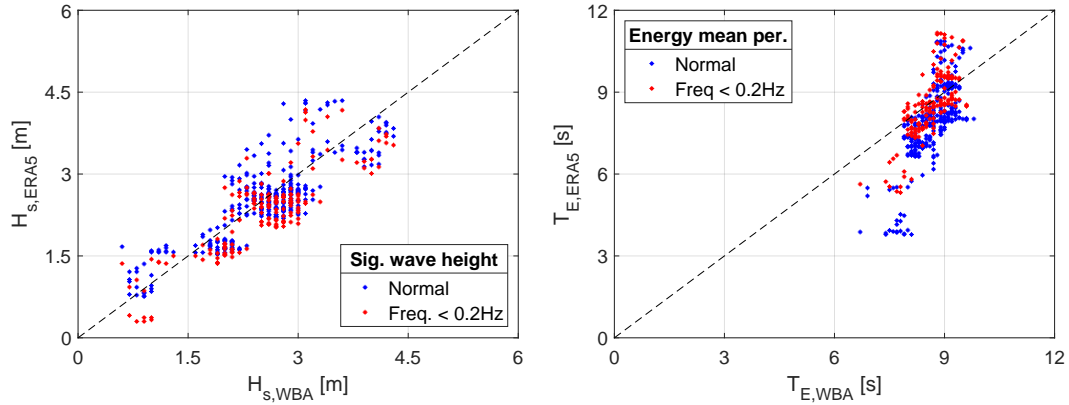


Figure 11: Comparisons between sea state parameters by the wave buoy analogy (WBA) and by ERA5 with two different sets of parameters for ERA5.

As mentioned in subsection 2.3, it will be left as a future task to make the detailed comparison between the directional wave spectra by the wave buoy analogy and the corresponding ones from ERA5. However, Figure 12 contains a few glimpses of arbitrarily selected examples; albeit the selected ones correspond to the three cases shown in Figure 8 and one additional case from the end of the data stream (sample 313). For completeness, the 1D spectra are included, although subfigures (a)-(c) are similar to the plots in Figure 8. Suffice it here to say that the (comparisons of the) directional spectra reveal an agreement of the detailed energy density distribution, somewhat in line with what can be expected for estimates *not* from the exact same physical position (nor time stamp); see further below. It is noted that the modal frequency of the directional spectra (ERA5 vs. WBA) to some degree matches; with the exception of sample 313 (subfigure d) which is an outlier in this respect. In further works, it should be attempted to study what the causes are for the observed inconsistencies. One explanation could be due to nonlinear effects (wave-ship interactions), which are not considered because of the use of (linear) transfer functions in the wave buoy analogy. However, it is also important to realise that it has not been attempted to interpolate the ERA5 spectra to the actual coordinates of the ship⁶. This means that the ERA5 wave spectra apply to the fixed grid points used in the ECMWF spectral wave model, and, therefore, the sole difference in physical location might be responsible for the (smaller) inconsistencies in the directional energy density distribution. The implication of (spatial) distance between estimates of directional wave

⁶Such an interpolation algorithm would require consideration of effects from wave dispersion, current, wind, etc.

509 spectra is also observed by, for instance, Stredulinsky (2010), where it can be seen that, albeit the
 510 mean wave direction agrees fairly well, notable differences may exist in the peak wave direction,

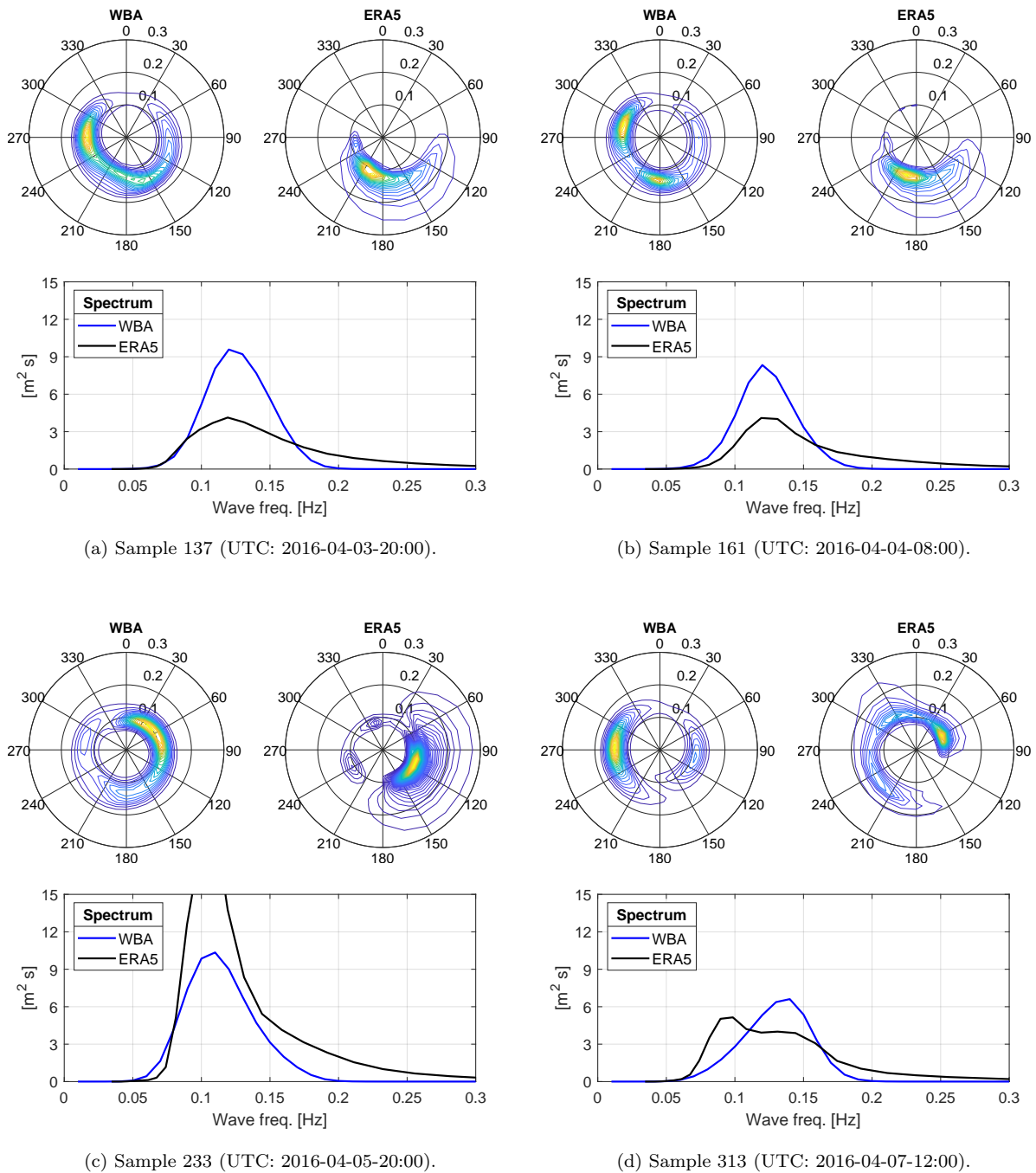


Figure 12: Examples of wave spectra with comparisons between the wave buoy analogy (WBA) and ERA5, both 2D and 1D spectra. Note, in the plots of 2D spectra, the directions are giving the directions where the energy is going.

511 when estimates by two wave buoys are compared; noting that the buoys are spaced only a few
512 kilometers (2-40 km) apart. In some of the presented cases (Stredulinsky, 2010), the two buoys
513 disagree up to 90 deg in the location of the peak.

514 *5.4. Remarks about validation of sea state estimates*

515 As indicated several times already, the comparison of ocean wave conditions is difficult, since
516 the ground truth is not available; although measurements from real wave buoys are often considered
517 close to. Nonetheless, in the preceding, the ERA5 was used as basis and, by all means, it seems
518 reasonable to assume that results of (third generation) spectral wave models are generally closer to
519 the ground truth than estimates by the wave buoy analogy. In this line, it seems relevant to mention
520 that different types of validation of the ERA5 data have been made by Hersbach et al. (2020) to
521 justify the use of ERA5 as basis. As a related comment, use of satellite data, i.e. space-borne SAR,
522 could also be interesting to look at for comparisons in future studies. In fact, satellite data was
523 used in the study by Nielsen (2006) that revealed a fair agreement. It is important, however, to
524 keep in mind that, like the wave buoy analogy, space-borne SAR cannot retrieve properly integral
525 wave parameters corresponding to the complete frequency range, since SAR observes only part of
526 the low frequency wave spectrum.

527 **6. Conclusions and final words**

528 In this paper, the wave buoy analogy has been applied to estimate the sea states encountered
529 by a container ship during a seven-days crossing of the Pacific Ocean. The basis for the estimation
530 was measurements from a motion response unit placed at a point off the centreline and close to the
531 forward perpendicular together with motion transfer functions calculated by strip theory. Thus,
532 instrumentation and necessary software are simple and inexpensive compared to the alternative
533 technology for real-time and on-site wave estimation from (in-service) ships; namely wave radar
534 systems (Nieto-Borge et al., 1999). The main conclusions of the study are listed below, and some
535 further works are suggested in addition to the future studies already discussed at various occasions
536 in the paper.

537 In the study, the sea state estimates - in terms of integral wave parameters - by the wave buoy
538 analogy were compared to two corresponding sets of estimates obtained from third generation
539 spectral wave models. The results of the spectral wave models were referred to by hindcasts,

540 and in the study the hindcasts collected from a commercial supplier and from the Copernicus
541 Climate Change Service Information were considered. Generally, good agreement exists among
542 the estimates, however, with the best correlation between the two sets of hindcast data; noting
543 that this is expected as the two sets are based on the same kind of modelling using the full energy
544 balance equation. From the quantified comparison between the wave buoy analogy and the two
545 sets of hindcast data, it was clear that the wave buoy analogy often produces directional wave
546 spectrum estimates having the same total *amount* of energy, given in terms of the significant wave
547 height (H_s), and with no particular trend. On the other hand, the estimation of the *distribution*
548 of energy (density), reflected by the mean energy period (T_E) and the mean wave direction (D_s),
549 is made with a reduced agreement, when results by the wave buoy analogy are compared to results
550 produced by hindcasts. The reason for this observation is mainly because of the filtering effect of
551 a ship; leading to a cut-off of the tail of the wave spectrum. Consequently, future studies on the
552 wave buoy analogy should try to address this problem. In this relation, existing work (Nielsen,
553 2007; de Souza et al., 2018; Nielsen et al., 2019a) focused on the use of different sensors and using
554 multiple size-varying ships could be further studied. Alternatively, in the same context, it could be
555 a possible solution to make a (mathematical) fit of the tail of the wave spectrum; emphasising that
556 the fitting should be made just for the (higher) frequencies where there generally is no significant
557 motion response in accordance with the transfer functions. This approach would be somewhat
558 consistent and similar to what is done in the spectral wave models that assume a high-frequency
559 tail above the last resolved frequency (0.55 Hz for ERA5).

560 In contrast to new developments directly focused on the wave buoy analogy, it would be of
561 interest to study how much variation the hindcast results, such as the ERA5 data, shows between
562 the single, discretised (geographical) coordinates in which the data is referred to. Herein, it is
563 understood that hindcast data, strictly speaking, is valid only in the points of the geographical
564 grid, in this study spaced by 0.5 degrees (longitudes and latitudes), while the hindcast estimate
565 at an arbitrary point, as an approximation, has been obtained by bilinear interpolation. This
566 approximation should be tested further by examining the hindcast estimates in several neighboring
567 grid points, and, in addition, by studying if the agreement between the wave buoy analogy and the
568 hindcast result(s) depends on the distance to the grid points. In fact, preliminary studies in this
569 direction have been initiated (Nielsen and Holt, 2020). As a somewhat related remark, it is noted

570 that when a ship advances with a speed of more than 20 knots, the distance covered in, say, 30
571 minutes is about 20 km. Obviously, the assumption about stationary conditions can be questioned;
572 despite a constant ship speed and heading, the seaway itself can change. It should therefore be
573 worthwhile to focus attention towards (higher order) spectral analysis methods applicable during
574 nonstationary conditions (Iseki and Terada, 2003; Iseki, 2010, 2012; Takami et al., 2020).

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